

# SEISRISK III: A Computer Program for Seismic Hazard Estimation

U.S. GEOLOGICAL SURVEY BULLETIN 1772





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**By BERNICE BENDER and DAVID M. PERKINS**

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# CONTENTS

Summary	1
Introduction	1
Modeling earthquake location uncertainty	2
Seismic hazard model for homogeneous sources	5
The fault-rupture model	6
Smoothing of accelerations calculated using fault ruptures	6
Magnitude smoothing of closest distance ruptures	7
Distance smoothing for ruptures along artificial parallel faults	8
Modeling acceleration variability	9
Calculating ground-motion exceedance probabilities	10
Model implementation—details and requirements	10
Geometry	10
Acceleration table levels	11
Source zone computations	11
Area computations	11
Grid spacing requirements for modeling uncertainty	12
Using sources with different $\sigma$ values	13
Fault computations	14
Rupture length computations	14
Magnitude smoothing	15
Attenuation function	15
Halving the magnitude intervals	15
Output options	15
Accumulating rates in successive runs	16
References cited	16
Appendix A	16
Appendix B	20
Appendix C	24

## FIGURES

1. Effect on source-zone seismicity of four different degrees of smoothing	3
2. Probabilistic accelerations for three different degrees of smoothing	4
3. Rupture spanning three fault segments	6
4. Acceleration exceedance curves calculated using three different magnitude increments	7
5. Effect of distance smoothing for traverse across three faults	8
6. Differences in acceleration levels calculated with and without attenuation variability	9
7. Four points defining new equator and seismic-felt area	11
8. Inputs for identifying quadrilaterals in a source-zone	12
9. Arcs used to approximate fraction of source zone at a given distance from site	12
10. Extension of local grid for calculating effect of future earthquake location uncertainty	13
11. Flowchart of source and fault computations	14
A1. Directions of increasing distance along fault from site	17
A2. Relocating fault segment to x-axis	18
A3. Effective local coordinates when rupture overlaps three fault segments	19



# SEISRISK III: A Computer Program for Seismic Hazard Estimation

By Bernice Bender and David M. Perkins

## SUMMARY

SEISRISK III is a revision of SEISRISK II, a computer program for seismic hazard analysis, described in Open File Report 82-293. Specifically, SEISRISK II and SEISRISK III were designed to compute maximum ground-motion levels that have a specified probability of not being exceeded during fixed time periods at each of a set of sites uniformly spaced on a two-dimensional grid. The principal differences between SEISRISK II and SEISRISK III are:

1. SEISRISK II assumes that seismicity within a seismic source zone is uniform; that is, each point within a source zone has the same probability of being the epicenter of a future earthquake. This assumption means that the projected rate of earthquakes changes abruptly at a source zone boundary; an effect of such a change in seismicity is that calculated probabilistic ground-motion levels may differ substantially at sites a few kilometers apart near a boundary. SEISRISK III retains the concept of seismic source zones but allows earthquakes within a zone to be normally rather than uniformly distributed. This allows some of the earthquakes that previously would have occurred within the zone to occur outside the zone, permitting seismicity to vary smoothly across the boundaries of the zone. The result is that calculated acceleration levels also vary more smoothly at sites near a boundary.

2. SEISRISK III does a partial "magnitude smoothing" that treats the "closest distance" ruptures as if they occurred over a range of magnitudes. In the fault-rupture model employed, ground motion (for example, acceleration) at a site resulting from a rupture along a fault is regarded as a function of closest site-to-rupture distance. This means that for the longer ruptures associated with higher magnitude earthquakes, a substantial fraction of the possible ruptures of a given magnitude along the fault may be at the same (closest) distance from the site, and the same ground-motion level will be calculated for all of these ruptures. Because the computations are performed for a discrete set of magnitudes, the ground-motion levels produced by the ruptures at closest distance for each of these magnitudes will be calculated as occurring more frequently than other ground-motion levels. This has the effect that the calculated probability of exceeding a specified ground-motion level does not vary smoothly as a function of ground-motion level. Increasing the number of magnitudes at which the calculations are performed re-

duces these irregularities but increases the computation time. By treating the closest distance ruptures for each magnitude as if they occurred over a range of magnitudes, SEISRISK III smooths the acceleration densities resulting from these ruptures.

3. The fault pattern may be quite complex within an active fault zone, and faults may be spread over a wide area. In both SEISRISK II and SEISRISK III, "artificial" parallel faults may be modeled when the presence of faulting is inferred from geologic considerations, and the orientation of the faults is assumed to be known. If a series of equidistant parallel faults are input into the program, SEISRISK III can do a partial "distance smoothing" to simulate a finer spacing between faults in order to better approximate a uniform distribution. (SEISRISK II did not permit this smoothing.)

4. Both SEISRISK II and SEISRISK III calculate probabilistic ground motions at all sites resulting from earthquakes in a single source, before proceeding to the next source. This requires retaining intermediate calculations for each site to accumulate ground-motions from successive sources. SEISRISK II wrote these intermediate results onto a disk; SEISRISK III saves the results in two-dimensional arrays in memory. Currently, the program requires two arrays dimensioned  $(56, N)$ , where  $N$ =number of sites on grid of sites for which calculations are performed.

## INTRODUCTION

In most seismic hazard analyses, earthquakes are modeled as

1. points located randomly within seismically homogeneous areal source zones; and (or)
2. finite length ruptures that occur randomly along linear fault segments.

Probabilistic models are used because sizes and locations of future earthquakes cannot be predicted precisely. Source zones are typically regarded as seismically homogeneous, that is, each point within a source zone is assumed to have the same probability of being the epicenter of a future earthquake, because the analyst has

no reason to assume that an earthquake is more likely to occur at one point than at another within the zone.

The analyst uses catalogs of recorded earthquakes, and whatever geologic, geophysical or other information is available to help him to determine seismic source zones. However, as Thenhaus (1983, p.31) noted, "The procedures used in delineating seismic source zones are ill defined. No single standard exists by which source zones across the nation can be drawn, primarily because of the nonuniform level of pertinent seismological, geological and geophysical information available for areas of vastly differing tectonic and geologic settings. The equivocal association of seismicity with geologic structure throughout most of the United States compounds the problem." Because the evaluation and interpretation of the various types of information available depends strongly on individual judgment or opinion, different experts may define source zones differently.

Assuming that seismicity is uniform within a source zone means that seismicity changes abruptly at each source-zone boundary. An effect of these abrupt changes in seismicity is that probabilistic ground-motion levels calculated at sites a few kilometers apart near a boundary may differ significantly. Bender (1986) noted changes of 50-80 percent or more in acceleration levels calculated for longer exposure times at sites 20 km apart near a boundary of a source zone. Such changes seem unreasonable if, in fact, source zones are not well determined.

To avoid sudden changes in seismicity at source-zone boundaries, SEISRISK III, unlike SEISRISK II, permits an option of "earthquake location uncertainty". Source-zone boundaries are specified in SEISRISK III as in SEISRISK II, but now each point within the zone is regarded as the mean or most likely location of a future earthquake; the locations of actual earthquakes are normally distributed with standard deviation  $\sigma$  about their mean locations ( $\sigma$  is an input parameter). This permits earthquake rates to vary smoothly near a boundary, without being significantly different in the center of the region from the rates assumed in the uniform-seismicity model (fig.1). As  $\sigma$  increases, a higher percentage of the earthquakes that were expected to occur within the source zone when no location uncertainty ( $\sigma=0$ ) is assumed, occur instead beyond the boundaries.

In the following sections, we discuss in more detail earthquake location uncertainty and how this uncertainty is implemented in SEISRISK III. We also discuss "magnitude smoothing" and "distance smoothing" for ground-motions resulting from ruptures along a fault, and the parameters that control this smoothing in the computer program. A listing of SEISRISK III (Appendix C) and information regarding data input and formats (Appendix B) are also included.

## TERMINOLOGY

In this report, we use acceleration as a generic measure of ground shaking; we could equally well have used velocity or any other measure. By acceleration occurrences in an interval  $a_{j-1} < a \leq a_j$  we mean the average annual occurrence rate of accelerations in the range  $a_{j-1} < a \leq a_j$  experienced at a specific site. Acceleration levels or probabilistic acceleration levels refer to the accelerations that are calculated as having a specified probability of not being exceeded during a given time period at particular sites.

## MODELING EARTHQUAKE LOCATION UNCERTAINTY

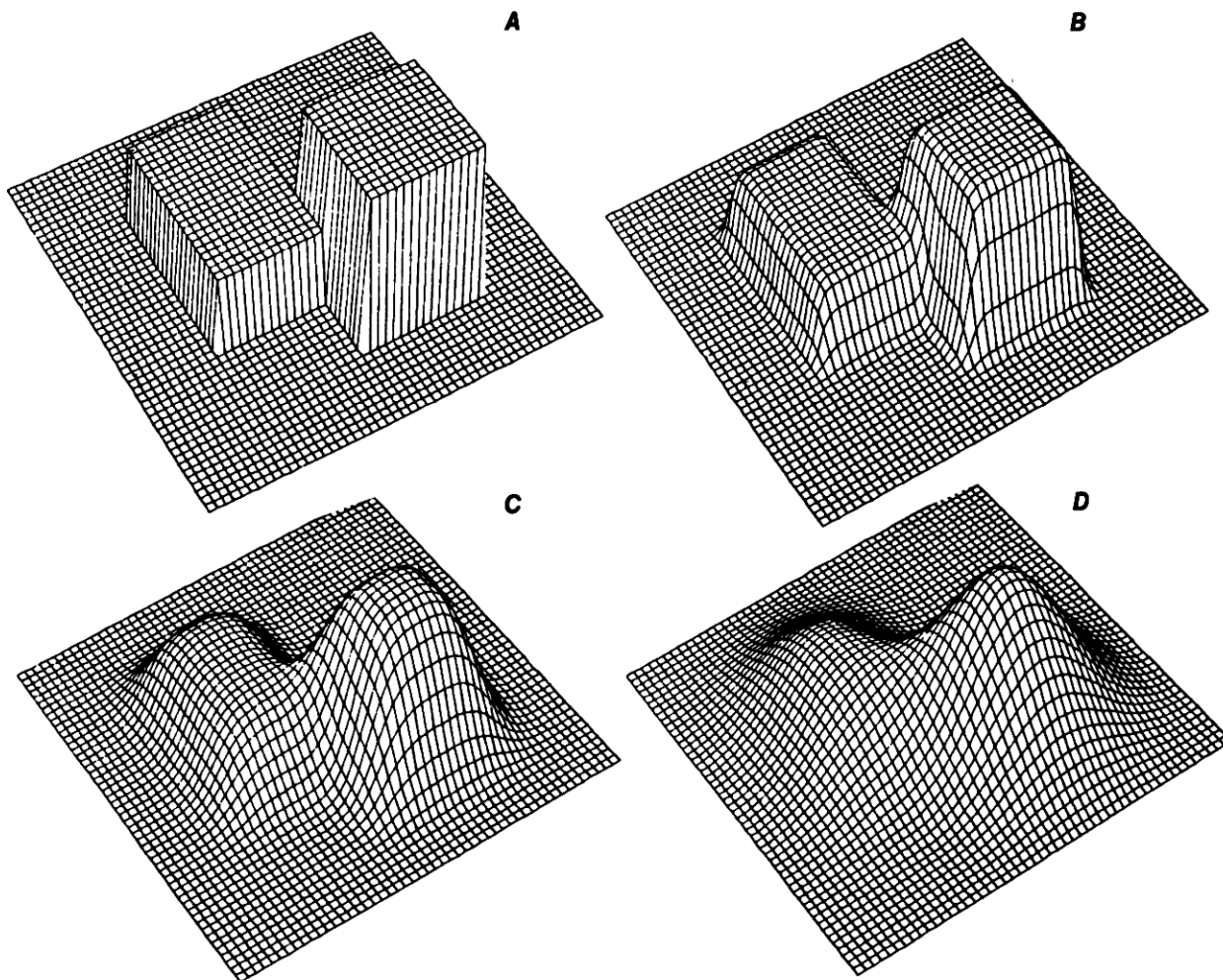
In the model implemented in SEISRISK III, the seismicity associated with each point in a source zone is regarded as normally distributed about that point. The possible earthquakes, rather than being uniformly distributed throughout a source region, have mean locations that are uniformly distributed, with scatter about each mean location given by a circular-normal distribution. Stated more precisely, in this model, if  $\sigma$  is the standard deviation in location, and if the expected (mean) location of an earthquake is at  $(X_e, Y_e)$ , the probability that the earthquake will occur instead in a small area  $A$  surrounding  $(X_e + \Delta x, Y_e + \Delta y)$  is given by

$$p_A(\Delta x, \Delta y) = \frac{A}{2\pi\sigma^2} \exp\left(-\frac{\Delta x^2 + \Delta y^2}{2\sigma^2}\right). \quad (1)$$

Figure 2 (A-C) shows acceleration levels calculated as having a 90-percent probability of not being exceeded during exposure times of 10, 50, 250, and 1000 years when no variability in earthquake location ( $\sigma = 0$ ) is assumed, and when  $\sigma = 10$  and 20 km (the attenuation function of Joyner and Boore (1981) with  $\sigma_a = 0.5$  in  $\log_e$  acceleration was used). The sites are located 5 km apart on a line through the centers of two contiguous source zones. Each zone is a square 60 km wide; the second zone has an earthquake rate 10 times higher than the first. Unless  $\sigma$  is a substantial fraction of the length or width of the source zone (for example,  $\sigma = 20$  km in fig.2), the acceleration levels calculated at sites near the center and at sites on the boundaries of the source zone are nearly the same for various values of  $\sigma$ . For a site outside the source, if  $\sigma \neq 0$  is assumed, some of the earthquakes will occur closer to the site than when  $\sigma = 0$  is assumed, and these closer earthquakes will produce higher accelerations at the site than will any earthquakes of the same magnitude occurring within the fixed source zone.

Figure 2D shows the difference in acceleration level calculated for the 1000-year exposure time at sites 5





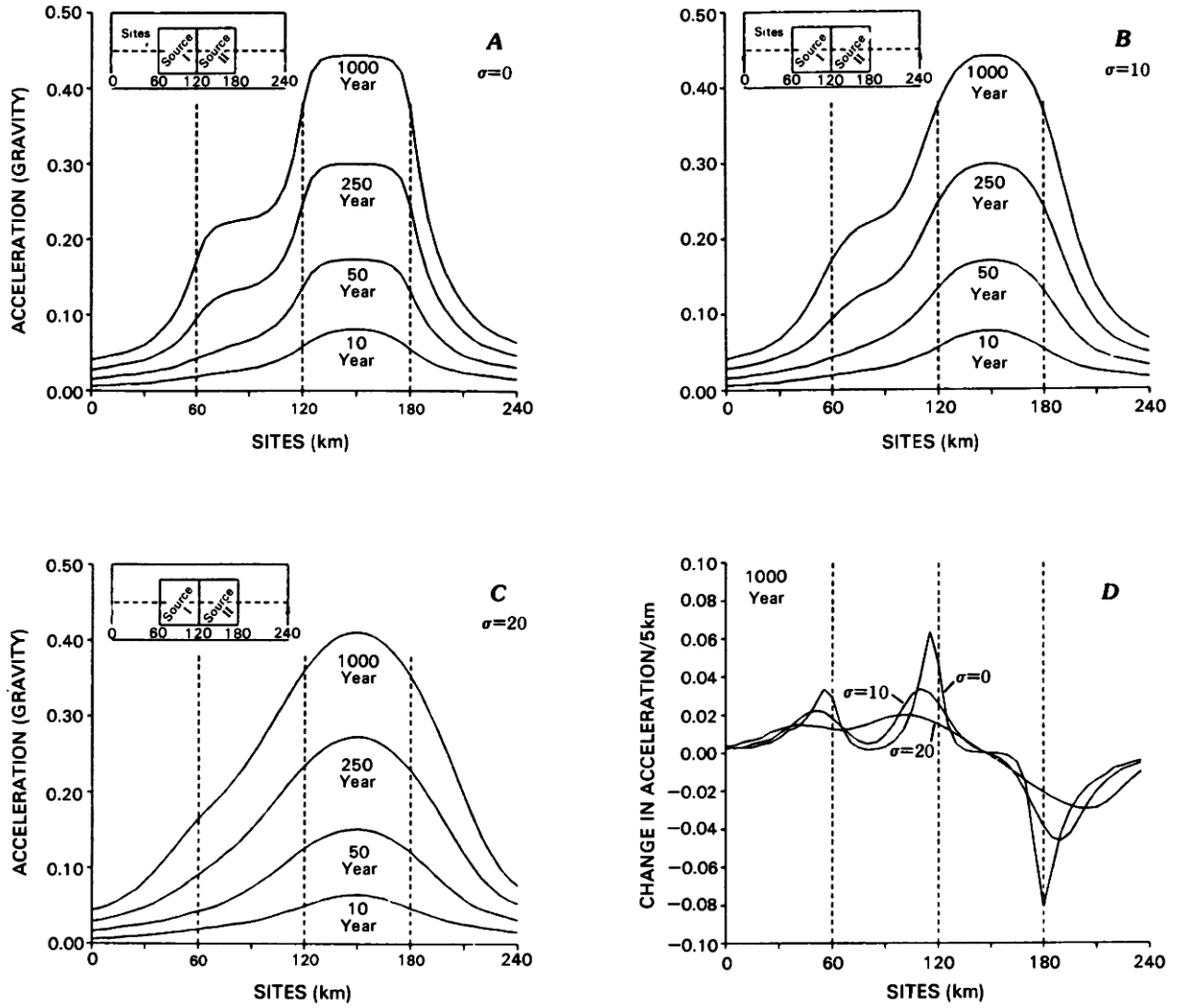
**Figure 1.** Earthquake rates within gridded areas (10 km x 10 km). In *A* earthquakes occur only within two well-defined seismically homogeneous source zones. In *B*, earthquake locations are normally distributed with standard deviation  $\sigma = 10$  km; in *C*,  $\sigma = 25$  km; in *D*,  $\sigma = 50$  km.

km apart on the line when  $\sigma = 0, 10$ , and  $20$  km; for example, a change in acceleration of  $+0.05$  g at the site at  $X$  km means that the acceleration level calculated at the site at  $X+5$  km is  $0.05$  g higher than the level calculated at  $X$ .

To explain how acceleration levels are calculated in SEISRISK III when earthquake-location variability is taken into account, we first observe that, if earthquakes are normally distributed, those that occur at distance  $(\Delta x, \Delta y)$  from their mean locations, in effect, create a new source zone translated by  $(\Delta x, \Delta y)$  from the original zone. This means that the possible locations of the source zone (rather than of the earthquakes) may be

regarded as having a normal distribution. More explicitly, the source zone is said to be located at  $(\Delta x, \Delta y)$  if an arbitrary point in the zone that would have been at  $(X_b, Y_b)$  if the zone were at its mean or most likely location, lies instead at  $(X_b + \Delta x, Y_b + \Delta y)$ . The probability that the zone will be located at  $(\Delta x, \Delta y)$ , is proportional to  $\exp\left(-\frac{\Delta x^2 + \Delta y^2}{2\sigma^2}\right)$ , where  $\sigma$  is now the

standard deviation in source location. Furthermore, the acceleration occurrences in an interval  $a_{j-1} < a \leq a_j$  that are calculated for a site at  $(X, Y)$  when the source location is moved to  $(\Delta x, \Delta y)$  are the same as those calculated for a site at  $(X - \Delta x, Y - \Delta y)$  when the source is at its original location. This means that the effect of source-zone location uncertainty at a site at  $(X, Y)$  can be produced by computing a weighted average of the acceleration occurrences at various sites, calculated using the original seismically homogeneous source zone. This last fact was used to implement earthquake location (or, equivalently, source-zone location) uncertainty in SEISRISK III as follows.



**Figure 2.** (A)-(C): Accelerations calculated to have a 90 percent probability of not being exceeded during various time periods, at sites on a line through the centers of two seismic sources. Source II has an earthquake rate 10 times higher than Source I. Earthquake locations are normally distributed with A,  $\sigma=0$  km; B,  $\sigma=10$  km; C,  $\sigma=20$  km. D, Differences in acceleration levels calculated above for 1000-yr period at sites 5 km apart on the line of sites.

SEISRISK III begins, as does SEISRISK II, by calculating expected annual accelerations at each site on a grid of sites, resulting from earthquakes in well-defined, seismically homogeneous zones. Sites are spaced at increments  $dx, dy = dx$ . The acceleration occurrences at a site at  $(X_1, Y_1)$  that result when earthquake location variability is taken into account is a weighted sum of the acceleration occurrences that were calculated at each point on the grid, assuming homogeneous

source zones. The weight assigned to acceleration occurrences in the interval  $a_{j-1} < a \leq a_j$  at the point at  $(X_1 + k_1 dx, Y_1 + k_2 dy)$  is proportional to

$$w(k_1, k_2) = \frac{dx dy}{2\pi\sigma^2} \exp\left(-\frac{k_1^2 dx^2 + k_2^2 dy^2}{2\sigma^2}\right). \quad (2)$$

The occurrences in the interval  $a_{j-1} < a \leq a_j$  that result when earthquake location variability is taken into account is given by

$$D = \frac{\sum_{k_1} \sum_{k_2} w(k_1, k_2) \rho[(X_1 + k_1 dx, Y_1 + k_2 dy), j]}{\sum_{k_1} \sum_{k_2} w(k_1, k_2)} \quad (3)$$

where  $\rho[(X_1 + k_1 dx, Y_1 + k_2 dy), j]$  = the acceleration occurrences calculated in the interval  $a_{j-1} < a \leq a_j$  at the site at  $X_1 + k_1 dx, Y_1 + k_2 dy$  assuming no earthquake location variability (i.e., homogeneous source zones).

## SEISMIC HAZARD MODEL FOR HOMOGENEOUS SOURCES

To determine acceleration rates in SEISRISK III that result when earthquake location variability is taken into account, we first compute acceleration rates, as in SEISRISK II, assuming that earthquakes occur randomly within seismically homogeneous source zones or along linear fault segments. In this section, we briefly describe the model before incorporating earthquake location uncertainty.

Earthquake occurrences are assumed to have a Poisson distribution, and rates that remain constant during the time periods of interest. Mean or median ground-motion from an earthquake is an increasing function of magnitude and decreasing function of site-to-source distance. Earthquakes occur as points within quadrilateral source zones or as finite-length ruptures along linear fault segments. Ruptures of a given length may, with equal likelihood, be centered at any point on the fault for which neither end of the rupture would extend beyond the end of the fault.

Earthquakes in a given zone or fault are restricted to occur within a specified magnitude range; the range may be different for different zones and faults. More specifically, for each zone (and fault) a maximum magnitude  $m_{max}$ , a minimum magnitude  $m_0$ , and a magnitude interval  $\Delta m$  are assumed. All magnitudes occurring within a magnitude interval are grouped at the center of the interval, that is, at a set of  $n$  distinct magnitudes  $m_j$ :

$$m_j = m_0 + (j+1/2)\Delta m, \quad 0 \leq j \leq n-1, \\ \Delta m = \frac{m_{max} - m_0}{n}. \quad (4)$$

The rate of earthquakes per unit time in the zone (or on the fault) in each magnitude interval is specified by the program user; it frequently follows (but is not required to follow) a relationship of the form

$$\log N_m = a - bm \quad (5)$$

where  $a$  and  $b$  are constants. If  $N_{m(j)}$  is the total number of earthquakes in the  $j^{th}$  magnitude interval, the fractional occurrences expected in any small subarea  $\Delta A$  is

$$\frac{\Delta A}{A} N_{m(j)} \quad (6)$$

where  $A$  = total area of zone.

Seismicity is assumed to remain constant during the time periods being considered; that is, the average rate of earthquakes per unit time for each magnitude interval does not change with time.

A table of median (or mean) ground-motion values (for instance, acceleration) for the desired attenuation function is entered into the program for a set of earthquake magnitudes and distances. The program assumes that ground-motion increases with magnitude and decreases with site-to-source distance; ground motions for various magnitudes and distances are obtained as needed by interpolating values in the table.

Given the source zones, earthquake rates and attenuation function, an average yearly rate of accelerations within each of 55 acceleration intervals is constructed for each site. (In SEISRISK II 100 intervals were used; the number was decreased to save time and space, but can be adjusted within the program.) Acceleration occurrences in the range  $a_{i-1} < a \leq a_i$  are accumulated into  $reg(i)$ , the accumulator for the  $i^{th}$  interval, as described below.

For the given attenuation function, let  $d_m(i)$  be the distance at which a magnitude  $m$  earthquake produces median acceleration  $a_i$  at a site. The probability that an acceleration in the range  $a_{i-1} < a \leq a_i$  occurs at the site is the probability that a magnitude  $m$  earthquake occurs at a distance  $d_m(i) \leq d < d_m(i-1)$  integrated (summed) over all permissible magnitudes. (Note  $d_m(i) \leq d_m(i-1)$  because shorter distances produce higher accelerations.) The probability that an earthquake of magnitude  $m$  occurring within the zone (or on the fault) occurs within the distance interval  $d_m(i) \leq d < d_m(i-1)$  is equivalent to the fraction of the area of the zone (or length of the fault) that lies within this interval. The area (or length) within the interval

$$A[m(i)] = f[\bar{d}_m(i), \Delta d_m(i)], \quad (7)$$

is a function of interval width

$$\Delta d(i) = d_m(i-1) - d_m(i) \quad (8)$$

and interval midpoint

$$\bar{d}_m(i) = \frac{d_m(i-1) + d_m(i)}{2}. \quad (9)$$

The rate of accelerations in the range  $a_{i-1} < a \leq a_i$  at the site resulting from magnitude  $m$  earthquakes within the zone then is equal to the rate of magnitude  $m$  earthquakes within  $A[m(i)]$ :

$$\rho_m(i) = \frac{A[m(i)]}{A} N_m \quad (10)$$

where  $N_m$  = rate of magnitude  $m$  earthquakes for the entire zone or fault

$A$  = total area of zone (or total fault length).



The program determines  $A[m(i)]$  for each magnitude and adds the corresponding rate  $\rho_m(i)$  into the  $i^{th}$  accumulator,  $reg(i)$ . The total rate accumulated for accelerations in the range  $a_{i-1} < a \leq a_i$  from one source zone then is the sum over magnitude of

$$\rho(i) = \sum_{j=1}^n \rho_{m_j}(i) \quad (11)$$

where  $n$ =total number of magnitude intervals.

The process is repeated for all  $1 \leq i \leq 55$ , for all sources.

These calculations assume a single value of ground motion (acceleration) for earthquakes of a given magnitude and distance. Calculations that take into account ground motion variability will be discussed later. First, we shall consider the calculation of (median) accelerations resulting from ruptures along faults and some ways of smoothing the calculated acceleration densities.

## THE FAULT-RUPTURE MODEL

As previously noted, both SEISRISK II and SEISRISK III allow earthquakes to be modeled as finite-length ruptures along linear fault segments. In both programs, acceleration occurrences resulting from fault ruptures are added to those resulting from earthquakes in source areas. However, unlike SEISRISK II, SEISRISK III permits a partial "smoothing" of acceleration occurrences at a site caused by ruptures along a fault. We shall discuss the fault-rupture model and the changes that have been made to smooth those occurrences that result from fault ruptures.

In the fault-rupture model, sets of one or more faults are defined as fault zones; if a fault zone contains more than one fault, earthquakes in the zone are distributed among the faults in proportion to the lengths of the faults. An individual fault of length  $L$  may consist of as many as 24 connected straight-line segments. (The maximum number of segments may be altered within the program.) A rupture may span several segments (fig.3). Ruptures occur with equal likelihood anywhere along the fault, so long as they are wholly contained within the fault. Rupture lengths are a function of magnitude and are lognormally distributed with median length  $l(m)$  and standard deviation  $\sigma_l$ .

As with source areas, the ground-motion at a site is a function of earthquake magnitude and distance. However, in the case of faults, distance from the site to a rupture is always taken to be the distance to the point on the rupture closest to the site. (The fault-rupture model is described in detail in Bender, 1984b.) All magnitude  $m$  ruptures along the fault whose closest distance

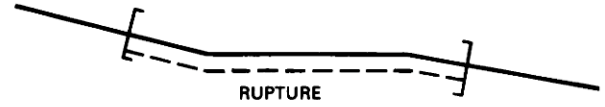


Figure 3. Fault consisting of three connected straight-line segments. A rupture may span several segments. Closest distance from rupture to site is used to determine ground motion at the site.

$d$  to the site is in the range  $d_m(j) \leq d < d_m(j+1)$  will produce acceleration  $a$ ,  $a_{j-1} < a \leq a_j$ , at the site. The calculations to determine site-to-rupture distance and the fraction of ruptures of each magnitude within each distance range are described in Appendix A.

## Smoothing of Accelerations Calculated Using Fault Ruptures

In the fault-rupture model, a substantial fraction of the ruptures of earthquakes of a given magnitude may occur at the same closest distance from the site and may produce the same median acceleration at the site. For example, assume the fault extends on the x-axis from  $0 \leq x \leq L$ . A rupture of length  $L/4$  may be centered at any point  $x$  along the fault such that  $1/8L < x \leq 7/8L$ . If the site is located on a line perpendicular to the center of the fault at  $(L/2, P)$ , all ruptures that have their centers in the range  $1/4L \leq x \leq 3/4L$  along the fault will be at the same (closest) distance  $P$  from the site. This means that two-thirds of the possible ruptures of length  $L/4$  will be at distance  $P$  from the site; for rupture length  $L/2$ , all possible ruptures will be at distance  $P$  from the site. Because longer ruptures are associated with higher magnitude earthquakes, the fraction of the ruptures that will be at closest distance  $P$  from the site increases with magnitude.

Because calculations are normally done only at a set of magnitudes  $m_i$  spaced  $\Delta m$  apart, ruptures of magnitude  $m_i$  earthquakes at distance  $P$  from the site will cause a "spike" of acceleration occurrences to be accumulated in some  $reg(j)$ . (Recall that occurrences of accelerations in an interval  $a_{j-1} < a \leq a_j$  are accumulated in a corresponding accumulator  $reg(j)$ .) The number of spikes in the acceleration occurrences at a site depends on the number of magnitudes used; the locations and sizes of the spikes depend upon the actual magnitude values. If no acceleration variability is assumed (for example, mean or median acceleration values only are used), or if  $\sigma_a$ , the standard deviation in log acceleration, is small, spikes in acceleration occurrences will result in sharp changes in the accelera-



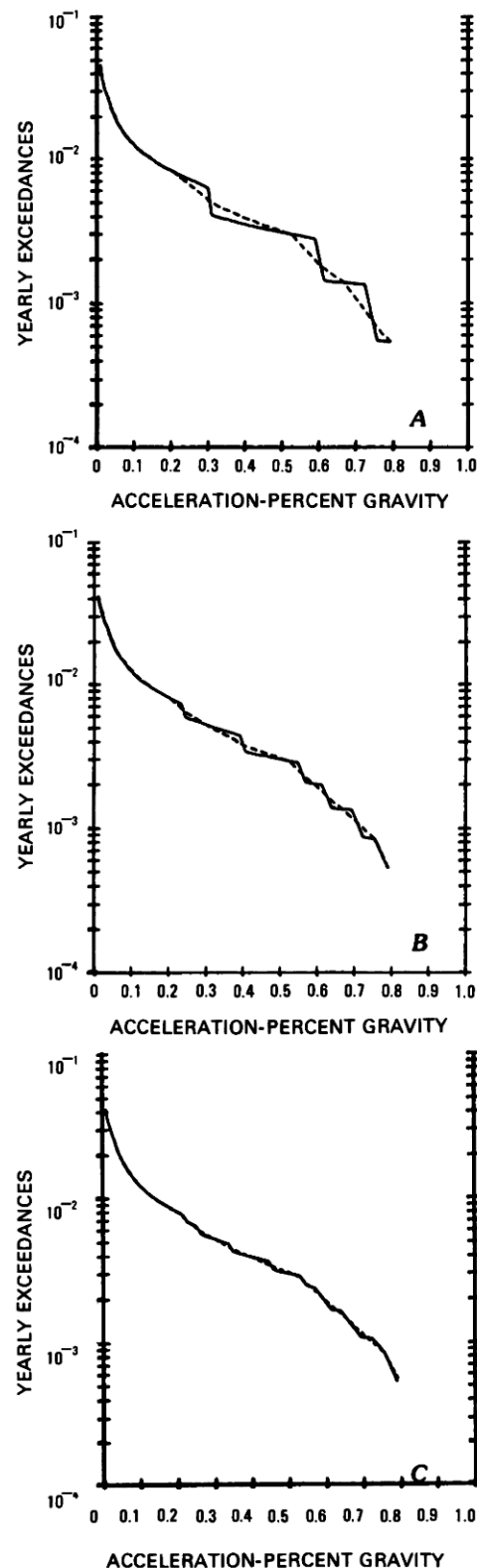
tion exceedance curves at the accelerations at which the spikes are present. Furthermore, these spikes may occur at different accelerations at different sites, causing the probabilistic accelerations calculated at different sites to vary irregularly. SEISRISK III does a partial smoothing of acceleration occurrences resulting from the ruptures at distance  $P$ , as described below.

### Magnitude Smoothing for Closest Distance Ruptures

Earthquakes at a fixed distance  $P$  from the site, with magnitudes in the range  $m(j-1) \leq m < m(j)$  will produce median accelerations in the corresponding interval  $a_{j-1} \leq a < a_j$ . If earthquakes within the magnitude interval  $m_i - \Delta m/2 \leq m_i \leq m_i + \Delta m/2$  are assumed to follow a Gutenberg-Richter magnitude-frequency relationship, the number  $h(j)$  of those earthquakes with magnitudes that lie between  $m(j-1) \leq m \leq m(j)$  can be determined ( $h(j)$  contains magnitudes in the range  $\max[m_i - \Delta m/2, m(j-1)] \leq m \leq \min[m(j), m_i + \Delta m/2]$ ). If  $f(i)$  is the fraction of ruptures of length  $l_i(m)$  of magnitude  $m_i$  earthquakes that occur at distance  $P$  from the site, approximately the same fraction of earthquakes will be at distance  $P$  for all magnitudes in the range  $m_i - \Delta m/2 \leq m_i \leq m_i + \Delta m/2$ , for small values of  $\Delta m$ . This means that a rate of accelerations  $r(i, j) = f(i)h(j)$  resulting from ruptures at distance  $P$  for magnitudes  $m_i - \Delta m/2 \leq m_i \leq m_i + \Delta m/2$  can be assigned to the acceleration interval  $a_{j-1} \leq a < a_j$ .

The solid lines in figure 4 show the yearly rate of exceedances of various accelerations calculated at a site at a perpendicular distance of 1 km from the center of a 300 km long fault, when the magnitude range is  $5.0 \leq m \leq 8.0$  and the magnitude increments are  $\Delta m = 0.6$ , 0.3, and 0.15, respectively. The dashed lines in figure 4 were obtained in each case by magnitude smoothing of the spikes of acceleration occurrences. Very little difference appears in the smoothed results for  $\Delta m = 0.6$ , 0.3 and 0.15, suggesting that for sites near the center of the fault a magnitude increment of  $\Delta m = 0.6$  might suffice if the spikes are smoothed in this manner.

**Figure 4.** Yearly acceleration exceedance rates calculated at a site at a perpendicular distance of 1 km from the center of a 300-km-long fault, using magnitude increments A.  $\Delta m = 0.6$ ; B.  $\Delta m = 0.3$ ; C.  $\Delta m = 0.15$ . The magnitude range is  $5.0 \leq m \leq 8.0$ . Attenuation curves are those of Schnabel and Seed, (1973). The dashed line in each case show the exceedances calculated when the acceleration density spikes resulting from ruptures that intersect the point on the fault closest to the site are smoothed for each magnitude.

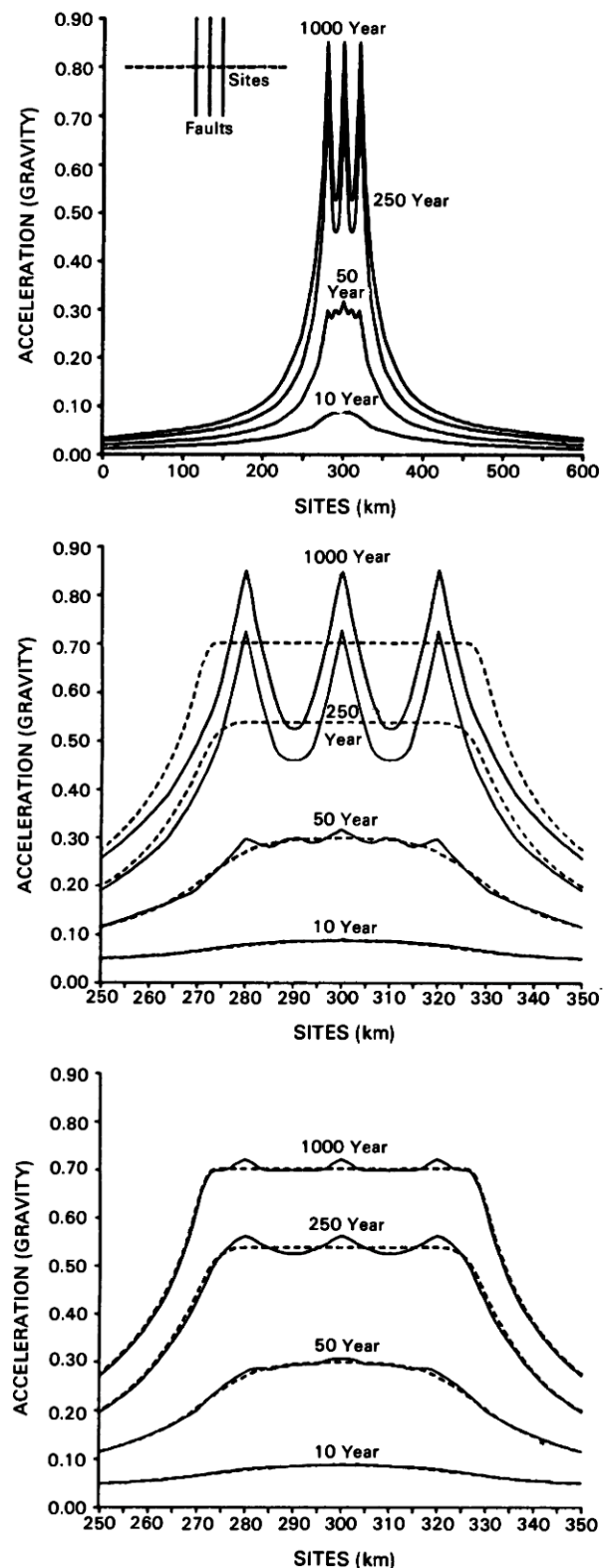


In smoothing acceleration occurrences, occurrences normally assigned to magnitude  $m_{max} - \Delta m/2$  are distributed throughout the interval  $m_{max} - \Delta m/2 \leq m \leq m_{max} + \Delta m/2$ ; in particular, some of the occurrences are assigned to the interval containing the acceleration produced at distance  $P$  by the maximum magnitude,  $m_{max}$ . However, when magnitude smoothing is not done, the highest magnitude used in the calculations continues to be  $m_{max} - \Delta m/2$ , (equation 4). For a site beyond the end of the fault at  $(X, P)$  (either  $X < 0$  or  $X > L$ ), there are no ruptures at distance  $P$ , and, hence, no acceleration occurrence spikes from ruptures at closest distance; therefore, no magnitude smoothing is done for acceleration occurrences at such a site. This implies, in effect, a lower maximum magnitude for sites beyond the end of the fault.

### Distance Smoothing for Ruptures Along Artificial Parallel Faults

The fault pattern may be quite complex within an active fault zone, and faults may be spread over a wide area. SEISRISK III allows one to model artificial parallel faults when the presence of faulting is inferred from geologic considerations and the orientation of the faults is assumed to be known. In this model, faults with a specified orientation are assumed to be equally likely anywhere within a given area; for example, faults parallel to the x-axis are equally likely to have their y-coordinate anywhere in the range  $0 \leq y \leq w$ . (The y-coordinate of the fault is uniformly distributed in this model.) To approximate a uniform distribution of faults, coordinates of evenly spaced parallel faults must be entered into the computer program. However, unless the spacing between faults is quite small (for example, 1 or 2 km), or unless some sort of smoothing is done, the acceleration levels calculated at sites a short distance apart can be quite different. Figure 5 illustrates how calculated acceleration levels may change as a function

Figure 5. Accelerations calculated to have a 90 percent probability of not being exceeded during various time periods at sites on a line perpendicular to the centers of three parallel faults, A, 20 km apart. B, the scale is expanded and the dashed lines show values obtained when faults are placed instead at increments of 1 km over a distance of 60 km and the original seismicity is divided evenly among the faults. C, the solid lines show the results when 20-km fault spacing is used in the calculations, but ruptures on each fault that occur "as close as possible" to the site are placed instead on a line perpendicular to the original faults; the dashed lines are the same as in B. The Schnabel and Seed (1973) attenuation curves are used.



of distance from a fault. Figure 5A shows levels calculated at sites on a line perpendicular to the centers of three parallel faults, 20 km apart. In figure 5B the scale was expanded; the solid lines again show acceleration levels for the three fixed fault locations; the dashed lines give results obtained when faults were placed 1 km apart and the seismicity adjusted accordingly.

A distance smoothing is desirable when a single fault is used to represent a uniform distribution of faults. For example, assume a fault that extends from  $(0, 0)$  to  $(L, 0)$  was entered into the program to represent faults with  $y$  coordinate anywhere in the range  $-y_0 \leq y \leq y_0$ . Accelerations are calculated only for ruptures along the fault at  $y=0$ . If the site is located "above" the fault at  $(X, P)$  ( $0 \leq X \leq L$ ), an acceleration density spike will result from ruptures at distance  $P$ . To smooth the accelerations resulting from these ruptures, assume that any rupture with end points  $(x_1, 0)$ ,  $(x_2, 0)$  along the original fault, at distance  $P$  from the site, may instead have end points  $(x_1, y)$ ,  $(x_2, y)$  on any fault such that  $-y_0 \leq y \leq y_0$ . The seismicity associated with such a rupture along a fault at  $y$  concentrated at the point on the fault closest to the site at  $(X, y)$ . This means that if faults [with end points at  $(0, y)$ ,  $(L, y)$ ,  $-y_0 \leq y \leq y_0$ ], are uniformly distributed parallel to the original fault, ruptures at closest distance on each fault will produce the same accelerations as points uniformly distributed on a line through the site perpendicular to, and extending  $\pm y_0$  km from, the original fault. Acceleration occurrences in each interval  $a_{i-1} \leq a \leq a_i$  may be determined from the frequency of the "point-source" earthquakes of each magnitude at the appropriate distances along this perpendicular line. In figure 6C, the ruptures at closest distance to each site were treated as point-source earthquakes along a line through the site and perpendicular to the faults; the calculation was done using  $\Delta m = 0.6$ .

## MODELING ACCELERATION VARIABILITY

The ground-motion computations that have been described assume a single value of acceleration results from earthquakes of each magnitude and distance. However, in fact, a range of accelerations results; this variability in acceleration is commonly modeled by assuming that accelerations from earthquakes of a given magnitude and distance are lognormally distributed with standard deviation  $\sigma_a$  in  $\log_e$  acceleration. Assuming a lognormal distribution of accelerations tends to cause the calculated probabilistic ground-motion levels at all sites to be higher than the levels calculated when only median accelerations are used (for example, Bender, 1984a).

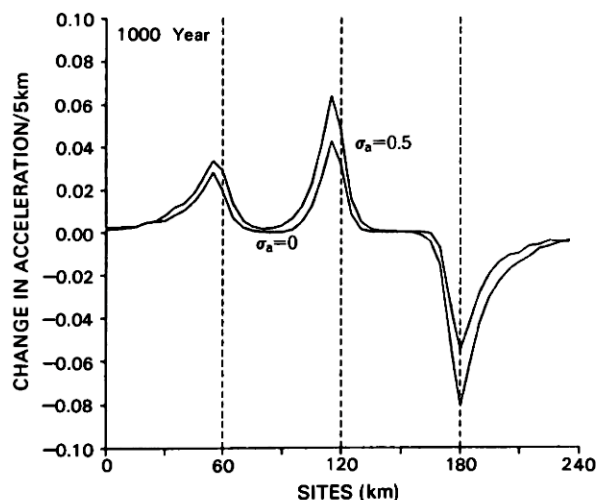


Figure 6. Differences in acceleration levels calculated for 1000-yr period at sites 5 km. apart, assuming  $\sigma_a=0$  and assuming  $\sigma_a=0.5$  in  $\log_e$ -acceleration in the attenuation function of Joyner and Boore (1981); no uncertainty ( $\sigma=0$ ) in earthquake location.

Figure 6 shows the differences in probabilistic acceleration levels calculated for a 1000-year exposure time at sites 5 km apart, when the sites and sources are the same as in figure 2. The calculation was done assuming no uncertainty in earthquake location using median values in the Joyner and Boore attenuation and using  $\sigma_a = 0.5$  in  $\log_e$  acceleration. (The calculation for  $\sigma_a = 0.5$  is the same as that shown in figure 2D for the 1000-year exposure period for  $\sigma = 0$ .) Figure 3 was included to illustrate that taking variability in the attenuation function into account, rather than reducing the differences in levels at sites near a boundary, may increase the site-to-site acceleration differences above those calculated when only median values are used in the calculations. This is discussed in detail in Bender (1986).

Assuming that  $\sigma_a$ , the standard deviation in  $\log$  acceleration, is independent of magnitude and distance permits SEISRISK III to take acceleration variability into account only after doing calculations using median acceleration values for all earthquakes. Median rather than mean acceleration values are used initially because  $\log$  acceleration (rather than acceleration) is assumed to be normally distributed; the median value of acceleration corresponds to the mean value of  $\log$ -acceleration.

Let us assume occurrences of median accelerations in each interval  $a_{j-1} < a \leq a_j$  have been accumulated in the corresponding  $reg(j)$ . Let  $regs(j)$  be a second



array in which to accumulate occurrences of accelerations  $a_{j-1} < a \leq a_j$  after acceleration variability has been taken into account. SEISRISK III redistributes a fraction of the acceleration occurrences accumulated in  $reg(j)$  to  $regs(k)$  (for each  $j$  and  $k$ ) as follows.

The entry in  $reg(j)$  is assumed to be concentrated at acceleration  $\bar{a}$  where

$$\log_e(\bar{a}) = \frac{\log_e[a_{j-1}] + \log_e[a_j]}{2}. \quad (12)$$

In a lognormal distribution, given  $\log_e \bar{a}$ , the probability of an acceleration in the range  $a_1 \leq a \leq a_2$  is the area under the normal probability curve:

$$p(a_1, a_2) = \frac{1}{\sqrt{2\pi}\sigma_a} \int_{\log_e a_1}^{\log_e a_2} \exp\left[-\frac{(x - \log_e \bar{a})^2}{2\sigma_a^2}\right] dx. \quad (13)$$

From equation 13, a fraction  $f(k)$  of the accelerations in  $reg(j)$  is placed in  $regs(k)$ ,

$$f(k) = p(a_{k-1}, a_k) \rho(j). \quad (14)$$

where  $\rho(j)$  = occurrences of (median) accelerations in the range  $a_{j-1} < a \leq a_j$  accumulated in  $reg(j)$ . Accounting for acceleration variability in this manner may be considerably faster computationally than methods which include acceleration variability for earthquakes of each magnitude and distance. Accounting for acceleration variability after calculations for median accelerations have been completed is possible because of the assumption that  $\sigma_a$ , the standard deviation in log acceleration is independent of distance and magnitude.

## CALCULATING GROUND-MOTION EXCEEDANCE PROBABILITIES

Given that average annual acceleration occurrence rates in the interval  $a_{j-1} < a \leq a_j$  have been accumulated in  $regs(j)$ ,  $1 \leq j \leq 55$ ,  $Ex[a_j]$ , the yearly occurrence rate of accelerations  $a > a_j$ , is simply

$$Ex[a_j] = \sum_{i=j+1}^{55} \rho^*(i) \quad (15)$$

where  $\rho^*(i)$  = accelerations in the range  $a_{i-1} < a \leq a_i$  accumulated in  $regs(i)$ . The yearly exceedance rate  $Ex(a)$  of acceleration  $a$  can be obtained by interpolating the  $Ex[a_j]$  in equation 15.

We are now in a position to determine the probability associated with the exceedance of any acceleration. Earthquake occurrences are assumed to have a Poisson distribution; that is, the probability that an earthquake will occur is the same at all times regardless of when the

last earthquake occurred. For a Poisson distribution with mean rate  $\mu$ , the probability of exactly  $k$  events during time  $t$  is

$$P(k, t) = \frac{(\mu t)^k \exp(-\mu t)}{k!}. \quad (16)$$

The probability of no event during the same time interval is, therefore,

$$P(0, t) = \exp(-\mu t). \quad (17)$$

Setting  $\mu = Ex(a)$ , the probability of not exceeding  $a$  in  $t$  years is

$$P(0, t) = \exp[-Ex(a)t]. \quad (18)$$

Taking the natural logarithm ( $\log_e$ ) of both sides of equation 18 gives

$$Ex(a) = -\frac{\log_e[P(0, t)]}{t}. \quad (19)$$

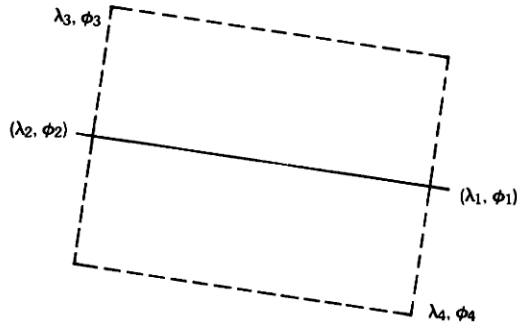
That is, the value of  $a$  for which equation 19 is satisfied is the acceleration which has probability  $p = P(0, t)$  of not being exceeded during the time interval  $t$ . (Equivalently,  $a$  has probability  $q = 1 - P(0, t)$  of being exceeded one or more times during time  $t$ .) By specifying any two of the three parameters ( $p, a, t$ ), one can determine the remaining parameter. However, in SEISRISK III a single probability level  $p$  and several values (up to 20) of time are specified and the corresponding values of  $a$  are obtained.

## MODEL IMPLEMENTATION-DETAILS AND REQUIREMENTS

In this section, we further describe the model and how it is implemented in SEISRISK III. Input parameters required to run SEISRISK III are identified and described briefly at the beginning of the program listing in Appendix A; here we give a more detailed explanation of the meaning and roles of some of these parameters.

### Geometry: Transformation to Local Equator

In both SEISRISK II and SEISRISK III, a pair of latitude and longitude values are selected (input) to be two points on a great circle that becomes the equator in a transformed coordinate system. The great circle is chosen to pass through the region of interest, and all subsequent input coordinates (for example, end points of source areas and fault segments) are transformed relative to the new equator. Furthermore, the subsequent calculations are done assuming a rectangular, rather than a spherical coordinate system; in the new system the (new) equator becomes the x-axis.



**Figure 7.** Four points define transformation and seismic-felt area. Great circle through  $(\lambda_1, \phi_1)$ ,  $(\lambda_2, \phi_2)$  becomes new equator.  $(\lambda_1, \phi_1) \rightarrow (d, 0)$ .  $(\lambda_2, \phi_2) \rightarrow (0, 0)$ .  $(\lambda_3, \phi_3)$  and  $(\lambda_4, \phi_4)$  become end points of rectangular area containing seismic felt points.  $(\lambda_3, \phi_3) \rightarrow (x_3, y_3)$ .  $(\lambda_4, \phi_4) \rightarrow (x_4, y_4)$ .

For points in the vicinity of the equator, calculations done using a rectangular coordinate system give a good approximation to those obtained using spherical coordinates. Because of convergence of parallels of latitude, the error could be substantial if the original  $(long, lat)$  coordinates are treated as being on a flat surface. However, this error is minimized when the coordinates are transformed to lie in the vicinity of a new equator, so that, for example, within a square region 12 degrees wide and centered on this equator, the maximum error introduced by using the straight line distance between two points  $(x_i, y_i)$ ,  $(x_j, y_j)$  instead of the great circle distance is less than 0.6 percent.

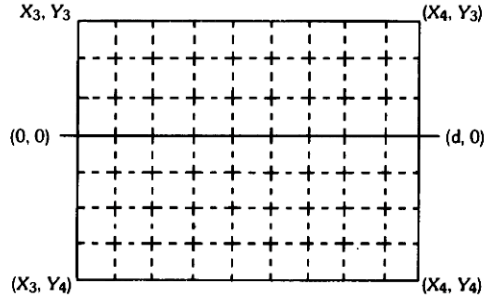
Sites at which accelerations are calculated lie on a rectangular  $m$  by  $n$  grid in which the site  $(x_i, y_j)$  is

$$x_i = ck(\text{longitude of row } i), \quad 1 \leq i \leq m$$

$$y_j = ck(\text{latitude of col } j), \quad 1 \leq j \leq n$$

where  $ck$  = conversion factor for degrees to kilometers at the equator.

Coordinates  $(long, lat)$  of two opposite corners (upper left, lower right) of the seismic felt area in the initial system and grid spacing are input to the program (fig. 7). Sites in the new coordinate system are located at even increments in  $\Delta x, \Delta y$  ( $\Delta y = \Delta x$ ) at intersections of grid lines. (In SEISRISK II,  $\Delta y = \Delta x$  was not required.) At the end of the computation, sites are transformed to the coordinates they would have in the original system. Sites may also be points evenly spaced on a number of lines; in this case,  $(long, lat)$  of two end points of each line, number of sites per line, and number of lines are specified as input. However, calculations for sites on a line are done assuming only seismically homogeneous source zones; earthquake location variability cannot be implemented for sites that are not on a grid of sites.



## Acceleration Table Levels

A table of 55 acceleration levels  $a_k$ , such that  $a_k < a_{k+1}$  for  $1 \leq k \leq 55$ , is constructed within the program. A scale factor (input) allows the original range of acceleration values in the table to be expanded or contracted. For a scale factor of one, 50 of the acceleration levels are in the range 0.02-1.0 g at increments of 0.02 g. Four acceleration levels are less than 0.02 g; accelerations at these levels are saved for calculations that are done when acceleration variability is taken into account; these accelerations are not printed in the output and, hence, are invisible to the analyst. A scale factor of 0.5 would permit values from 0.01g-0.5g. (The table is created in the subroutine "box" and other values may be substituted by replacing this subroutine.)

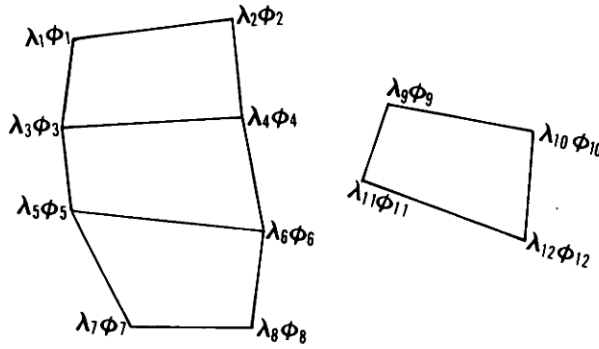
Given the table of acceleration levels  $a_k$ , a corresponding array  $reg(k)$  is set aside to accumulate fractional acceleration occurrences, with occurrences in the range  $a_{j-1} < a \leq a_j$  being placed in  $reg(j)$ . Accelerations less than  $a(1)$  are placed in  $reg(1)$  and those exceeding the highest entry  $a(max)$  (where  $max = 54$ ) are placed into  $reg(max+1)$ . Earthquakes as point events within quadrilateral source areas and as finite length ruptures along linear fault segments may contribute to accelerations at a site. We shall discuss these two cases separately.

## Source Zone Computations

### Area Computations

A seismic source zone is defined as a seismically homogeneous area enclosed by one or more arbitrary quadrilaterals, connected or disjoint (fig. 8).

Recall that an acceleration "box" represents occurrences of accelerations in some range  $a_{j-1} < a \leq a_j$ . For a specified magnitude, acceleration box boundaries  $[a_{j-1}, a_j]$  correspond to earthquakes at distances in the range  $r(j) \leq r < r(j+1)$  from the site. The contribution to accelerations  $a_{j-1} < a \leq a_j$  produced by magnitude  $m$  earthquakes in a seismically homogeneous quadrilateral is equal to the rate per unit area within the



**Figure 8.** Seismic source zone consisting of two sets of quadrilaterals

Input: For each quadrilateral set

(1) jseg ifr tot

jseg=number+1 of quadrilaterals in this set

ifr identifies current set,  $1 \leq \text{ifr} \leq \text{tot}$

tot=total number of sets of quadrilaterals

(2) pairs of quadrilateral corner points.

In this example, input is

4 1 2

λ<sub>1</sub> φ<sub>1</sub> λ<sub>2</sub> φ<sub>2</sub>

λ<sub>3</sub> φ<sub>3</sub> λ<sub>4</sub> φ<sub>4</sub>

λ<sub>5</sub> φ<sub>5</sub> λ<sub>6</sub> φ<sub>6</sub>

λ<sub>7</sub> φ<sub>7</sub> λ<sub>8</sub> φ<sub>8</sub>

2 2 2

λ<sub>9</sub> φ<sub>9</sub> λ<sub>10</sub> φ<sub>10</sub>

λ<sub>11</sub> φ<sub>11</sub> λ<sub>12</sub> φ<sub>12</sub>

quadrilateral, multiplied by the area within the quadrilateral at distances  $r(j) \leq r < r(j+1)$  from the site. The area of the quadrilateral within the distance range  $r(j) \leq r < r(j+1)$  is approximately the area contained within a ring of radius  $r$  and width  $dr = [r(j+1) - r(j)]$  centered at the site (fig.9).

The distance range over which earthquakes produce accelerations in the interval  $a_{j-1} < a \leq a_j$  depends on magnitude. The accelerations from earthquakes of each magnitude in the appropriate distance range within a source area are summed to obtain the total contribution from these earthquakes added into  $reg(j)$ . The procedure is repeated for all acceleration intervals, and subsequently for all sources.

To save computation time, for a particular site and source, both SEISRISK II and SEISRISK III create a table of distance versus arc-length (for a predetermined set of radii) and then later interpolate in this table to obtain areas at desired distances. In this case

$$\text{Area} = \phi r dr \quad (20)$$

where  $\phi$ =arc length in radians of that portion of the circle of radius  $r=r_j$  that lies wholly within the quadrilateral  
 $dr$ =interval width.

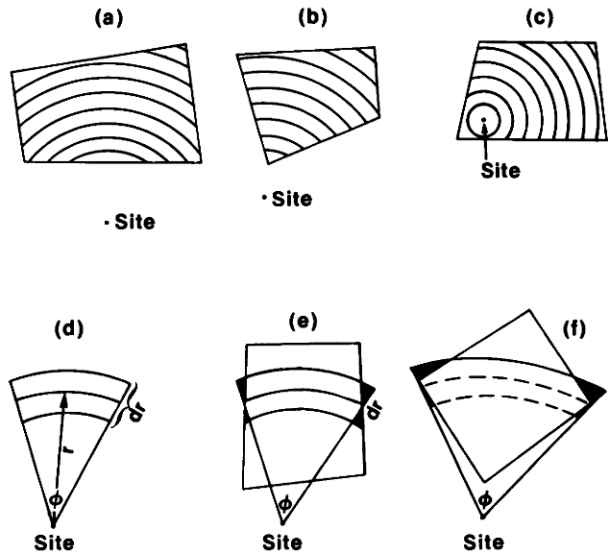
The quantity  $\phi r$  is tabulated as a function of  $r$  for seven evenly spaced radii  $r$  per quadrilateral and for two to eight additional radii to vertices and edges, depending upon the location of the site relative to the source. The accuracy of the calculated areas depends on the geometry of the source, location of the site, and on the number of radii used in the calculations.

To test the accuracy of the arc-area computations for a fixed value of  $itst$ , areas for a set of 5856 quadrilateral-site pairs (used in a California study) were evaluated using  $itst=5, 7, 9, 11$ , and  $13$ . For  $itst=7$ , the

("arc estimate" area - "true" area) / "true" area exceeded 0.03 only five times and never exceeded 0.04; for  $itst=5$ , the error exceeded 0.04 seven times; for  $itst=9, 11$  and  $13$ , the largest error was 0.02. We concluded that  $itst = 7$  gave sufficient accuracy. The number of evenly spaced radii is set by parameter  $itst = 7$  in subroutine RRISK;  $itst$  may be increased by recompiling the program, but this should probably be done only under unusual circumstances (for example, for testing purposes), because using larger values of  $itst$  increases the computation time.

### Grid Spacing Requirements For Modeling Uncertainty

SEISRISK III initially calculates expected acceleration occurrences at each site on a grid assuming that seismicity is uniform throughout a source zone. Then to determine acceleration occurrences at a site at  $(X_1, Y_1)$ ,



**Figure 9.** Arcs of circles with centers at the site are used to approximate the area of a quadrilateral.  $\text{Area} = \phi r dr$  in *D*, *E*, and *F*. Shaded area in *E* and *F* is the difference between the quadrilateral area and arc area. In *E* errors nearly cancel; in *F* the arc area is an overestimate.

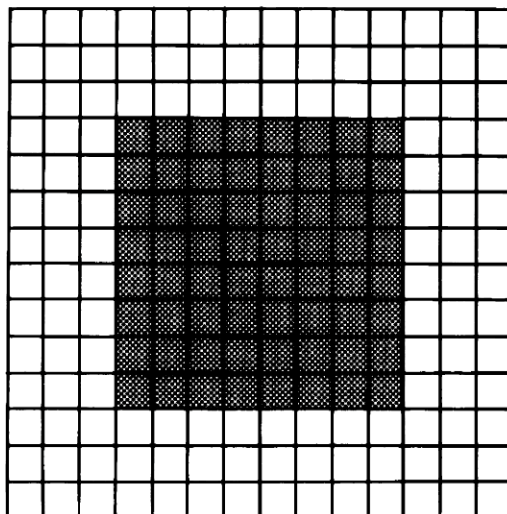


with earthquake location uncertainty included, SEISRISK III computes a weighted average of acceleration occurrences previously calculated at various sites assuming uniform seismicity. The fractional contribution or weight assigned to the accelerations calculated at a site at  $(X_i, Y_k)$  depends on  $\sigma$  and on the distance between  $(X_1, Y_1)$  and  $(X_i, Y_k)$  (equation 2).

The grid spacing or distance between sites is of major importance in these calculations. Accelerations at sites at distances greater than  $2.0\sigma - 2.5\sigma$  from  $(X_1, Y_1)$  have low weights and little effect on the weighted accelerations at  $(X_1, Y_1)$ . This means that if the spacing between sites is large relative to  $\sigma$ , the only accelerations that will have much weight at  $(X_1, Y_1)$  are those originally calculated at  $(X_1, Y_1)$  (for uniform seismicity), and, in this case, the weighted results will give a poor approximation to the true effects of earthquake location uncertainty. In tests to determine the maximum acceptable grid spacing, we found that we could use a spacing  $dx$  as large as  $\sigma$  and still obtain reasonable results; in these tests, for  $dx=\sigma$  results were quite consistent with those obtained using  $dx=\sigma/2$ .

In order to take into account accelerations at sites at distances  $2\sigma$  from every site of interest, SEISRISK III extends the gridded region to include additional sites (fig.10). If the sites of interest are at row and column intersections on an  $(n)(n)$  grid, and the distance between rows and columns is  $dx$ , then the grid must be extended in each direction by at least  $n_{rc} = 2.0\sigma/dx$  rows and columns; that is, the calculations must be done at  $(n + 2n_{rc})^2$  sites in order to model effects of earthquake location variability at each of the original  $n$  sites. The number of additional sites can be quite large if  $\sigma$  is small relative to  $dx$ . For example, if  $dx=10$  and  $\sigma=40$ , the number of sites increases from  $n^2$  to  $(n+16)^2$ .

Increasing the number of sites at which the calculations are performed increases the storage requirements as well as the computation time. Because acceleration occurrences calculated for uniform seismicity at various sites are used in calculating the weighted averages, we retain these occurrences in a two-dimensional array. This requires an array of dimension of at least  $(II, JJ)$ , where  $II$ =number of acceleration levels in the *reg* array, and  $JJ$ =the number of sites (including the sites added for earthquake location variability) at which the calculation is performed ( $JJ \geq (n + 2n_{rc})^2$ ). Furthermore, a second array of the same size is set aside in which to accumulate the weighted acceleration occurrences. In SEISRISK III, the two arrays RAWBIN and SMOBIN for saving the accelerations are dimensioned (55,1600). A total of 1600 sites (including the sites added for earthquake location variability) can be processed. (Dimensions of arrays SMOBIN and RAWBIN can be adjusted to change the number of sites permitted; the parameter *mazar* must be set equal to the site dimension.) Sav-



**Figure 10.** Sites of interest located at intersections of grid lines within darkened area. If the spacing between grid lines is  $dx$  and the standard deviation in earthquake location is  $\sigma$ , the gridded region is extended in each direction by  $n_{rc} = 2.0\sigma/dx$  (rounded to the next highest integer) rows and columns.

ing results in two-dimensional arrays contrasts with the procedure in SEISRISK II, in which acceleration densities calculated at a single site were written onto disc (logical unit 3 or logical unit 4) rather than being saved in a large array.

### Using Sources with Different $\sigma$ Values

SEISRISK III permits earthquakes associated with different source-zones to have different values of  $\sigma$ . If  $\sigma$  is the same for earthquakes in a number of zones, weighted averages of accelerations can be calculated after the calculations assuming uniform seismicity are completed for earthquakes in all of the zones. On the other hand, if a different  $\sigma$  is associated with earthquakes in each source zone, the weighting (equation 2) of accelerations resulting from earthquakes in a zone must be performed after the calculations for uniform seismicity have been done for that zone.

To instruct the program when to calculate weighted averages, a value of an input parameter *num* is entered for each source zone along with other data for that zone; *num*=98 indicates averaging should be performed for all accelerations accumulated since the averaging was last performed. The current value of  $\sigma$  is used in this averaging, and one need enter a value of  $\sigma$  only for source zones for which *num* = 98. Figure 11A gives an abbreviated flowchart for the source zone computations.

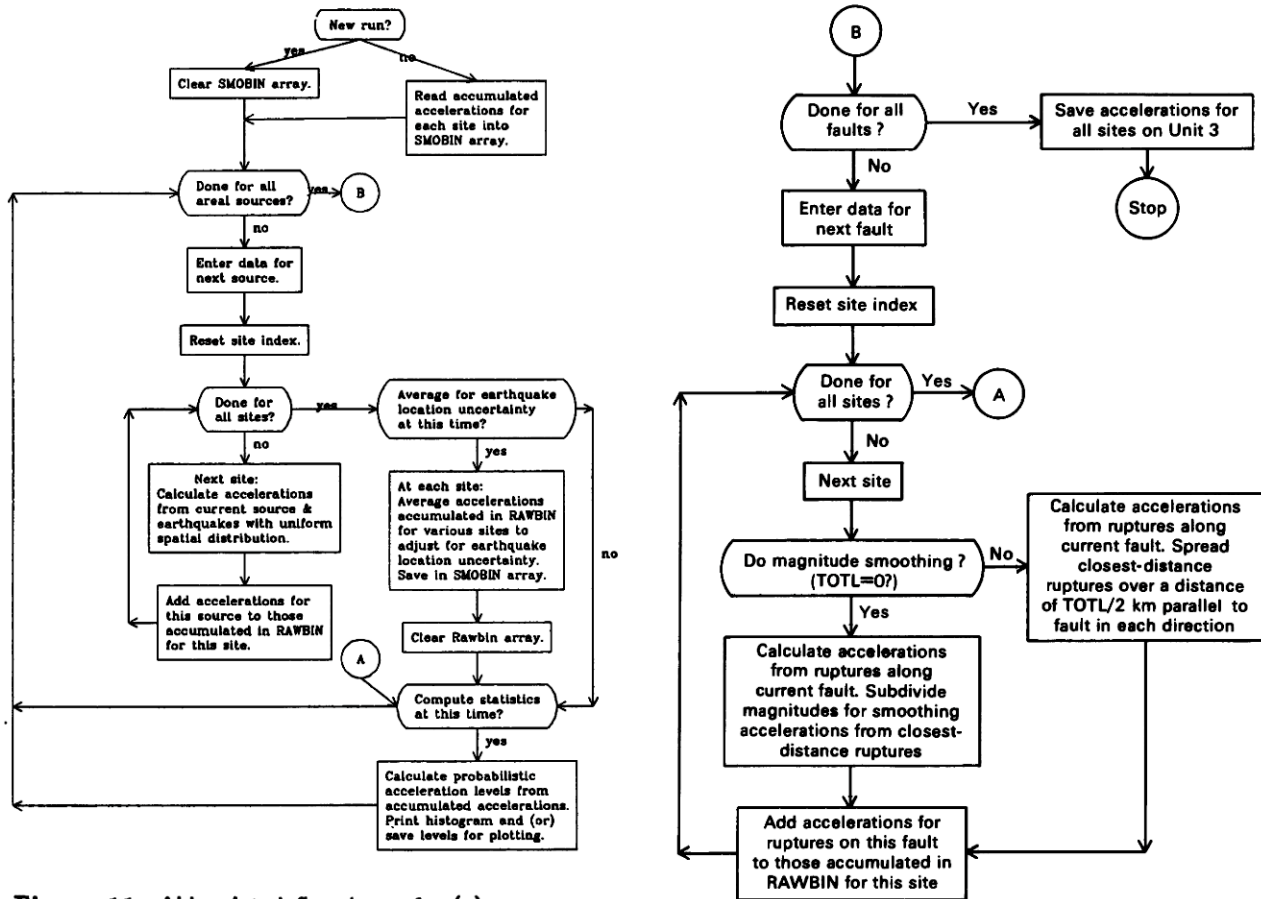


Figure 11. Abbreviated flowcharts for (a) source-zone calculations and (b) fault-rupture calculations.

When considering using different values of  $\sigma$  for earthquakes in different zones, we should keep the following considerations in mind.

1. The number of rows and columns and the spacing between sites remains the same throughout the analysis. The value of  $n_{rc}$ , the number of additional rows and columns by which the initial gridded area must be extended in each direction to account for earthquake location uncertainty, is determined by the largest value of  $\sigma$  to be used for any source-zone. However,  $n_{rc}$  is calculated before the source zone inputs are processed, and therefore, the largest value of  $\sigma$  must be identified in the program initially.
2. If considerably different values of  $\sigma$  are assigned to earthquakes in different areas, the grid spacing must be sufficiently small to adequately represent the intended variability for those earthquakes with the smallest  $\sigma$ . At the same time, a sufficient number of marginal sites (determined by  $n_{rc}$ ) must be provided to permit accurate calculations when  $\sigma$  is large. Accomodating both these requirements can lead to evaluating accelerations at considerably more sites than necessary, both when  $\sigma$  is small (because of an unnecessarily large value of  $n_{rc}$

in this case) and when  $\sigma$  is large (because of the small increment  $dx$  relative to  $\sigma$ ).

3. Allowing different values of  $\sigma$  can lead to an anomalous ridge of seismicity along one side of a joint boundary of contiguous source zones. In general, we suggest that two source zones with different  $\sigma$  values be separated by at least two to three times the difference in the  $\sigma$  values.

## Fault Computations

### Rupture Length Computations

A rupture of a magnitude  $m$  earthquake is defined to have length  $l_i(m)$  with probability  $p(i)$ , where

$$\log_{10}[l_i(m)] = a + bm + \eta(i)\sigma_i \quad (21)$$

where  $a$  and  $b$  are constants.

$p(i)$  = probability of rupture of length  $l_i(m)$

$$p(i) = \frac{1}{\sqrt{2\pi}} \int_{f(i)}^{f(i+1)} \exp\left(-\frac{x^2}{2}\right) dx \quad (22)$$



where  $f(i)$  is selected so that  $\eta(i)$  is

$$f(i) < \eta(i) \leq f(i+1); \quad f(1) = -\infty; \quad f(n_r + 1) = \infty;$$

$$\sum_{i=1}^{n_r} p(i) = 1.$$

$n_r$  = number of rupture lengths per magnitude.

SEISRISK III allows either one rupture length or five rupture lengths per magnitude depending on whether median rupture lengths only are assumed or rupture length variability is to be modeled in the length versus magnitude relationship. Thus

$$\eta(i) = \begin{cases} 0, & \text{if only median rupture length is used} \\ & \text{(one rupture length per magnitude),} \\ \text{five values in the range } -2 \leq \eta(i) \leq 2 & \\ & \text{(five lengths per magnitude).} \end{cases} \quad (23)$$

The values of  $a$ ,  $b$  and  $\sigma_l$  may be provided as input for the rupture-length magnitude relationship (equation 21). If  $a$ ,  $b$  and  $\sigma_l$  are all zero or blank, previous values of these quantities are used. If no values are provided, default values of Bonilla and Buchanan (1970) are used:

$$a = -1.085 \quad b = 0.389 \quad \sigma_l = 0.52$$

If  $a \neq 0$ ,  $b \neq 0$ , and  $\sigma_l = 0$  are input, a single rupture-length value is used.

#### Magnitude Smoothing:

##### Calculation of Gutenberg-Richter $b$ -Value

The Gutenberg-Richter  $b$ -value used to determine the number of earthquakes in an arbitrary magnitude interval is calculated from the number  $N(m_1)$  of earthquakes at magnitude  $m_1$  and  $N(m_2)$ , the number at magnitude  $m_2$ ; that is,

$$b = -\frac{1}{\Delta m} \log \left[ \frac{N(m_2)}{N(m_1)} \right]. \quad (24)$$

where  $m_1$  is the lowest magnitude used in the calculations and  $m_2 = m_1 + \Delta m$  (equation 4). However, if only a single magnitude is specified, neither the magnitude interval nor the  $b$ -value can be calculated, and hence default values of  $b$  and  $\Delta m$  are provided. (The default values are  $b = 1.0$  and  $\Delta m = 0.6$ .)

##### Distance Smoothing Option

To instruct the program to do distance (rather than magnitude) smoothing, the distance  $d$  between the parallel faults must be input.

$$totl = \begin{cases} d & \text{distance between faults if this is a set of} \\ & \text{"dummy" faults used to approximate} \\ & \text{a uniform field of faults.} \\ 0 & \text{if individual fault.} \end{cases}$$

If  $totl = 0$  (or is blank), the program assumes that the faults are individual faults, and the previously discussed magnitude (rather than distance) smoothing is performed. Distance smoothing over a range of  $\pm d/2$  km for a single fault can be forced by setting  $totl = d$  for a single fault. The program cannot do both distance and magnitude smoothing. Figure 11B shows an abbreviated flowchart for the fault rupture calculations.

#### Attenuation Function: Input Table, Interpolation, Scale Factor

The program computes ground-motion (acceleration) by interpolating in a table of acceleration as a function of magnitude and distance. Interpolations are linear in distance and linear in log acceleration and magnitude; in SEISRISK II, interpolations were linear in log magnitude rather than in magnitude.

The table of ground-motions is read in for up to eight magnitudes and 20 distances. Velocity or some other ground-motion parameter may be substituted for acceleration. If the ground-motion values read in are in a range different from 0.02 to 1.0, the appropriate scale factor must be input to make the "box" levels compatible with the desired values.

#### Halving the Magnitude Intervals

The program requires that earthquake rates be entered for each source at equally spaced magnitude intervals. The magnitudes that are entered are assumed to be values at the centers of the intervals. If input parameter  $inos = 1$ , the program will divide each magnitude interval into two parts and perform the calculations for magnitudes at the midpoints of the new intervals. Earthquake rates are calculated for each interval from the total earthquakes assigned to the original intervals using a Gutenberg-Richter  $b$ -value based on the number of earthquakes in the first two (original) magnitude intervals, as in equation 24.

#### Output Options

As stated earlier, we seek the acceleration that, with probability  $q$ , will be exceeded (or probability  $p = 1 - q$  will not be exceeded) during  $t$  years. We have seen that this is the value  $a$  of acceleration for which  $Ex(a)$  the yearly exceedance rate of  $a$  at the site is

$$Ex(a) = \frac{\log_e(p)}{t}. \quad (25)$$

Assuming yearly exceedance rates (as in equation 15) have been calculated for each of 55 acceleration levels, the solution  $a$  that corresponds to the specified  $(p, t)$  may be found by interpolating between the calculated values.

Solutions are computed for up to 20 time periods  $t$  using rates derived for median accelerations only and also for the same set of times using rates that include acceleration variability.

The results, including the acceleration histogram, may optionally be written to file 16 for later printing. Summary results may be saved on file 2 for future plotting.

Accelerations from the earthquakes within a source area or along a set of faults are calculated for all sites before proceeding to the next source area or fault, with acceleration occurrences being added to those obtained previously for each site. Therefore, partial results may be examined at any stage. Output options are given for each source set by the input parameter *iprint*.

*iprint*=

- 1: continue to next set; no print, no statistics
- +1: compute solutions  $a$ ; write histogram and solutions to file 16, then continue to next set
- +2: same as +1 but in addition save the solutions on file 2
- +3: no write to file 16; save solutions on file 2.

### Accumulating Rates In Successive Runs

Rates that have been accumulated in each acceleration interval for each site are saved in file 3 at the end of a run. These values may be read back into the 2-dimensional array SMOBIN for a subsequent run. (The saved rates are rates that have been adjusted for earthquake location uncertainty.) Doing successive runs makes it possible, for example, to accumulate accelerations from sources having different attenuation functions. The equator, maximum  $\sigma$ , gridded region, and rows and columns containing sites of interest should not be changed from their original values in a run continuation.

$$isw = \begin{cases} 1 & \text{if this is a run continuation} \\ 0 & \text{if this is a new run.} \end{cases}$$

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## APPENDIX A

### Details of Fault-Rupture Computations

A fault consists of one or more connected line segments. A rupture may be wholly contained within one segment or overlap any number of fault segments. It may be as long as the fault, but cannot extend beyond the end of the fault. The distance from a rupture to a site is the distance from the site to the point on the rupture closest to the site.

A conceptually simple approach for calculating accelerations at a site resulting from ruptures along the fault pretends that ruptures of length  $l_i(m)$  of magnitude  $m$  earthquakes occur only at a set of  $n$  distinct locations, with rupture centers evenly spaced at increments  $\Delta l_i(m)$  on the fault. In this case, the  $k^{th}$  rupture center is at a distance  $d_i(k)$  from one end of the fault where

$$d_i(k) = \frac{l_i(m)}{2} + (k-1)\Delta l_i(m) \quad (A1)$$

and

$$\Delta l_i(m) = \frac{L - l_i(m)}{n-1}, \quad 1 \leq k \leq n. \quad (A2)$$

The fractional seismicity  $C(i, m)$  associated with one such rupture is

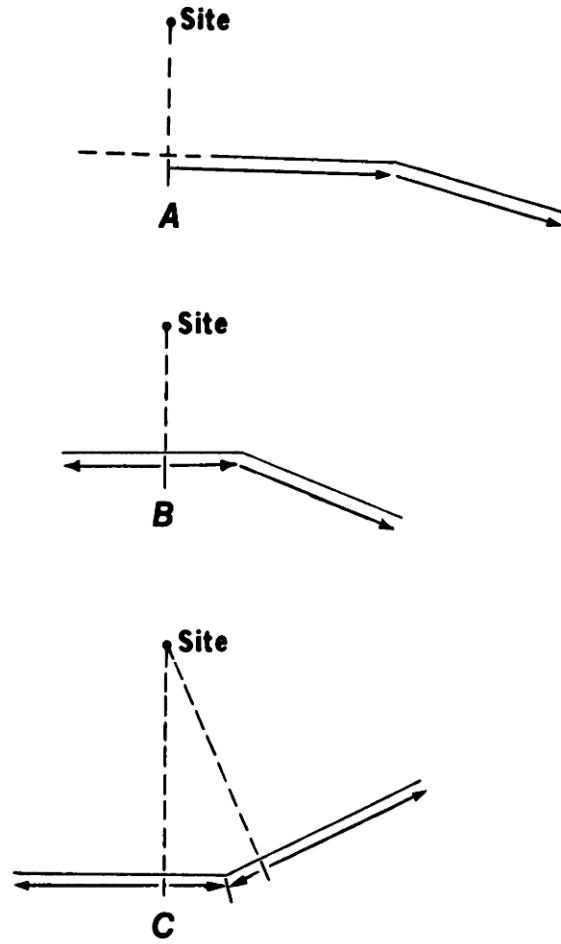
$$C(i, m) = \frac{\rho(m)p(i)}{n} f, \quad (A3)$$

$\rho(m)$ =rate of magnitude  $m$  earthquakes for the set of faults,

$f$ =length of this fault/total length of faults in set,

$p(i)$ =probability of rupture of length  $l_i(m)$ , for either 1 or 5 rupture lengths per magnitude.

For each rupture, the closest distance to the site is determined (fig.A1). For one such rupture of magnitude  $m$  and length  $l_i(m)$ , the acceleration interval is found that contains the acceleration produced by a rupture of this magnitude at this distance. The fractional seismicity  $C(i, m)$  is added to the accumulator  $reg(j)$ , for the appropriate  $j$ .



**Figure A1.** Arrows indicate direction of increasing distance along the fault from the site.  
**A** Closest distance from site to fault is at left end of fault. Distance increases monotonically along the fault as one moves in the direction of the arrows.  
**B** Distance to site decreases as one moves along the fault from left end of fault to point A, then increases monotonically as one continues to other end.  
**C** Distance to site alternately decreases and increases as one moves along the fault.

The computer time required may be substantial if the computation of nearest distance from the site to a rupture is done repeatedly for different rupture locations along the fault. Some transformations are made in SEISRISK III that simplify the site-to-rupture computations. The following describes these transformations.

For an arbitrary fault segment  $1 \leq i \leq n$ , let

$$L(i) = \sqrt{[x_1(i) - x_2(i)]^2 + [y_1(i) - y_2(i)]^2} \quad (A4)$$

= length of segment  $i$ ,

where  $[x_1(i), y_1(i)], [x_2(i), y_2(i)]$  = coordinates of end-points of segment  $i$ .

The coordinates are transformed so that fault end points

$$\begin{aligned} [x_1(i), y_1(i)] &\rightarrow (0, 0) \\ [x_2(i), y_2(i)] &\rightarrow [L(i), 0] \end{aligned} \quad (A5)$$

and the site at  $(F_x, F_y)$

$$(F_x, F_y) \rightarrow [X(i), P(i)]$$

where

$$\begin{aligned} X(i) &= [F_x - x_1(i)] \cos \alpha + [F_y - y_1(i)] \sin \alpha, \\ P(i) &= -[F_x - x_1(i)] \sin \alpha + [F_y - y_1(i)] \cos \alpha. \end{aligned} \quad (A6)$$

= perpendicular distance from site to line  $i$ .

$$\cos \alpha = \frac{x_2(i) - x_1(i)}{L(i)}, \quad \sin \alpha = \frac{y_2(i) - y_1(i)}{L(i)}.$$

After  $L(i)$ ,  $X(i)$  and  $P(i)$  have been determined, the fault segment may be regarded as located on the  $x$ -axis extending from  $(0,0)$  to  $[L(i), 0]$  and the site located at  $[X(i), P(i)]$  (fig.A2).

If the fault consists of  $N$  connected line segments, SEISRISK III computes and saves  $L(i)$ ,  $X(i)$ ,  $P(i)$  for each segment,  $1 \leq i \leq N$ . Assume an arbitrary rupture has one end point on segment  $j$  and the second on segment  $k$ , where  $1 \leq j \leq k \leq N$ . The coordinates of the end points of the part of the rupture that lies on segment  $i$  ( $j \leq i \leq k$ ) are defined as  $[R_1(i), 0], [R_2(i), 0]$  (fig.A3).

$$0 \leq R_1(i) \leq R_2(i) \leq L(i) \quad \text{for } i=j \text{ or } i=k, \quad (A7)$$

$$R_1(i) = 0, R_2(i) = L(i) \quad \text{for } j < i < k.$$

If a rupture extends over three or more fault segments, the rupture overlap with an interior segment is the entire segment.

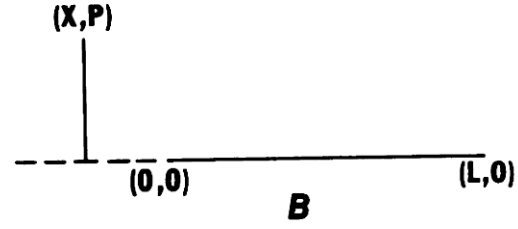
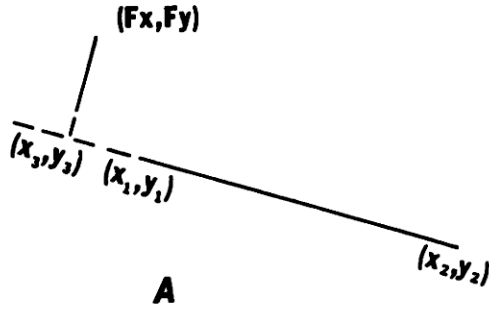
In the new system, the relative end and overlap points  $R_1(i), R_2(i)$  are determined from the total distance  $\delta$  along the fault of the initial end point (start) of the given rupture and from the length of the rupture and fault segments, by calculating  $Rem(j)$  for each segment  $j$ , where  $Rem(j)$  is defined as

$$Rem(j) = \delta - \sum_{i=0}^{j-1} L(i), \quad \text{where } L(0) = 0.$$

The first segment containing the rupture beginning at  $\delta$  is the segment  $j$  for which  $Rem(j) > 0$  and  $Rem(j+1) \leq 0$ . (For this  $j$ ,  $R_1(j) = Rem(j)$ .) The segment containing the final end point of a rupture of length  $l$  is determined similarly by substituting  $(\delta+l)$  for  $(\delta)$  above. If the rupture overlaps segment  $i$ , let  $D^2(i)$  = square of the closest distance from the site to the rupture on segment  $i$ . If the rupture extends over several segments, the value of  $D^2(i)$ , that is the minimum for all such segments, is required.  $D^2(i)$  is defined as follows:

$$D^2(i) = \begin{cases} [R_1(i) - X(i)]^2 + P(i)^2 & \text{if } X(i) < R_1(i), \\ [X(i) - R_2(i)]^2 + P(i)^2 & \text{if } X(i) > R_2(i), \\ P(i)^2 & \text{if } R_1(i) \leq X(i) \leq R_2(i). \end{cases} \quad (A8)$$





**Figure A2.** Fault segment and site. *A* original coordinates. *B* relocated so that segment lies on X-axis. *A* Original coordinates. Segment of fault extends from  $(x_1, y_1)$  to  $(x_2, y_2)$ ; site is at  $(F_x, F_y)$ . Perpendicular from site intersects extended fault line at  $(x_3, y_3)$ . *B* Relocated. Fault and site moved so that fault extends from  $(0, 0)$  to  $(L, 0)$ . Site moved to  $(X, P)$ .

Thus, having calculated and saved  $L(i)$ ,  $P(i)$ ,  $X(i)$ , determining the (square of the) closest distance requires only a simple computation for each fault segment that the rupture overlaps.

Assigning the fractional seismicity (equation A3) associated with a rupture to a single acceleration interval results in chunks of seismicity in a set of intervals. This means that the calculated acceleration densities do not vary smoothly as a function of acceleration level. In some cases, SEISRISK III can avoid doing calculations for a distinct set of ruptures and instead directly calculate the fraction of ruptures of magnitude  $m$  earthquakes that produce accelerations in each interval,  $a_{j-1} < a \leq a_j$ . When the more direct approach cannot be employed, SEISRISK III calculates site-to-rupture distance for  $n$  ruptures of magnitude  $m$  earthquakes, evenly spaced along the fault (equation A1). If rupture  $i$  is at distance  $d_i$  from the site, and rupture  $i+1$  is at distance  $d_{i+1}$  from the site, SEISRISK III assumes a fraction  $1/n$  of the possible ruptures are in the distance range  $d_i \leq d \leq d_{i+1}$  (assuming  $d_i \leq d_{i+1}$ ). It then finds the accelerations  $a(i)$  and  $a(i+1)$  produced by magnitude  $m$  ruptures at distances  $d_i$  and  $d_{i+1}$ . Assuming  $a(i+1) \leq a_j \leq a_{j+1} \leq a(i)$ , the fractional seismicity assigned to the acceleration interval  $a_j \leq a \leq a_{j+1}$  is given by

$$C(i, m) \frac{a_{j+1} - a_j}{a(i) - a(i+1)}, \quad (A9)$$

where  $C(i, m)$  is defined as in equation A3.

We next consider two cases in which a more direct calculation of fractional accelerations can be done for each acceleration interval. We use the notation and definitions already developed.

Case 1.  $X(i) < 0$  for all  $1 \leq i \leq n$  (or  $L(i) < X(i)$  for all  $1 \leq i \leq n$ ). The point on the rupture nearest the

site is at one end of the rupture with the distance to the site increasing (or decreasing) monotonically along the rupture.

If  $X(i) < 0$  for all  $i$ , for any rupture overlapping segments  $j$  through  $k$ , where  $j \leq k$ , then  $[R_1(j), 0]$  is the point closest to the site. Similarly if  $X(i) > L(i)$  for all  $i$ , then  $[R_2(k), 0]$  is the point of the rupture closest to the site.

Assume  $X(i) < 0$ , for all  $i$ . ( $L(i) < X(i)$  could be considered in a symmetrical argument.) In this case the rupture point nearest to the site is  $R_1(i)$  for the lowest  $i$  containing the rupture. Recall that, for a rupture of length  $l$ , rupture centers may be located at any distance  $d_c$  along the fault for which

$$\frac{l}{2} \leq d_c \leq L - \frac{l}{2} \quad (A10)$$

or, equivalently, the lowest end point of a rupture may be at any distance  $\delta$  for which

$$0 \leq \delta \leq L - l.$$

The acceleration produced by a rupture of magnitude  $m$  and length  $l_i(m)$  decreases as the rupture moves along the fault from a maximum acceleration  $a_{max}[l_i(m)]$  for a rupture with closest end point at  $\delta = 0$  to a minimum acceleration  $a_{min}[l_i(m)]$  for a rupture with closest end point at  $\delta = L - l(m)$ . Therefore, for an acceleration interval such that

$$a_{min}[l_i(m)] \leq a_{j-1} < a \leq a_j \leq a_{max}[l_i(m)]$$

a corresponding interval exists along the fault

$$h_m(j) \leq \delta \leq h_m(j-1)$$

for magnitude  $m$  earthquakes in which (closest) end points of ruptures producing these accelerations may be located. Because these ruptures are at a distance  $d_m(j) \leq d \leq d_m(j-1)$  from the site,

$$h_m(j) = \sqrt{d_m(j)^2 - P(i)^2} + X(i); \quad [X(i) < 0] \quad (A11)$$

$$h_m(j-1) = \sqrt{d_m(j-1)^2 - P(i)^2} + X(i). \quad (A12)$$

Ruptures at distance along the fault  $h_m(j) \leq \delta < h_m(j-1)$  therefore contribute to accelerations in the range  $a_{j-1} < a \leq a_j$  at the rate

$$\text{rat}[a_j, l_i(m)] = \rho(m) \frac{h_m(j-1) - h_m(j)}{L - l_i(m)} f p(i) \quad (\text{A13})$$

where  $f$ ,  $p(i)$  and  $\rho(m)$  are defined as in equation A3.

For  $n$  rupture lengths, given that

$$l_1(m) > l_2(m) \dots > l_n(m), \quad (\text{A14})$$

a corresponding section of fault  $0 \leq \delta \leq L - l_i(m)$  may contain closest end points of ruptures and

$$L - l_n(m) > L - l_{n-1}(m) > \dots > L - l_1(m). \quad (\text{A15})$$

Now define

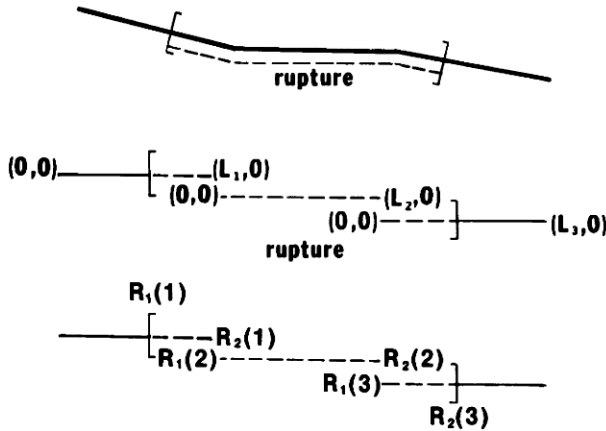
$$h_{m,i}(j) = \begin{cases} h_m(j) & \text{if } h_m(j) < L - l_i(m) \\ L - l_i(m) & \text{otherwise, } 1 \leq i \leq n. \end{cases}$$

The total contribution to accelerations  $a_{j-1} < a \leq a_j$  is the sum over rupture lengths for magnitude  $m$  ruptures:

$$\text{rat}(a) = \rho(m) \sum_{i=1}^n \frac{h_{m,i}(j-1) - h_{m,i}(j)}{L - l_i(m)} f p(i) \quad (\text{A16})$$

where  $h_{m,i}(0) = 0$ .

Thus, one need only compute one set of distances  $h_m(j)$  per magnitude along the fault rather than repeating the process for each rupture length, so that including rupture length variability in the calculation does not substantially increase the computation time.



**Figure A3.** Rupture overlaps three fault segments. Each segment is relocated to lie on the X-axis. End points of the part of the rupture on segment  $i$  are  $[R_1(i), 0]$  and  $[R_2(i), 0]$ .

Case 2. The distance from the fault to the site decreases monotonically as one moves a distance  $d_c$  along the fault to some point on segment  $k$   $[(I_k, 0)$  in the new system] and then increases steadily as one continues to the other end. This corresponds to:

$$X(i) < 0 \text{ for } 1 \leq i \leq k-1,$$

$$X(k) = I_k = d_c - \sum_{i=1}^{k-1} L(i),$$

$$X(i) > L(i) \text{ for } k+1 \leq i \leq n.$$

In this case, all ruptures of length  $l_i(m)$  with mid-points at a distance  $\delta$  along the fault

$$d_c - \frac{l_i(m)}{2} < \delta < d_c + \frac{l_i(m)}{2} \quad (\text{A17})$$

overlap the point  $d_c$  and hence are regarded as being at the same closest distance from the site. The acceleration at the site produced by these magnitude  $m$  ruptures is the acceleration at distance  $P(k)$ , the perpendicular distance to the  $k^{\text{th}}$  segment. Because ruptures cannot extend past the end of the fault, the location of possible rupture centers is restricted to a part of the fault and the contribution at  $P(k)$  becomes

$$\text{rat}[a(j), l_i(m)] = \rho(m) f \sum_{i=1}^n \frac{D_{u_i}(m) - D_{l_i}(m)}{L - l_i(m)} p(i), \quad (\text{A18})$$

where

$$D_{u_i}(m) = \min[d_c + \frac{l_i(m)}{2}, L - \frac{l_i(m)}{2}], \quad (\text{A19})$$

$$D_{l_i}(m) = \max[d_c - \frac{l_i(m)}{2}, \frac{l_i(m)}{2}].$$

Now the bookkeeping becomes more complicated, but we can look at all ruptures of length  $l_i(m)$  with centers at distance  $\delta$  on the fault

$$\frac{l_i(m)}{2} \leq \delta \leq L - \frac{l_i(m)}{2} \quad (\text{A20})$$

and remove those already accounted for, those with centers

$$d_c - \frac{l_i(m)}{2} \leq \delta < d_c + \frac{l_i(m)}{2}. \quad (\text{A21})$$

The remaining ruptures may be separated into two sets: those (if any) with centers

$$\frac{l_i(m)}{2} \leq \delta < d_c - \frac{l_i(m)}{2} \quad (\text{A23})$$

and those for which

$$d_c + \frac{l_i(m)}{2} < \delta \leq L - \frac{l_i(m)}{2}.$$

The last two intervals correspond to ruptures whose closest end points to the site are at  $\delta$ , where

$$\min[l_i(m), d_c] \leq \delta < d_c \quad (A24)$$

for one set of ruptures, and at

$$d_c < \delta \leq \max[d_c, L - l_i(m)] \quad (A25)$$

for the other set. These two sets of ruptures do not overlap and the fault may be regarded as consisting of two distinct segments in which the distance from the rupture to the site decreases (or increases) monotonically as one moves along the fault. Hence, each of these segments may be treated separately as in Case 1, and the accelerations resulting from ruptures along these segments may be added to those already accounted for.

## APPENDIX B

### Sample Inputs and Outputs

Sample inputs and outputs are provided in this appendix. An explanation of the input variables (file 15) and formats is provided at the beginning of the program listing in Appendix C.

The output file (file 16) gives ground-motion statistics for each site. The site coordinates (longitude, latitude) in decimal degrees appear first, followed by the shortest distance from the site to the closest fault. Then two arrays are given, one for zero attenuation variability and one for attenuation variability  $\sigma_a$ . The columns

in each array give for the ground-motion (g.m.) in the  $i^{\text{th}}$  row:

occ/yr: occurrence rate per year of earthquakes producing ground motions at the site in the range  $\text{g.m.}(i-1) < \text{g.m.} \leq \text{g.m.}(i)$ ;

exc/yr: exceedance rate per year of  $\text{g.m.}(i)$ ;

r(events): the average number of events required to produce an exceedance of  $\text{g.m.}(i)$ ,  

$$r(\text{events}) = \frac{\text{total yearly events}}{\text{exc/yr of g.m.}(i)};$$

r(yrs): return period in years,

$$r(\text{yrs}) = \frac{1}{\text{exc/yr of g.m.}(i)}.$$

Following the arrays is "total yearly events," the total number of earthquakes from all sources in this run. This total should be nearly the same at all sites; slight differences between sites are due to inaccuracies in calculating fractional areas and assigning proportional seismicity to these areas. Large differences probably indicate an error in the computation.

The concluding lines for each site read

"xxx ext prob=yyy for zzz years",

that is, xxx is the extreme probability that ground-motion yyy will not be exceeded in zzz years.

Included also in this appendix are a listing of a program to read (binary) output file 02 and the output of that program for the sample problem. File 02 contains only the summary information needed for mapping purposes: site coordinates (long, lat) and ground-motions calculated at the specified probability level for the times of interest (for both zero-attenuation variability and one value of  $\sigma_a$ ).

File 15--Inputs

```

SEIRISK III Sample Problem
0 20
  90 3 10 50 250
1 0 5 0
123 0 38 0 120 0 38 0
123 00 39 00 120 00 37 00 .100
5 6 5 6
1
3
123 39 123 38
6 12
  wataashh79' 8 5 7 6 4 6 5 6 5 2 5 2 5 2 5 2
3 22 74 73 67 45 195 135 135 135 135 135 135 135 135 135 135
6 43 64 62 53 36 047 02 02 02 02 02 02 02 02 02 02
16 09 49 43 32 19 052 02 0052 0052 0052 0052 0052 0052 0052 0052
32 18 36 28 17 09 02 0052 0052 0052 0052 0052 0052 0052 0052
64 3 21 14 06 03 0051 0013 0013 0013 0013 0013 0013 0013 0013
96 5 12 07 03 0023 0023 0023 0023 0023 0023 0023 0023 0023
160 9 045 025 015 0053 0053 0053 0053 0053 0053 0053 0053 0053
321 8 013 0076 0026 0012 00021 00001 00001 00001 00001 00001 00001 00001
643 0 0034 0019 00065 0003 00005 00001 00001 00001 00001 00001 00001 00001
1288 0 00085 00047 00016 00007 00001 00001 00001 00001 00001 00001 00001 00001
2570 0 00021 00012 00004 00002 00001 00001 00001 00001 00001 00001 00001 00001
5140 0 0001 00006 00002 00001 00001 00001 00001 00001 00001 00001 00001 00001
00 100 -1 zn01
3 1
122 20 37 22 122 00 37 32
123 85 38 77 123 68 39 05
124 12 39 32 123 89 39 40
  106 281 74 1 97
6 100 5 500 4 900 4 300 zn02
00 65 2 -1
2 1 1
120 05 34 09 119 26 34 20
120 40 33 50 119 77 33 39
  031 073 172 405 955 2 248
7 1 0 -1 zn03 20
120 40 34 68 120 00 34 78
120 56 34 93 120 28 35 00
120 78 35 20 120 38 35 39
121 00 35 48 121 00 36 00
121 51 36 00 121 47 36 50
121 82 36 45 121 80 36 80
122 00 36 73 122 21 37 11
2 2 2
121 80 36 80 121 77 36 88
122 21 37 11 122 20 37 22
0198 0641 2074 6712
6 100 5 500 4 900 4 300 ft01-1.085 .389 .50
99 1 0 -1
12 1 1
117 37 34 17 118 09 34 59 118 67 34 79 119 03 119 03 119 03 119 03 119 03 119 03 119 03
119 26 34 90 119 54 35 06 121 27 36 66 121 69 121 69 121 69 121 69 121 69 121 69 121 69
122 10 37 27 123 77 39 01 124 01 39 36 124 16 124 16 124 16 124 16 124 16 124 16 124 16
  0021 0057 015 040
8 500 7 900 7 300 6 700

```



## File 16--Outputs

```

SEISRISK III Sample Problem
ismw0: new run--no previous results included
extreme probability 0.900
scale factor for ground motion "box" levels= 1.00
coordinates input in decimal degrees
variability in attenuation, sigma= 0.50
grid oriented parallel to great circle thru ( 123.00, 38.00), ( 120.00, 38.00)
corners of gridded area--upper left= 123.00, 39.00
lower right= 120.00, 37.00
longitude increment= 0.1000 (decimal degrees)
latitude increment = 0.1000 (decimal degrees)
gridded region contains 27 rows, 30 cols, including border
for this run begin at row 5 and row 6, begin col 5 and col 6
new coordinates (km) gridded area
upper left= 294.73 0.00; lower right= 144.70 -111.30
sites are also located on 1 line(s)
3 sites per line
line 1 end points at 123.000, 39.000 and 123.000, 38.000
attenuation function watashh79

```

dist(km)	8.50	7.60	6.60	5.60	5.20	4.20
3.22	0.74000	0.73000	0.67000	0.45000	0.19500	0.07200
6.43	0.64000	0.62000	0.53000	0.36000	0.13500	0.04700
16.09	0.49000	0.43000	0.32000	0.19000	0.05200	0.02000
32.18	0.36000	0.28000	0.17000	0.09000	0.02000	0.00520
64.30	0.21000	0.14000	0.06000	0.03500	0.00510	0.00130
96.50	0.12000	0.07000	0.03000	0.01380	0.00230	0.00042
160.90	0.04500	0.02500	0.01000	0.00500	0.00083	0.00010
321.80	0.01340	0.00740	0.00260	0.00120	0.00021	0.00003
643.00	0.00340	0.00190	0.00065	0.00030	0.00007	0.00001
1288.00	0.00085	0.00047	0.00016	0.00007	0.00001	0.00001
2570.00	0.00021	0.00012	0.00004	0.00002	0.00001	0.00001
5140.00	0.00010	0.00006	0.00002	0.00001	0.00001	0.00001

```

yrnoc= 100. iprint=-1 for area zn01
122.200 37.220 122.000 37.320
123.890 38.970 123.680 39.050
124.120 39.320 123.890 39.400
nr of levels of seismicity = 4
zn01 beta= -1.6319
before normalizing to rate/year
earthquake rate / year
occurrences= 0.019700 0.007400 0.002810 0.001060
magnitudes= 4.30 4.90 5.50 6.10
zn01 area= 5457. sq km, rate/sq km= 0.36101E-05 for mags 4.00- 4.60

yrnoc= 65. iprint=-1 for area zn02
120.050 34.090 119.260 34.200
120.400 33.500 119.770 33.390
nr of levels of seismicity = 6
zn02 beta= -1.4268
before normalizing to rate/year
earthquake rate / year
occurrences= 2.248000 0.955000 0.405000 0.172000 0.073000 0.031000
magnitudes= 6.70 7.30 7.90 8.50 9.10 9.70
zn02 area= 5212. sq km, rate/sq km= 0.66157E-05 for mags 4.00- 4.60

yrnoc= 1. iprint=-1 for area zn03
smooth for earthquake location uncertainty sigma= 20.0
120.400 34.680 120.000 34.780
120.560 34.930 120.280 35.000
120.780 35.200 120.380 35.390
121.000 35.480 121.000 36.000
121.510 36.000 121.470 36.500
121.820 36.450 121.800 36.800
122.000 36.730 122.210 37.110
121.800 36.800 121.770 36.880
122.210 37.110 122.200 37.220
nr of levels of seismicity = 4
zn03 beta= -1.9574
earthquake rate / year
occurrences= 0.671200 0.207400 0.064100 0.019800
magnitudes= 4.30 4.90 5.50 6.10
zn03 area= 10075. sq km, rate/sq km= 0.66621E-04 for mags 4.00- 4.60

yrnoc= 1. iprint=-1 for area ft01
Distance between dummy faults= 0.0
fault 1 of 1
117.37 34.17, 118.09 34.59, 118.67 34.79, 119.03 34.85,
119.26 34.90, 119.54 35.06, 121.27 36.66, 121.69 36.94,
122.10 37.27, 123.77 39.01, 124.01 39.36, 124.16 39.93,
nr of levels of seismicity = 4
ft01 beta= -1.6347
earthquake rate / year
occurrences= 0.040000 0.015000 0.005700 0.002100
magnitudes= 6.70 7.30 7.90 8.50
fault rupture length parameters a1= -1.085 b1= 0.389 sig1= 0.50

yrnoc= 10. iprint= 2 for area ft02
Distance between dummy faults= 0.0
fault 1 of 4
122.40 38.10, 122.90 38.66, 123.38 39.10, 123.94 39.79,
fault 2 of 4
122.60 38.68, 123.10 39.20, 123.74 39.95,
fault 3 of 4
123.16 38.16, 122.46 38.80, 122.75 38.98, 123.18 39.56,
123.40 39.98,
fault 4 of 4
122.50 39.10, 122.80 39.25, 122.95 39.43,
nr of levels of seismicity = 3
ft02 beta= -1.7455
before normalizing to rate/year
earthquake rate / year
occurrences= 0.057000 0.020000 0.007200
magnitudes= 6.70 7.30 7.90
fault rupture length parameters a1= -1.085 b1= 0.389 sig1= 0.50

```



```

c program to read summary file02 created by seisrisk ii
c
c file 02 contains for each site the level of ground motion
c that has probability "perct" of not being exceeded during
c "ntim" years. Each of "ntim" values of "jtim"
c and "jtim" values are given first for "zero variability in
c attenuation" for all times "jtim", then for "variability in
c attenuation" sigma = sd for all values of "jtim".
c
c first record contains:
c 1row1,1row2,1col1,1col2 = first and last rows, first and last
c columns in gridded area containing sites for which
c ground motions were calculated in seisrisk ii.
c indvpt = total number of sites on lines
c ntimes = number of exposure times
c jtimes = exposure times in years
c sd = standard deviation in attenuation variability.
c perct: ground motion calculated has probability "perct"
c of not being exceeded during "jtim" years
c
c all records after first:
c 1row1,1row2,1col1,1col2 = first and last rows, first and last
c columns in gridded area containing sites for which
c ground motions were calculated in seisrisk ii.
c "perct" probability and "jtim(k)" years where
c i=1,2,...,ntim; k=i for zero attenuation variability
c i=ntim+1, ntim+2,...,2*ntim; k=i-ntim
c for variability in attenuation sigma = sd
c
c dimension sol(48),jtim(24)
c read (2) 1row1,1row2,1col1,1col2,indvpt,ntim, (jtim(jv),jv=1,ntim)
c &ntim),perct,sd
c ntim=2*ntim
c write(44,10), (jtim(jv),jv=1,ntim), (jtim(jv),jv=1,ntim)
c 10 format(30x,'exposure times',/,' long lat -dty)
c 20 read (2,44:40), 10long,10lat,10sol(jt,jt=1,ntim2)
c write(44,30), 10long,10lat,10sol(jt,jt=1,ntim2)
c 30 format(14x,27f,3d,6p4,3)
c go to 20
c 40 call exit
c end

```

The following are the summarized results that were written to file 02 for the sample problem

Long	lat	10yr	50yr	250yr	10yr	50yr	250yr
122.48	38.61	0.095	0.280	0.500	0.105	0.324	0.676
122.35	38.61	0.079	0.224	0.418	0.087	0.295	0.671
122.48	38.51	0.104	0.277	0.418	0.115	0.323	0.614
122.35	38.51	0.086	0.224	0.418	0.096	0.301	0.602
123.00	39.00	0.159	0.351	0.598	0.138	0.422	0.824
123.00	38.50	0.239	0.398	0.454	0.110	0.320	0.634
123.00	38.00	0.219	0.443	0.499	0.203	0.570	0.925

g.m.	acc/yr	sec/yr	r(events)	r(yr)	variability in atten.	sigma=0	sigma=50
0.02	1.08957	0.03459	32.5	28.3	0.0844	0.0245	0.0245
0.04	0.10339	0.02421	46.4	41.7	0.0357	0.0245	0.0245
0.06	0.00327	0.01235	91.0	81.0	0.0411	0.01794	0.01794
0.08	0.00241	0.00594	113.2	100.7	0.0403	0.01390	0.01390
0.10	0.00220	0.00774	145.3	129.3	0.0381	0.01199	0.01199
0.12	0.00292	0.00482	164.8	146.6	0.0303	0.00906	0.00906
0.14	0.00112	0.00370	197.2	175.4	0.03152	0.00754	0.00754
0.16	0.00101	0.00469	239.9	213.4	0.01116	0.00638	0.00638
0.18	0.00077	0.00392	286.8	253.1	0.00992	0.00546	0.00546
0.20	0.00083	0.00309	363.4	323.2	0.00674	0.00473	0.00473
0.22	0.00048	0.00261	430.7	383.1	0.00650	0.00412	0.00412
0.24	0.00029	0.00232	484.3	430.7	0.00402	0.00352	0.00352
0.26	0.00021	0.00211	533.3	474.3	0.00336	0.00284	0.00284
0.28	0.00018	0.00193	583.9	519.3	0.00331	0.00254	0.00254
0.30	0.00013	0.00180	625.8	556.6	0.00227	0.00227	0.00227
0.32	0.00015	0.00164	684.4	608.7	0.0023	0.00227	0.00227
0.34	0.00018	0.00146	747.8	668.7	0.0023	0.00184	0.00184
0.36	0.00013	0.00124	813.0	748.9	0.0018	0.00164	0.00164
0.38	0.00027	0.00119	943.4	839.1	0.0014	0.00130	0.00130
0.40	0.00024	0.00095	1180.6	1050.1	0.0014	0.00123	0.00123
0.42	0.00021	0.00074	1518.3	1350.4	0.0011	0.00112	0.00112
0.44	0.00018	0.00056	2002.5	1781.2	0.0010	0.00102	0.00102
0.46	0.00014	0.00042	2672.0	2376.7	0.0009	0.00093	0.00093
0.48	0.00014	0.00028	4015.7	3571.9	0.0008	0.00085	0.00085
0.50	0.00009	0.00019	5936.6	5280.4	0.0007	0.00077	0.00077
0.52	0.00006	0.00013	8530.0	7604.9	0.0007	0.00070	0.00070
0.54	0.00012	0.00001	9999.9	9999.9	0.0006	0.00064	0.00064
0.56	0.00001	0.00000	9999.9	9999.9	0.0006	0.00059	0.00059
0.58	0.00000	0.00000	9999.9	9999.9	0.0005	0.00054	0.00054
0.60	0.00000	0.00000	9999.9	9999.9	0.0005	0.00049	0.00049
0.62	0.00000	0.00000	9999.9	9999.9	0.0004	0.00044	0.00044
0.64	0.00000	0.00000	9999.9	9999.9	0.0004	0.00044	0.00044
0.66	0.00000	0.00000	9999.9	9999.9	0.0004	0.00044	0.00044
0.68	0.00000	0.00000	9999.9	9999.9	0.0004	0.00044	0.00044
0.70	0.00000	0.00000	9999.9	9999.9	0.0003	0.00038	0.00038
0.72	0.00000	0.00000	9999.9	9999.9	0.0003	0.00038	0.00038
0.74	0.00000	0.00000	9999.9	9999.9	0.0003	0.00033	0.00033
0.76	0.00000	0.00000	9999.9	9999.9	0.0003	0.00033	0.00033
0.78	0.00000	0.00000	9999.9	9999.9	0.0002	0.00027	0.00027
0.80	0.00000	0.00000	9999.9	9999.9	0.0002	0.00025	0.00025
0.82	0.00000	0.00000	9999.9	9999.9	0.0002	0.00023	0.00023
0.84	0.00000	0.00000	9999.9	9999.9	0.0002	0.00021	0.00021
0.86	0.00000	0.00000	9999.9	9999.9	0.0002	0.00018	0.00018
0.88	0.00000	0.00000	9999.9	9999.9	0.0002	0.00016	0.00016
0.90	0.00000	0.00000	9999.9	9999.9	0.0001	0.00013	0.00013
0.92	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
0.94	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
0.96	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
0.98	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.00	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.02	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.04	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.06	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.08	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.10	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.12	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.14	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.16	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.18	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.20	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.22	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.24	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.26	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.28	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.30	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.32	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.34	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.36	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.38	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.40	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.42	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.44	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.46	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.48	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.50	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.52	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.54	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.56	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.58	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.60	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.62	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.64	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.66	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.68	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.70	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.72	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.74	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.76	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.78	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.80	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.82	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.84	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.86	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.88	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.90	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.92	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.94	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.96	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
1.98	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.00	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.02	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.04	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.06	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.08	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.10	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.12	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.14	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.16	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.18	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.20	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.22	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.24	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.26	0.00000	0.00000	9999.9	9999.9	0.0001	0.00011	0.00011
2.28	0.00000	0.00000	9999.9	9999			

### Listing of Computer Program SEISRISK III

24 Seisrisk III: A Computer Program for Seismic Hazard Estimation



```

c
c*** iscom = dimension of following sin.cos arrays
common/sinco/sinx(500),cosx(500),sinu(500),cosu(500)
dimension x(4),y(4),z(4),noc(24),fm(25)
common/lrnp/dls(24),disq(24,26),jseg,jsegm,
& x1(24,26),y1(24,26),z1(24,26),sta(24,26)
common/lrdis/dls(24),perp(24),min(24),persq(24)
dimension jm(26),disrem(24)
common/accel/gm(25),exdif(101),atah(20,25),aclim(2)
c arrays on previous line inserted for spreading out accelerations
real noc
integer dsu
data lnsel,lns,npls/0,0,0/
data idsu/'decimal degrees', 'degrees, minutes'
data r,pi2,pi,pih/6378.,6.2831854,3.1415927,1.5707963/
rad=pi/180.
c input on unit 15
c rewind 15
c saved on unit 3 for later use or continuation
c rewind 3
allerr=0.
dlevsv=0.
flevst=0.
levnom=0.
tot5=0.
20 format(a80)
30 format(16,30) ttitle
write(16,30) ttitle
30 format('1 a80)
c isw tells whether to include data from a previous run, zero if no.
c sigma=maximum value of sigma (km, in earthquake location)
read(15,*)isw,sigmax
if(isw.eq.0) write(16,40)
40 format(' isw=0: new run--no previous results included')
if(isw.eq.1) write(16,50)
50 format(' isw=1: run continuation; add to previous results')
if(isw.eq.0) go to 70
idrc=1
60 read(3,end=70) (smobin(i,rp,idrc),i,rp=1,maxdim)
idrc=idrc+1
go to 60
70 rewind 3
c prob=extreme probability in decimal
c ntims are number of 'jtimes' for which calculation is done
c jtimes are the durations for which the extreme motions are
to be calculated at the prob extreme probability level
read(15,*)prob,ntims,(jtim(jv),jv=1,ntims)
if (prob.lt.1.) go to 90
write(16,80) prob
80 format(' prob =f8.3' must be decimal less than one')
call exit
90 continue
100 format(' extreme probability f6.3
& for exposure times (years) '2(10i5//)
ex=alog(prob)
ex=alog(ex)
do 110 jv=1,ntims
110 ftim(jv)=jtim(jv)
c
c read(15,*) scale,dsu,sd,inos
write(16,120) scale,idsu(dsu+1),idsu(dsu+1),sd
120 format(' scale factor for ground motion "box" levels=
&f7.2', coordinates input in ' a16,
& coordinates are printed in 'a16,
& variability in attenuation, sigma=f6.2)
c scale=multiplicative factor for acceleration levels
c dsu=0 if input is decimal degrees; =1 for degrees,minutes
c sd=standard deviation for log acceleration
c read (x1,y1), (x2,y2) along, lat in dec. degrees of 2 points
c to transform to equator
c (x1,y1) to (0.,0)(x2,y2) to (dist. between points,0)
130 read(15,*) x1,y1,x2,y2
write(16,140) x1,y1,x2,y2
140 format(' grid oriented parallel to great circle thru ('
& f7.2',f6.2'),(f7.2',f6.2')')
c convert to decimal degrees unless dsu=0.
if(dsu.ne.0) call cond(x1,x2,y1,y2)
x1=x1*rad
x2=x2*rad
y1=y1*rad
y2=y2*rad
c transform to equator
call inleqr(x2,y2,x1,y1)
c
c call toeqr(x1,y1,x1t,y1t)
c call toeqr(x2,y2,x2t,y2t)
c read in upper left and lower right corners of affected rectangle
c -- degrees, minutes is 20.30=20 degrees,30 minutes (dsu=1)
c (decimal degrees 20.50 = 20 1/2 degrees (dsu=0) )
c opposite corners of gridded area (in latitude and longitude)
c gridded area becomes rectangle surrounding new equator.
c sites are at uniform increments in latitude and longitude.
c within this gridded area in new coordinate system.
c in order to limit affected area you can specify
c beginning and ending rows and columns as irow1 and 2
c and icol1 and 2 (next input line)
c phi=phi+latitude
c phi=phi+longitude
c
c read(15,*) f11,phi1,f12,ph2,phinc
write(16,160) f11,phi1,f12,ph2,phinc
160 format(' corners of gridded area--upper left=f7.2',f7.2,
&f23r' lower right=f7.2',f7.2)
c
c convert to decimal degrees (unless dsu=0)
if(dsu.eq.0) go to 170
call cond(phi,f11,ph2,f12)
call condec(phinc)
170 continue
c
c determine number of subregions in map area and set up subscripting
c irowno of rows
c icolno of cols
write(16,180) phinc,phinc
180 format(' longitude increment=f7.4' (decimal degrees)'
& latitude increment =f7.4' (decimal degrees)')
phinc=phinc*rad
f1inc=phinc

```



```

dphinc=dphinc+r
dphinc=dphinc/2.
dflinc=dflinc+r
fl1=fl1+r
ph1=ph1+r
fl2=fl2+r
ph2=ph2+r
c transform to equator and x-y rectangular coordinate system
c convert to kilometers
c call toeqr(fl1,ph1,flam1,ph1)
c call to eqr(fl2,ph2,flam2,ph2)
if(ph1.gt.0) go to 210
if(ph12.gt.0) go to 220
if(ph12.gt.0) go to 220
c apparent error in inputs or xformatation
190 write(16,200)
200 format(' probable input error---felt region upper left, '
& ' lower right', must be on opposite sides of new equator')
call exit
210 if(ph12.le.0) go to 230
go to 190
220 phs=ph1
ph1=ph12
ph12=phs
fls=flam1
flam1=flam2
flam2=fls
c have proper orientation---continue
230 ph1=ph1
if(flam1.lt.0) go to 240
fl1=flam1
c find total number of cols
go to 260
240 write(16,250)flam1,flam2
250 format(' apparent error in xforming lat and long limits '
& ' to eqr---', flam1='f9.3', fl2='f9.3')
call exit
260 continue
ph1r=ph1r
ph12r=ph12r
fl1r=fl1r
flam2r=flam2r
c adjust to include area beyond edges for later smoothing
c determine nbord=number of additional rows and columns at
c which ground motions must be calculated to account for
c earthquake location variability.
nbord=2.*sigmax/dphinc+.01
if(sigmax.eq.0.) go to 280
if(nbord.ge.2) go to 280
write (16,270) dphinc,sigmax,nbord
270 format(' increments between sites =e12.5' too large for '
& ' sigmax=e12.5' nbord='i5')
call exit
280 ph1=ph1+nbord*phinc
if (flam2.gt.fl1) fl1nc=fl1nc
fl1=fl1+nbord*fl1nc
flam2=flam2+nbord*fl1nc
icols=(ph1-ph12)/phinc+.01+nbord
icols=(fl1-flam2)/fl1nc+.01+nbord
write(16,290)icols,icols,nbord
290 format(' gridded region contains 'i4,' rows, 'i4' cols '
& ' including border 'i4' rows and cols')

```

```

read(15,*)row1,row2,icol1,icol2
c limits rows and columns of felt points for which calculation is
c to be made.
write(16,300)row1,row2,icol1,icol2
300 format(' for this run begin at row 'i3' end row 'i3'
& ', begin col 'i3' end col 'i3')
c for earthquake location variability, compute at more sites
c calculate at additional rows and cols
row2=row2+2*nbord
icol2=icol2+2*nbord
icols=icol2-icol1+1
icols=row2-row1+1
if(row1.eq.0) row1=0
if(icol1.eq.0) icols=0
if(row1.le.row2) go to 320
write(16,310) row1
310 format(' row1 = 'i4' too large---error stop')
call exit
320 if(icol1.le.icol2) go to 340
write(16,330) icol1
330 format(' icol1 = 'i4' too large---error stop')
call exit
340 if(icols*rows.le.maxr) go to 370
350 write(16,360)rows,icols,indvpt,maxdim
360 format(' rows='i4' icols='i4' pts on line='i4'
& ' exceeds maxdim of 'i4')
call exit
370 continue
c read in number of line segments containing sites at which
c calculation is to be done (in addition to or instead of
c computing for sites on a fixed grid)
c zero if no line segments
uplv=ph1r
uplx=fl1r
boty=ph12r
botx=flam2r
write(16,380) uplx,botx,uplv,boty
380 format(' new coordinates (km) gridded area '
& ', upper left='2f9.2', lower right='2f9.2')
read(15,*)indv
write(16,390) indv
390 format(' sites are also located on 'i3' line(s)')
if (indv.eq.0) go to 450
c calculate accelerations at nvs sites on each of indv lines
c read in end point pairs (long,lat) xel,yel,xz2,yz2
c for each of indv lines
nvs sites are evenly spaced on line*
nvs=2 gives accelerations at end points of line*
nvs=3 acceleration at 2 end points plus center of line,etc.
read(15,*) nvs
write(16,400) nvs
400 format('i4' sites per line')
indvpt=0
do 450 i=1,indv
read(15,*) xel,yel,xz2,yz2
write(16,410) i,xel,yel,xz2,yz2
410 format(' line 'i2' end points at 'f7.3','f7.3' and 'f7.3'
& ', 'f7.3')

```



```

if(yrnoc.eq.0.) go to 2950
line=1
710 write(16,690)yrnoc,iprint,dumid
720 if(line.eq.1) go to 830
if(num.eq.98) write(16,703) als
705 format(' smooth for earthquake location uncertainty sigma='
1 f5.1)
nbro=0
ist=0
iend=0
c read identifiers for source area inputs
730 read(15,*)jseg,ifr,itot
write(16,740) ifr,itot
740 format(' source '12' of '12)
c jseg=number of pairs of quadrilateral end points in this
c single source
c itot=number of sources in this source area
c if=identifies which source, ifr=1,2,...,itot
if (itot.le.10) go to 760
750 format(' itot =14' to large--max 10 jseg=14
& , ifr='14)
call exit
760 ist=iend+1
iend=ist+jseg-1
if (iend.le.50) go to 780
write(16,770)
770 format(' too many quadrilaterals in source--max 50')
call exit
c read in boundaries for quadrilaterals for all subareas
c in this source
c read limits of seismic area x(i)=lambda(i), y(i)=phi(i)
c upper right(x,y) then upper right(x,y) --1st quadrilateral
c 2nd line lower left, lower right 1st quadrilateral=
c 2nd card upper left, upper right 2nd quadrilateral, etc.
c
780 do 810 i=ist,iend
read(15,*) (xsav(i,1),ysav(i,1), i=1,2)
790 format(4f6.0)
800 format(1h 4f10.3)
write(16,800)(xsav(i,1),ysav(i,1), i=1,2)
810 continue
if(ifr.eq.itot) go to 820
nbro=nbro+1
ibr=(nbro)isend
go to 730
820 num=iend
go to 960
c fault line read
830 j=1
write(16,840) totl
840 format(' Distance between dummy faults='f5.1)
dsun=0.
c fault inputs
850 read(15,*) jseg , ifr, itot
write(16,860) ifr, itot
860 format(' fault '12' of '12)
c jseg=number of end points of connected segments
c for the current fault
c jsegone plus number of segments
c
c itot=number of distinct faults in this zone
c ifr identifies current fault, ifr=1,2,...,itot
870 write(16,880) jseg, itot
880 format(' jseg='14' (max 24) itot='14' (max 26)--stop')
890 if(itot.gt.26) go to 870
c read in end points long, lat of each segment
read(15,*) (xl(i,j),yp(i,j), i=1,jseg)
write(16,900)(xl(i,j),yp(i,j), i=1,jseg)
900 format(1h 4 f10.2, f8.2, ',')
jml(j)=jseg
do 920 i=1,jseg
c convert to decimal degrees unless dsu=0
if(dsu.eq.0) go to 910
call condec(yp(i,j))
call condec(xl(i,j))
910 continue
c transform to equator--rectangular coordinates
yp(i,j)=yp(i,j)*rad
xl(i,j)=xl(i,j)*rad
call to eqr(xin,ypin,xout,yout)
xl(i,j)=xouter
yp(i,j)=youter
920 continue
c compute line parameters, lengths for fault segments
jsegm=jseg-1
dtot(j)=0.
do 930 i=1,jsegm
xdelta=xl(i+1,j)-xl(i,j)
ydelta=yp(i+1,j)-yp(i,j)
disq(i,j)=xdelta*xdelta+ydelta*ydelta
dist(i,j)=sqrt(disq(i,j))
coa(i,j)=(xl(i+1,j)-xl(i,j))/dist(i,j)
sia(i,j)=(yp(i+1,j)-yp(i,j))/dist(i,j)
dtot(j)=dtot(j)+dist(i,j)
dsun=dsun+dtot(j)
if(ifr.eq.itot) go to 950
940 j=j+1
go to 850
c read in num of occurrences for each magnitude (12f6.2)
c both area and fault input
950 jtot=j
960 read (15,1170) (noc(i),i=1,12)
lev=12
970 if(noc(lev).ne.0.0) go to 990
lev=lev-1
if(lev.gt.0) go to 970
write(16,980)
980 format(' apparent input err for no of occurrences at each level')
go to 2990
990 continue
write(16,1000)lev
1000 format(' nr of levels of seismicity = '12)
c read in corresponding mag. interval center points. (lev of them)
1010 format(' before normalizing to rate/year')
1020 format(' earthquake rate / year')
read(15,1170)(fm(i),i=1,lev)
c magnitudes must be evenly spaced and in increasing order.

```

```

c if not entered in increasing order, program reverses order
c of magnitudes and corresponding earthquake occurrences
c at each magnitude.
c Don't confuse this with the original table of ground
c motions as a function of magnitude and distance.
c The magnitudes in that table must be read in decreasing
c (rather than increasing) order.
  if(fm(1),lt,fm(2)) go to 1040
  if(lev,eq.1) go to 1040
  lev=lev/2
  ll=lev
  do 1030 lm=1,levh
    sav=fm(l)
    fm(l)=fm(ll)
    fm(ll)=sav
    sav=nmoc(l)
    nmoc(l)=nmoc(ll)
    nmoc(ll)=sav
  1030 ll=ll-1
  1040 continue
  if(lev.gt.1) go to 1050
c we cannot calculate beta from numbers of earthquakes
c if only one level--so give a default beta so we
c can later spread out
  beta=-2.0
  fm(2)=fm(1)+.6
  delm=.6
  go to 1060
1050 continue
  dlev=abs(fm(1)-fm(2))
  beta=log( nmoc(2)/nmoc(1) )/dlev
1060 write(16,1070) dmid, beta
1070 format(1x,e4,' beta=',f8.4)
  if(ynoc,eq.1.) go to 1090
  write(16,1010)
  this is a normalization to rate per year
  do 1080 lm=1,lev
    nmoc(lm)=nmoc(lm)/ynoc
  1080 continue
1090 write(16,1020)
  write(16,1100)(nmoc(l),l=1,lev)
1100 format(' occurrences='12f10.6)
  write(16,1110)(fm(l),l=1,lev)
1110 format(' magnitudes='12f10.2)
  sumnoc=0.
  do 1120 lm=1,lev
    sumnoc=sumnoc+nmoc(lm)
  1120 sumnoc=sumnoc/nmoc(lm)
c evaluate at twice as many magnitudes as initially read
  in if inos=1
    if(inos.ne.1) go to 1140
    dlev=(fm(2)-fm(1))/2.
    zacc=exp(abs(beta)*dlev)
    nmoc(1)=nmoc(1)*zacc/(1.+zacc)
    sumnoc2=nmoc(1)
    lev=2*lev
    fm(1)=fm(1)-dlev/2.
    do 1130 ll=2,lev
      fm(ll)=fm(ll-1)+dlev
      nmoc(ll)=nmoc(ll-1)/zacc
    1130 nmoc2=sumnoc2+nmoc(ll)
    write(16,1150) nmoc2, sumnoc2
    write(16,1100) (nmoc(l),l=1,lev)
    write(16,1110) (fm(l),l=1,lev)
  1140 continue
  1150 format(' total eqs='2e12.5)
  1160 continue
  1170 format(12f6.2)
c
  if (line,eq.0) go to 1250
c determine which rupture length magnitude relationship to use
  if(als,eq.0) go to 1240
  bl=als
  bl=bls
  sigl=sigl
c compute break length
  1180 write(16,1190) al,bl,sigl
  1190 format(' fault rupture length parameters al='f7.3' bl='f7.3'
    &,' sigl='f6.2)
  linset=1
  if (sigl.gt.0) go to 1230
  if which single break default do we want
  if bl=0, al non zero use old Algermissen-Perkins default values
  c use parameter values supplied for one break only
  1200 do 1210 i=1,lev
    1210 rlls(i,i)=10.*(al+bl*fm(i))
    1220 prls(i,i)=1.
    nrls=1
    go to 1350
  c set up rupture lengths and probabilities as fct of magnitude
  c nrls lengths per magnitude
  c al, bl, sigl non zero
  1230 call pbreak(fm,al,bl,sigl,rlls,prls,lev,nrls)
  go to 1350
  1240 if (nrls,eq.1) go to 1200
  if(linset,eq.1) go to 1230
  c compute default values
  al=-1.085
  bl=-.389
  sigl=.52
  linset=1
  go to 1180
c convert quad corner points to decimal degrees (if dsw =1)
  1250 do 1280 ii=1,num
    do 1280 i=1,2
      if(dsw,eq.0) go to 1260
      call condec(xsav(i,ii))
      call condec(ysav(i,ii))
  1260 continue
      xinx=xsav(i,ii)*rad
      yiny=xsav(i,ii)*rad
      call to eqr(xin,yin,rout,yout)
      xsav(i,ii)=rout*rad
      ysav(i,ii)=yout*rad
      write(16,1270) xsav(i,ii),ysav(i,ii),ii
  1270 format(' xsav,ysav='2f9.3' ii='i3)
  1280 continue
c
  1290 nm=num-1

```



```

1440 dtab(j,1)=rtab(mdis)
1450 go to 1500
1450 if(tadif(i,1).ne.0) go to 1490
c no change in acceleration at this distance
1460 if(i.eq.2) go to 1480
1470 write(16,1470) i,1,j
1480 format(' apparent error indist-table i,1,j='3i4)
1490 call exit
1480 dtab(j,1)=0.
1490 go to 1500
1490 dtab(j,1)=rtab(i)-rtab(i-1))*ata(i,1)-aclog(j))
&/atadif(i,1)
if(dtab(j,1).lt.0) dtab(j,1)=0.
1500 continue
1510 continue
1510 dlevsvdlev
1510 levsvdlev
1520 continue
1520 continue
c add half values of mag for later spreading out
dmag=fm(2)-fm(1)
bmag=fm(1)-dmag/2.
levp=lev+1
do 1530 i=1,levp
gm(i)=bmag
bmag=bmag+dmag
do 1570 il=1,levp
do 1540 il=2,jent
if(gm(il).gt.tm(il)) go to 1550
1540 continue
il=jent
do 1550 i=1,mdis
stah(j,1)=atab(j,il-1)+(gm(il)-tm(il-1))/tmdiv(il)
&*atab(j,il)-atab(j,il-1)
1560 continue
1570 continue
c take exp differences for later accel spread out
c the "exdif" values are used in subroutine "spread"
if (lev.eq.1) betad=2.0
if(lev.ne.1) betad=log(noc(1)/noc(2))
betabeta/dmag
fmul=1./1.-exp(-betadmag)
frak=exp(-betadmag/100.)
frak=frak
exdif(1)=0.
do 1580 i=2,101
exdif(i)=(1.-frak)*fmul
1580 frak=frak*frak
dtab=atab(1,1)
if(dlev.lt.0) dtab=dtab(1,lev)
c find entry for each magnitude in distance-acceleration table
c for zero distance (highest acceleration for that magnitude)
do 1600 l=1,lev
j=maxp
j=mxp
1590 if(dtab(j,1).ne.0.) go to 1600
j=j-1
go to 1590
1600 maxa(1)=j
c
irow1=irow1

```

```

1300 continue
c set line parameters for subsorce
1310 call setreg(ii)
do 1320 i=1,2
x(i)=xav(i,ii)
x(i+2)=xav(i,ii+1)
y(i)=yav(i,ii)
y(i+2)=yav(i,ii+1)
1320 y(i+2)=yav(i,ii+1)
c find area of subregion ii
s1=abs(x(i)*y(2)-y(3))+x(2)*y(3)-y(1))+x(3)*y(1)-y(2)))
s2=abs(x(i)*y(2)-y(3))+x(2)*y(3)-y(4))+x(3)*y(4)-y(2)))
sarea(ii)=(s1+s2)/2.
stot=stot+sarea(ii)
1330 continue
c
c find earthquake rate per unit area for lowest magnitude
c rate will be printed but not used in program
delm=(fm(2)-fm(1))/2.
fm1=fm(1)-delm
fm2=fm(1)+delm
eua=eua/stot
write(16,1340) dumid,stot,eua,fm1,fm2
1340 format(1h a4, 'area',f9.0, 'sq km, rate/sq km='e12.5,
&, ' for mags',f5.2,'-f5.2)
1350 dlev=fm(1)-fm(2)
c create table of log accelerations for set of magnitudes
c in (tm) for the original set of mdis distances as a function
c of distance
c next three lines test whether table has already
c been computed--if so skip over
if(lev.gt.levno) go to 1360
if(fm(1).ne.flevst) go to 1360
if(dlev.eq.dlevst) go to 1520
1360 do 1370 il=2,jent
1370 tmdif(il)=tm(il)-tm(il-1)
do 1420 il=1,lev
do 1380 il=2,jent
if(fm(il).gt.tm(il)) go to 1390
1380 continue
il=jent
1390 il=il-1
do 1400 j=1,mdis
ata(j,1)=atab(j,il)+(fm(il)-tm(il))/tmdiv(il)
&*atab(j,il)-atab(j,il))
do 1410 j=2,mdis
1410 atadif(j,1)=ata(j,1)-ata(j-1,1)
1420 lsub(1)=il
c find distance corresponding to each acceleration in
c table ac (set up in subroutine box) for each magnitude
do 1510 l=1,lev
do 1500 j=1,maxp
do 1430 i=2,mdis
if(aclog(j).gt.ata(i,1)) go to 1450
1430 continue
c
beyond table

```

```

row2s=row2
icol1s=icol1
icol2s=icol2
idrc=0
do 2770 ispt=1,2
  if(ispt.eq.1) go to 1610
  case of calculation done for single points on line
  if(indvpt.eq.0) go to 2770
  row1s=1
  row2s=1
  icol1s=1
  icol2s=indvpt
  go to 1620
1610 if(row1.eq.0) go to 2770
1620 do 2760 ira=row1s,row2s
  c
  ymid=(phi1-(ira-1.)*phinc)*r
  do 2750 ii=icol1s,icol2s
  c
  idrc=idrc+1
  if(ispt.eq.2) go to 1630
  fil=ii
  xmid=(fil-(ii-1.)*flinc)*r
  go to 1640
1630 xmid=xt2(ii)
  ymid=yt2(ii)
1640 if(isw.eq.0) go to 1660
  c
  after first time thru, read accelerations calculated
  c for previous sources at this site (saved in 2 dimensional
  c array raubin) into reg (1 dimensional) array.
  c
  do 1650 iqrp=1,maxdim
  1650 reg(iqrp)=raubin(iqrp,idrc)
  go to 1680
  c
  first time thru clear reg array if not continuation run
  c
  1660 do 1670 ll=1,maxp
  1670 reg(ll)=0.
  regg(idrc)=9999.999
  c set regg to min dist from fault
  c
  c for fixed affected area in row ira, col ii
  c
  c source area computation
  1680 if(line.eq.1) go to 2120
  do 2110 k=1,nm
  if(nbr.eq.0) go to 1700
  do 1690 iq=1,nbr
  if(k.eq.ibr(iq)) go to 2110
  1690 continue
  1700 approx=0.
  do 1710 i=1,2
  xll(i)=xsav(i,k)
  xll(i+2)=xsav(i,k+1)
  ql(i)=ysav(i,k)
  ql(i+2)=ysav(i,k+1)
  1710 continue
  r(3)=xll(1)

```

```

yr(3)=yl(1)
yr(1)=xl(2)
yr(1)=yl(2)
yr(2)=xl(4)
yr(2)=yl(4)
yr(4)=xl(3)
yr(4)=yl(3)
aread=area(k)/stot
c determine whether site is inside or outside k th quadrilateral
c subregion. find distance rc to closest endpoint
c and rf to farthest end point. Set up table of distance versus
c arc length for distances in the range rc to rf using
c subroutines outside or inside.
c itst=no of radii at fixed increments between rc and rf
c at which arc lengths are to be evaluated in table
c itst=7 set in subroutine risk--recompile to change
c call risk(xmid,ymid,in,k,dis,rc,rf,itst)
if(in.eq.1) go to 1740
if(rc.lt.dmax) go to 1730
do 1720 ll=1,lev
  1720 reg(ll)=reg(1)+aread*roc(ll)
  go to 2110
1730 rc=rc
  dial(1)=rc
  angl(1)=0.
  itst=itst+1
  ibeg=2
  go to 1750
1740 dial(1)=0.
  angl(1)=0
  dial(2)=rc
  angl(2)=pi2*rc
  ibeg=3
  if(rc.eq.0) ibeg=2
  rc=0.
  itst=itst+2
  c compute distance versus arc length table for even increments
  c in distance
  1750 do 1810 ii=ibeg,itst
  if(in.eq.0.) go to 1760
  call inside(xmid,ymid,k,dis,angle)
  go to 1780
  1760 call outside(xmid,ymid,k,dis,angle,kerr)
  1770 format('err outside for ira='i4' ii='i4)
  1780 continue
  dial(ii)=dis
  angl(ii)=angle+dis
  if(kerr.eq.0) go to 1810
  write (16,1770) ira,ii
  do 1790 imk1=1,ii
  1790 write(16,1800) dial(imk1),angl(imk1)
  1800 format(' dist, ang='2e12.5)
  1810 dis=dis+stsize
  itst=itst+1
  dial(itst)=r*phi
  angl(itst)=0.
  c compute distance versus arc length table for distances
  c to selected vertices and perpendiculars to edges
  rcsq=rc*rc
  rfsq=r*r

```

```

c acceleration a: ac(i-1) < a < ac(i) for each entry in acceleration
c table ac. equivalent to finding fractional source area at distance
c from site between successive entries in acceleration-distance
c table: dtab(i,1) < d < dtab(i+1,1) for magnitude 1.
c interpolates for distance is disl vs angl table
c area=integral(angle+dist*delte(dist)) from dtab(i-1,1) to dtab(i,1)
do 2090 ll=1,lev
stnoc=ncoc(11)/stot
isud=0
maxm=maxa(11)
if(maxm.le.0.) go to 1960
approx=0.
if(in.eq.0) go to 1950
jlo=maxm
go to 1980
1950 if(rcs.lt.dtab(1,1)) go to 1970
1960 reg(1)=reg(1)+aread*ncoc(11)
go to 2080
1970 jlo=indpt(dtab(1,1),rcs,maxm)
1980 jlo=jlo
iq=2
do 2060 jk=1,jlo
if(jk.gt.1) go to 1990
dbot=rcs
ddis=dtab(j,11)
if(dtab(j,11).lt.rf) goto 2010
reg(j+1)=reg(j+1)+sarea(k)*stnoc
go to 2080
1990 if(dtab(j,11).lt.rf) go to 2000
dbot=dtab(j+1,11)
ddis=rf
isud=1
go to 2010
2000 ddis=dtab(j,11)
dbot=dtab(j+1,11)
anarea=0.
do 2020 iq=1,itst
if(dbot.le.disl(iq)) go to 2030
2020 continue
iq=itst
distance exceeds last value-error since table set up to handle this
2030 dtop=dsl(iq)
if(dsl(iq).gt.ddis) dtop=ddis
dmid=(dtop+dbot)/2.
deir=dtop-dbot
fr=(dmid-dsl(iq-1))/(dsl(iq)-dsl(iq-1))
ang=angl(iq-1)+(angl(iq)-angl(iq-1))*fr
anarea=anarea+ang*deir
if(dtop.eq.ddis) go to 2050
dbot=dtop
iq=iqt1
if(iq.le.itst) go to 2030
write(16,2040)
2040 format(' err in arc area computation')
go to 2030
c interpolate in angle , distance table
2050 approx=approx+anarea
reg(j+1)=reg(j+1)+anarea*stnoc
jm=j-1
if(isud.eq.0) go to 2060

```

```

do 1940 i=1,4
ire=0
if(in.eq.0) ire=2
if(sllq(i).le.rcsq(1.)) go to 1930
if(sllq(i).ge.rfsq(1.)) go to 1930
dis=stc(sllq(i))-001
if(dis.le.rc) goto 1930
add angle, distance corresponding to vertex
if (in.eq.0) go to 1830
call inside(xmid,ymid,k,dis,angle)
go to 1860
1820 if(mina(i).ne.3) go to 1940
c closest side next
dis=abs(perpa(i))
if(dis.le.rc(3)) go to 1940
if(dis.ge.rf(3)) go to 1940
dis=dis-.001
call inside(xmid,ymid,k,dis,angle)
ire=1
go to 1860
1830 continue
call outsid(xmid,ymid,k,dis,angle,kerr)
go to 1860
1840 if(mina(i).ne.3) go to 1940
dis=abs(perpa(i))
ire=1
do 1850 iql=1,4
if(abs(sllq(iql)-perpaq(i)).le.2.) go to 1940
1850 continue
if(dis.le.rc(1.)) go to 1940
if(dis.ge.rf(1.)) go to 1940
dis=dis-.001
call outsid(xmid,ymid,k,dis,angle,kerr)
1860 continue
if(kerr.eq.0) go to 1880
write (16,1770) ira,ii
do 1870 imk1,iii
1870 write(16,1800) disl(imk1),angl(imk1)
1880 continue
do 1890 ll=1,itst
if(dis-dsl(11)) 1900,1920,1890
1890 continue
c insert new distance, angle
1900 jk=itst
do 1910 mm=1,itst
angl(jk+1)=angl(jk)
disl(jk+1)=disl(jk)
1910 jk=jk-1
disl(11)=dis
angl(11)=angle*dis
itst=itst-1
1920 if(ire.eq.0) go to 1820
if(ire.eq.2) go to 1840
go to 1940
1930 if(in)1840,1840,1820
1940 continue
itst=itst+1
disl(itst)=rf
angl(itst)=0
c for each magnitude, source, determine fractional source area giving

```

```

      arerr=(approx-sarea(k))/sarea(k)
      allerr=allerr+1.
      if (abs(arerr).lt. .05) go to 2090
      tot5=tot5+1.
      if(abs(arerr).le. .10) go to 2090
      write(16,2100) arerr,ire,ii,rc,r,f,k,ll,in,dumid
      go to 2090
2060  iqs=iq
      c need to add remaining area if any to lowest acceleration
      dif=sarea(k)-approx
      if(dif.lt.0) go to 2080
2070  reg(1)=reg(1)+dif*stnoc
2080  approx=sarea(k)
2090  continue
2100  format(' int err='f6.3' row 'i2' col 'i2' rc='
      & f7.2' rf='f7.2' k='i2' ll='i2' in='i1,ix,a4')
2110  continue
      go to 2730
      c fault computation
2120  do 2720 j=1,jtot
      jseg=jseg+1
      jseg=jseg-1
      call cidis(imid,ymid,j)
      do 2160 ll=1,lev
      c fill in rupture lengths--make certain that rupture length
      c does not exceed total fault length
      nrl=nrls
      do 2130 ir1=1,nrl
      rll(ir1,11)=rlls(ir1,11)
2130  prl(ir1,11)=prls(ir1)
      if(rll(1,11).lt.dtot(j)) go to 2160
      rll(1,11)=dtot(j)
      if(nrl.eq.1) go to 2160
2140  if(rll(2,11).lt.dtot(j)) go to 2160
      prl(1,11)=prl(1,11)+prl(2,11)
      nrl=nrl-1
      if(nrl.eq.1) go to 2160
      do 2150 ir2=1,nrl
      prl(ir2,11)=prl(ir2,11)
2150  rll(1,11)=rll(1,11)+prl(ir2,11)
      go to 2140
2160  nrlsv(11)=nrl
      pt1=sqrt(shdis)
      if(pt1.lt.regg(idrc)) regg(idrc)=pt1
      if (isid.eq.0) go to 2490
      c distance is monotonically increasing for all segments
      c (the decreasing case has been converted to increasing)
      c or single shortest distance is interior to fault
      do 2480 ll=1,lev
      nrl=nrlsv(11)
      frcon=dtot(j)/dsunmnc(11)
      maxm=mxax(11)
      jlo=indat(dtab(1,11),pt1,maxm)
      if (jlo.gt.0) go to 2170
      c distance beyond table--lump acceleration in first box
      reg(1)=reg(1)+frcon
      go to 2480
2170  continue
      if (isid.eq.1) go to 2190
2180  ibeg=1

```

```

      ifin=jsegm
      ixt=1
      dnear=dtot(j)
      inear=jsegm
      go to 2310
2190  ibeg=1
      ifin=inear
      ixt=2
      c dnear=distance along fault from one end to point on fault
      c closest to the site
      c dnear=total length of segments 1,2,...inear
      c inear defined sub cldis--in common/extra/
      dnear=0
      do 2200 iq=1,inear
      dnear=dnear+distl(iq)
      dfer=dtot(j)-dnear
      totac=0.
      do 2220 ir1=1,nrl
      smd=rll(ir1,11)
      if(smd.lt.dtot(j)) go to 2210
      c break is entire segment
      totac=totac+frcon*prl(ir1,11)
      sumr=totac
      go to 2220
2210  smd2=smd/2.
      c find fraction of fault at closest distance to site
      c for interior site
      arcn=smd2
      arcf=smd2
      if(dnear.lt.smd) arcn=dnear-smd2
      if(dfer.lt.smd) arcf=dfer-smd2
      arct=arcn+arcf
      if(arct.lt. .0001) go to 2220
      totac=totac+frcon/(dtot(j)-smd)*arct*prl(ir1,11)
2220  continue
      sumr=totac
      totac=totacs+totac
      jlop=jlop+1
      if(totl.gt.0) go to 2230
      c spread dummy faults over "totl" km distance if "totl" > 0
      c if "totl" = 0, spread over magnitude instead
      call spread(reg,jlop,11,totac,pt1)
      go to 2310
2230  width=totl/2.
      sumr=sumr/totl
      add=0
      if(abs(pt1).gt.width) go to 2280
      qtot=width*abs(pt1)
      dsvs=0.
      jlo2=mxm
      do 2270 iqt=1,2
      jrp=jlo2
      do 2250 jk=1,jlo2
      d13=dtab(jrp,11)
      if(d13.gt.qtot) go to 2260
      cary=d13-dsvs
      reg(jrp+1)=reg(jrp+1)+sumr*cary
      add=add+sumr*cary
      jrp=jrp-1
2240  format(' add='e12.5)

```



```

2250 davs=d13
2260 jrp=0
2260 cary=qtot-davs
2260 add=dd+sum*cary
2260 jrp=jrp+1
2260 reg(jrp+1)=reg(jrp+1)+sum*cary
2260 davs=0
2270 qtot=width-abs(pt1)
2270 jrp=jlo2
2270 jrp=jrp+1
2280 go to 2310
2280 dstart=abs(pt1)-width
2280 qtot=width-dstart
2280 jlo2=indpt(dtab(i,11),dstart,maxm)
2280 davs=dstart
2280 jrp=jlo2
2280 do 2290 jk=1,jlo2
2280 d13=dtab(jrp,11)
2280 if(d13.gt.qtot) go to 2300
2280 cary=d13-davs
2280 reg(jrp+1)=reg(jrp+1)+sum*cary
2280 add=dd+sum*cary
2290 jrp=jrp+1
2290 davs=d13
2300 jrp=0
2300 cary=qtot-davs
2300 reg(jrp+1)=reg(jrp+1)+sum*cary
2300 c find frac of fault in distance range dtab(jr,11) < d < dtab(jr+1,11)
2300 for each magnitude li
2310 do 2470 i thru i,ixt
2310 jlo=jlo
2310 ilo=1
2310 ihi=ibeg
2310 casum=0
2310 ir=1
2310 dsave=dls(iii)
2320 frac=0
2320 jk=1
2320 smd=r11(ir,11)
2320 if(smd.lt.dtot(j)) go to 2330
2320 if(ist.eq.2) go to 2450
2320 add in entire fault length
2320 totacc=frcon*pr11(ir,11)
2320 jlo=jlo+1
2320 call spread(reg,jlo,11,totacc,pt1)
2320 go to 2450
2330 cuml=dnear-smd
2330 if(cuml.le.0) go to 2450
2330 do 2340 i=1,nr1
2330 frac=frac*pr11(i,11)/(dtot(j)-r11(i,11))
2330 frac=frac*frcon
2330 dlen=0
2330 do 2350 i=ibeg,ifin
2330 drem(iq)=disti(iq)
2330 dlen=dlen+drem(iq)
2330 if (dlen.ge.cuml) go to 2370
2330 continue
2350 write(16,2360)dlen,ibeg,ifin,cuml
2360 format(' error for fault line dlen=e12.5' ibeg='13
& ifin='13' cuml=e12.5)
2370
2370 call exit
2370 distem(iq)=cuml-dlen+drem(iq)
2380 ifins=iq
2390 dend=distem(iii)-dls(iii)
2390 dsave=starting distance along segment iii
2400 continue
2410 jk=ilo,jlo
2410 d13=sqrt(dtab(jr,11)**2-persq(iii))
2410 if(d13.gt.dend) go to 2440
2410 carea=d13-dsav
2410 casum=casum+carea
2410 reg(jr+1)=reg(jr+1)+carea*frac
2410 jrp=jr-1
2410 dsave=d13
2420 at top of table---remaining dist beyond table
2420 do 2430 i=r11,nr1
2420 cuml=dnear-r11(i,11)
2430 reg(i)=reg(i)+(cuml-casum)*pr11(iq,11)*frcon
2430 & /(dtot(j)-r11(iq,11))
2430 go to 2460
2440 carea=dend-dsav
2440 casum=casum+carea
2440 reg(jr+1)=reg(jr+1)+carea*frac
2440 dsave=dend
2440 c at end of segment iii
2440 if(iii.eq.ifins) go to 2450
2440 ihi=ihl+1
2440 dsave=dls(iii)
2440 ilo=jk
2440 go to 2390
2450 ir=ir+1
2450 ilo=jk
2450 if(ir.le.nr1) go to 2320
2450 if(i thru.eq.ixt) go to 2480
2460 ir=1
2460 ibeg=inear+1
2460 ifin=jsegm
2460 dnear=dtot(j)-dnear
2470 continue
2480 go to 2720
2480 c more than one turning point
2480 c brute force. find contribution for single break
2480 c increment break centers by distance ddel along fault from
2480 c one end to other.
2490 continue
2490 ddel=10.
2490 if(pt1.gt.100) ddel=20.
2490 if(pt1.lt.15.) ddel=5.
2490 do 2710 il=1,lev
2490 nr1=nr1sv(il)
2490 do 2710 ir=1,nr1
2490 smd=r11(ir,11)
2490 pntoc=noc(11)*pr11(ir,11)
2490 d=dtot(j)-smd
2490 nd=dj/ddel+1.5
2490 if(nd.gt.1) go to 2500
2490 nd=2
2490 if(dj.gt.1.) go to 2500
2490 break is entire fault length

```



```

      eqsum=0.
      dumfac=1.
      c compute icon based on sigma and dphinc
      c probably enough to go out 2 sigma on either side of point
      c find how many rows or columns that corresponds to
      icon=2*sl/dphinc+1.01
      c if(icon.lt.5) icon=5
      c if icon<5, set icon=5
      irm=irow2-nbord
      icm=icol2-nbord
      isub1=irow1+nbord
      isub2=icol1+nbord
      do 2830 iia=isub2,icm
      j1=iia-icon
      j2=iia+icon
      if(j2.gt.icol2) j2=icol2
      if(j1.lt.icol1) j1=icol1
      do 2830 iras=isub1,irm
      i1=iras-icon
      i2=iras+icon
      if(i1.lt.irow1) i1=irow1
      i2=irow1+icon
      if(i2.gt.irow2) i2=irow2
      do 2800 iqr=i1,maxp
      add=0.
      do 2790 i=1,i2
      wtt=0.
      do 2790 j=1,j2
      ksq=(iras-i1)*2+(iia-j1)*2
      if(ksq.gt.imaxsq) go to 2790
      c we need to convert to proper subscript in rabin
      c in rabin, first all cols for one row
      c (i-1)*icol(s)+jrow and col with single subscript in rabin
      isub=((i-irow1)*icol+1-j-icol1+1
      add=add+fac(ksq)*rabin(iq,isub)
      wtt=wtt+fac(ksq)
      2790 continue
      reg(iq)=0
      if(wtt.ne.0) reg(iq)=add/wtt
      eqsum=eqsum+reg(iq)
      2800 continue
      2810 continue
      isub=((iras-irow1)*icol+1+iia-icol1+1
      do 2820 iqr=i1,maxdim
      smobin(iq,isub)=reg(iq)+smobin(iq,isub)
      2830 continue
      c now delete entries from rabin array
      do 2840 isub=i,nrc
      do 2840 iqr=i,maxdim
      rabin(iq,isub)=0.
      2850 continue
      if(iprint.lt.0) go to 2940
      c output all results to this point
      do 2930 ispt=i,2
      if(ispt.eq.1) go to 2860
      if(indvpt.eq.0) go to 2930
      irow1s=i
      irow2s=i
      icol1s=i
      icol2s=indvpt
      go to 2870
      2860 if(irow1.eq.0) go to 2930
      irow1s=irow1+nbord
      irow2s=irow2-nbord
      icol1s=icol1+nbord
      icol2s=icol2-nbord
      2870 do 2920 ira=irow1s,irow2s
      do 2920 i1=icol1s,icol2s
      if(ispt.eq.1) go to 2890
      xouts=xpt1(ii)
      youts=ypt1(ii)
      isub=irows*icol+ii
      go to 2900
      c back transform point from equator to dec degrees
      2890 call backe(ii,ira,xouts,youts)
      c results calculated and printed in subroutine out
      isub=(ira-irow1)*icol+ii-icol1+1
      2900 do 2910 iqr=i,maxp
      2910 reg(iqr)=rabin(iqr,isub)+smobin(iqr,isub)
      rdis=reg(isub)
      if(iprint.le.2) write(16,30) rdis
      call out(reg,naf,xouts,youts,sd)
      2920 continue
      2930 continue
      2940 continue
      isw=1
      go to 680
      c finished
      2950 nvals=icolslstirout+indvpt
      c write to unit 3 for possible run continuation
      c may delete to save time and space if no run
      c continuations are desired.
      do 2970 i=1,nvals
      do 2960 iqr=i,maxdim
      2960 reg(iqr)=rabin(iqr,i)+smobin(iqr,i)
      write(3) reg
      2970 continue
      c transfer to unit 3 for possible continuation or restart
      2980 continue
      2990 continue
      c write(16,2930) tot5,allerr
      3000 format(/' sources with error in area > .05='#7.0
      &' out of '#7.0' sources ')
      3010 format(' number of brute force points='i4)
      stop
      end
      subroutine cond(x1,x2,y1,y2)
      call condec(x1)
      call condec(x2)
      call condec(y1)
      call condec(y2)
      return
      end
      subroutine condec(phi)
      i=phi
      phi=float(i)+(phi-float(i))/.60
      return
      end
      c *****

```

```

c subroutine lin(x1,y1,x2,y2,a,b,c,x3,y3)
c coefficients a, b, c are determined so that points which are in the
c same half plane as x3,y3 will be a positive distance from the line
c joining x1,y1 and x2,y2
c
c   a x + b y + c = 0
c   if(x1.eq.x2) go to 10
c   a=(y1-y2)/(x1-x2)
c   b=-1
c   c=y1-ax1
c   go to 20
10  a=1.
c   b=0.
c   c=-x1
20  d=ax3+by3+c
c   if(d.gt.0.) return
c   a=-a
c   b=-b
c   c=-c
c   return
c
c *****
c subroutine linint(reg,t,sol)
c calculates acceleration "sol" that has probability
c "prob" of not being exceeded during "t" years at
c site.
c expected yearly accelerations in the range
c ac(j-1) < a < ac(j) are in reg(j).
c linint is called from subroutine out.
c common/tims/dummy(42),iprint
c common/ext/prob,ex,ex1
c dimension reg(106)
c i=maxp-1
c sum=0.
10  if(i1.le.2) go to 20
c   if(reg(i1).gt.0.) go to 40
c   i=i1-1
c   go to 10
20  sol=0.
c   if(iprint.le.2) write(16,30)
30  format(' sol not obtained for time='f7.0)
c   return
40  sumsum=reg(i1)
c   if(tsum.gt.ex) go to 50
c   i=i1-1
c   if(i1.ge.2) go to 40
c   go to 20
50  if(sum.eq.reg(i1)) go to 70
c   y2=log(t*(sum-reg(i1)))
c   y1=log(t*sum)
c   sol=ex1
60  a=(y1-y2)/(ac(i1)-ac(i1))
c   b=y1-ac(i1-1)
c   sol=(ex1-b)/a
c   return
70  y1=t*sum
c   y2=0.
c   sol=ex1
80  format(' log interpolation fails---use linear')

```

```

c *****
c subroutine box(scale,sd)
c determine boundaries ac(i) for acceleration levels
c accelerations in the range ac(i-1) < ac(i) will be
c accumulated in reg(i)
c scale multiplies the basic levels computed
c scale=1 for accelerations in the range .02 to 1.
c common /quad/maxq,ac(105),dumlog,aclog(105)
c common/wds/maxw,maxp,maxsp
c maxsp=104
c max=54
c maxp=55
c ac(5)=.02*scale
c dif=ac(5)
c do 10 i=6,55
c   ac(i)=ac(i-1)+dif
10  continue
c
c -----note important-----
c the following inserted to allow acceleration boxes for small
c values below range of interest for sd computation
c this should be changed later as desired
c
c   j=4
c   efac=exp(-sd/2.)
c   do 20 i=1,4
c     ac(j)=ac(j+1)*efac
20  j=j-1
c
c for later attenuation variability, expand acceleration range
c used only in subroutine "out"
c   y=alog(ac(max))
c   w=exp(y+4.4*sd)
c   w1=exp(y+2.0*sd)
c   dimension of ac, aclog is 105, we cannot exceed this dimension
c   n=maxsp-max
c   n2=n/2
c   i2s=max+n2
c   dif=(w1-ac(max))/n2
c   do 30 i=maxp,i2s
c     ac(i)=ac(i-1)+dif
30  continue
c   i2sp=i2s+1
c   dif=(w-ac(i2s))/n2
c   do 40 i=i2sp,maxsp
c     ac(i)=ac(i-1)+dif
40  continue
c   return
c
c *****
c subroutine inieqr(x1,y1,x2,y2)
c read along, ylat in radians for 1st point
c read along, ylat in radians for 2nd point
c points 1 and 2 to be transformed to equator
c point 1 will be transformed to (0,0)
c point 2 will be transformed to (d,0)
c where d=great circle distance between (x1,y1) and (x2,y2)
c output is transformation matrices f (forward) f1 (backward)
c generated and stored in common
c common/eqr/ r1(3,3),r2(3,3),r3(3,3),d(3,3),f(3,3),vec(3)
c common/inv/f1(3,3),res(3)

```



```

cpi=cos(y1)
cp2=cos(y2)
cli=cos(x1)
spi=sin(y1)
sp2=sin(y2)
sli=sin(x1)

c initialize matrices
do 10 i=1,3
  do 10 j=1,3
    r1(i,j)=0.
  r2(i,j)=0.
  r3(i,j)=0.
  r1(i,1)=1.
  r2(2,2)=1.
  r3(3,3)=1.
  r2(1,1)=cp1
  r2(3,3)=cp1
  r2(1,3)=sp1
  r2(3,1)=sp1
  r3(1,1)=cl1
  r3(2,2)=cl1
  r3(1,2)=sli
  r3(2,1)=sli
  r3(2,1)=sli
  r3(2,1)=sli
  s1=cp1*sp2-sp1*cp2*cos(x2-x1)
  c1=cp2*sin(x2-x1)
  den=sqrt(c1*c1+s1*s1)
  sinth=s1/den
  cothc1=den
  r1(2,2)=cothc1
  r1(3,3)=cothc1
  r1(2,3)=sinth
  r1(3,2)=-sinth
  r1(3,2)=-sinth
  call matmpy(r1,r2,d)
  call matmpy(d,r3,f)

c transformation matrix in f! inverse matrix in f1
c compute inverse
do 20 i=1,3
  do 20 j=1,3
    r1(i,j)=0.
  r2(i,j)=0.
  r3(i,j)=0.
20 continue
  r1(3,3)=1.
  r2(2,2)=1.
  r3(1,1)=1.
  r1(1,1)=cl1
  r1(2,2)=cl1
  r1(2,1)=sli
  r1(1,2)=-sli
  r2(1,1)=cp1
  r2(3,3)=cp1
  r2(1,3)=sp1
  r2(3,1)=sp1
  r3(2,2)=cothc1
  r3(3,3)=cothc1
  r3(3,2)=sinth
  r3(2,3)=-sinth
  call matmpy(r1,r2,d)
  call matmpy(d,r3,f1)
return

c *****
end
subroutine toeq(r3,y3,flam6,phi6)
c transform point at (x3,y3) (long,lat in radians) to new
c equator using transformation defined in subroutine ineqr
common/eq/ r1(3,3),r2(3,3),r3(3,3),d(3,3),f(3,3),vec(3)
common/inv/f1(3,3),res(3)
c mid=cos(x3)
c mid=cos(y3)
c mid=sin(x3)
c mid=sin(y3)
vec(1)=cxmid*cyamid
vec(2)=sxmid*cyamid
vec(3)=syamid
call vecmpy(f,vec,res)
phi6=asin(res(3))
cp6=cos(phi6)
cl6=sin(phi6)
s16=res(1)/cp6
s16=res(2)/cp6
if(cl6.gt.1.) cl6=1.
flam6=acos(cl6)
if(s16.lt.0.) flam6=-flam6
return
end
c *****
subroutine backw(ix,iy,fl1,phi1)
c for printout and identification: point defined by
c row ix and col iy in new coordinate system is transformed
c to location (long, lat, in dec degrees) in original
c coordinate system.
common/inv/f1(3,3),res(3)
dimension vec(3)
c need cos, sins in common array
common/sincos/sinx(500),cosx(500),siny(500),cosy(500)
data rad/.0174533/
s16=sin(ix)
cl6=cos(ix)
cp6=cos(iy)
sp6=sin(iy)
vec(1)=cl6*cp6
vec(2)=s16*cp6
vec(3)=sp6
call vecmpy(f1,vec,res)
if(abs(res(3)).lt.1.0) go to 30
if(res(3).ge.1.) go to 10
phi1=90.
go to 20
10 phi1=90.
20 flam6=0.
return
30 phi1=asin(res(3))
cpi=cos(phi1)
phi1=phi1/rad
cli=phi1/rad
sli=res(1)/cpi
sli=res(2)/cpi
fli=acos(c1i)
if(sli.lt.0.) fli=-fli
phi1=fli/rad
return
end

```

```

c *****
c      subroutine matmpy(a,b,c)
c      multiply two 3 x 3 matrices A x B = C
c      dimension a(3,3),b(3,3),c(3,3)
c      do 10 i=1,3
c      do 10 j=1,3
c      c(i,j)=0.
c      do 10 k=1,3
c      c(i,j)=c(i,j)+a(i,k)*b(k,j)
c      return
c      end
c *****
c      subroutine vecmpy(f,vec,r)
c      vector multiplication used by subroutines backw, toeqr
c      dimension f(3,3),vec(3),r(3)
c      do 10 i=1,3
c      r(i)=0.
c      do 10 k=1,3
c      r(i)=r(i)+f(i,k)*vec(k)
c      return
c      end
c *****
c      subroutine csqdis(x,xi,yi,xr,yr,isub)
c      for a quadrilateral source area, calculates
c      nearest distance from site at (x,y) to each side
c      and distance to each vertex.
c      called from subroutine risk
c      dimension xl(4),yl(4),xr(4),yr(4)
c      common/slmp/fsa(4,50),fsb(4,50),fsc(4,50),rta(4,50),dist(4,50)
c      &,disq(4,50)
c      common /slnds/dls(4),perp(4),sl(4),slsq(4),min(4),persq(4)
c      &,rsq(4),dip(4)
c      do 10 i=1,4
c      perp(i)=(fsa(i, isub)*x+fsb(i, isub)*y+fsc(i, isub))/rta(i, isub)
c      do 20 i=1,4
c      slsq(i)=(xi-xl(i))*2*(y-yl(i))*2
c      rsq(i)=(xi-xr(i))*2*(y-yr(i))*2
c      20 continue
c      do 50 i=1,4
c      persq(i)=perp(i)*perp(i)
c      dl1=slsq(i)-persq(i)
c      dl2=rsq(i)-persq(i)
c      if(dl1.lt.0.) dl1=0.
c      if(dl2.lt.0.) dl2=0.
c      if(dl1+dl2.le.disq(i, isub)) go to 40
c      shortest distance external to segment
c      if(dl1.lt. dl2) go to 30
c      min(i)=1
c      dip(i)=sqrt(dl2)
c      dls(i)=dist(i, isub)+dip(i)
c      go to 50
c      30 min(i)=2
c      dis(i)=sqrt(dl1)
c      go to 50
c      40 min(i)=3
c      dis(i)=sqrt(dl1)
c      dip(i)=0.
c      50 continue
c      return
c      end

```

```

c *****
subroutine rrist(xnot,ynot,in,sub,r,rc,rp,ist)
common/inout/ic,ics,ifar,azimf,ssize
common/xyarea/ysav(2,50),ysav(2,50),sarea(50)
common/ylinds/dls(4),perp(4),sl(4),slsq(4),min(4),persq(4)
&,sraq(4),dip(4)
&,common/slinp/aa(4,50),bb(4,50),cc(4,50),rta(4,50),dist(4,50),
&,disq(4,50)
do 10 i=1,2
xl(i)=xsav(i,sub)
xl(i+2)=xsav(i,sub)
yl(i)=ysav(i,sub)
yl(i+2)=ysav(i,sub)
10 continue
xr(3)=xl(1)
yr(3)=yl(1)
xr(1)=xl(2)
yr(1)=yl(2)
xr(2)=xl(4)
yr(2)=yl(4)
xr(4)=xl(3)
yr(4)=yl(3)
c determine if any sides are vertical lines
do 70 li=1,4
if(xl(li).eq.xr(li)) go to 20
aa(ii,sub)=yl(ii)-yr(ii)/(xl(ii)-xr(ii))
bb(ii,sub)=-1.
cc(ii,sub)=yl(ii)-aa(ii,sub)*xl(ii)
go to 30
20 aa(ii,sub)=1.
bb(ii,sub)=0.
cc(ii,sub)=xl(ii)
30 if(ii.gt.2) go to 40
d=aa(ii,sub)*xl(3)+bb(ii,sub)*yl(3)+cc(ii,sub)
go to 50
40 d=aa(ii,sub)*xl(2)+bb(ii,sub)*yl(2)+cc(ii,sub)
50 if (d.gt.0.) go to 60
aa(ii,sub)=-aa(ii,sub)
bb(ii,sub)=-bb(ii,sub)
cc(ii,sub)=-cc(ii,sub)
60 disq(ii,sub)=(xr(ii)-xl(ii))*2+(yr(ii)-yl(ii))*2
dist(ii,sub)=sqrt(disq(ii,sub))
70 rta(ii,sub)=sqrt(aa(ii,sub)*aa(ii,sub)+bb(ii,sub)*
&bb(ii,sub))
return
end
function indpt(y,uval,n)
c points in descending order--non zero values from points
c 1 to n: find first value in table exceeding yval
dimension y(102)
if(yval.lt.y(1)) go to 10
indpt=0
return
10 if(yval.gt.y(n)) go to 20
indpt=n
return
20 num=
idiv=2
30 nh=num/idiv
if(nh.le.1) go to 50
if(yval.lt.y(nh)) go to 40

```

```

      idiv=2*idiv
      go to 30
    c      yval lies approx between nh and 2*nh or 2*nh
    40 idiv=2*idiv
      nh2=nh+num/idiv
      if(nh.eq.nh2) goto 80
      if(yval.gt.y(nh2)) go to 40
      yval further down on table than y(nh2)
      nh=nh2
    c      go to 40
    50 do 40 i=1,n
      if(y(i).lt.yval) go to 70
    60 continue
    70 indpt=i-1
      return
    80 continue
      do 90 i=nh,n
      if(y(i).lt.yval) go to 70
    90 continue
  end
  c *****
  c subroutine cldis(x,y,jl)
  c containing fault segment i of j1 th fault (for each i)
  c determines whether closest point to fault is at
  c low end (min(i)=2)
  c interior (min(i)=3)
  c high end (min(i)=1)
  c set isid=0 if multiple turning points
  c isid=1 if single closest distance from site to fault
  c interior to fault
  c isid=2 if closest distance from site to fault at one end
  c of the fault
  c dimension dlp(20)
  common/linps/distl(24,26),disq(24,26),jseg,jsegm,
& x1(24,26),yp(24,26),coa(24,26),sia(24,26)
  common/extra/isid,shdis,dist(24),linear
  common /lindis/dis(24),perp(24),min(24),persq(24)
  shdis=1.e9
  do 10 i=i1,jsegm
    dist(i)=distl(i,jl)
    i1=0
    i2=0
    i3=0
    isid=0
    do 20 i=i1,jseg
      alsq=(x-x1(i,jl))*2+(y-yp(i,jl))*2
      if(alsq.lt.shdis) shdis=alsq
    20 continue
      do 30 i=i1,jsegm
        xd=x-x1(i,jl)
        yd=y-yp(i,jl)
        perp(i)=xd*sia(i,jl)+yd*coa(i,jl)
        dist(i)=xd*coa(i,jl)-yd*sia(i,jl)
        persq(i)=perp(i)*perp(i)
        if(dls(i).lt.0) go to 30
        if(dls(i).le.dist(i)) go to 40
        shortest distance external to segment
        min(i)=1
        i1=i1+1
      30 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=dist(k)+1-dls(k)
        dist(k)=dls(k)
        dls(k+1)=0
      40 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      50 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      60 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      70 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      80 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      90 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      100 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      110 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      120 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      130 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      140 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      150 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      160 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      170 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      180 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      190 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      200 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      210 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      220 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      230 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      240 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      250 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      260 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      270 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      280 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      290 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      300 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      310 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      320 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      330 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      340 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      350 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      360 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      370 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      380 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      390 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      400 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      410 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      420 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      430 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      440 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      450 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      460 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      470 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      480 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      490 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      500 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      510 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      520 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      530 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      540 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      550 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      560 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      570 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      580 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      590 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      600 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      610 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      620 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      630 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      640 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      650 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      660 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      670 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      680 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      690 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      700 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      710 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      720 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      730 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      740 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      750 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      760 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      770 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      780 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      790 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      800 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      810 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      820 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      830 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      840 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      850 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      860 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      870 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      880 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      890 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      900 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      910 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      920 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      930 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      940 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      950 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      960 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      970 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      980 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
      990 i=i1,jsegm
        jseg=jseg+1
        dist(k+1)=0
    c      we have single turning point
    isid=1
    linear=k-1
    if(min(k).ne.3) go to 140
    linear=k
    adjust to 2 sep segments
    j=jseg
    do 130 i=k,jsegm
      dls(j)=dls(j-1)
      persq(j)=persq(j-1)
      dist(j)=dist(j-1)
      j=j-1
    130 j=j-1
    jseg=jseg
    jseg=jseg+1
    dist(k+1)=dist(k)+1-dls(k)
    dist(k)=dls(k)
    dls(k+1)=0
  end

```



```

140 iup=linear
   go to 60
end
c *****
c subroutine pbreak(fm,al,bl,sl,prls,prls,lev,nrls)
c dimension fm(12),prls(12),rlls(10,12)
c dimension pmls(5),prang(5)
c data pmls/1.575,.74,0.,-.74,-1.575/
c data prang/1151.,2295.,3108.,2295.,1151/
c data nrls/5/
c evaluate al+bl*fm(1)+pmls(i)*sigl
c where i<1<lev and i<1<nrl
c prang probability associated with jth break
c for each magnitude; compute nrl rupture lengths
c : : : : : must be ordered in descending order---largest break
c : : : : : length first for each magnitude
c do 10 j=1,lev
c   conval + bl*fm(j)
c   do 10 i=1,nrl
c     c=con+pmls(i)*sigl
c     10 rlls(i,j)=10.**c
c     do 20 i=1,nrl
c       prls(i)=prang(i)
c       nrls=nrl
c     return
c   end
c *****
c subroutine bbgau(x,gg)
c computes normal probability integral gg from -(infinity) to xx
c sets gg=1 if xx<-6, gg=0 if xx>+6
c uses approximation Handbook of Mathematical Functions-
c NBS Applied Math Series 55--p299 7.1.26 for error function
c ax=abs(x)/1.4142136
c if(ax<4.2426408) 20,20,10
c 10 gg=0.
c   if(x.lt.0) gg=-x1.
c   return
c 20 continue
c   d=1.+(((((4.30638e-5*ax+2.765672e-4)*ax+1.520143e-4
c   & )*ax+.2705272e-3)*ax+4.2282012e-2)*ax+.0523078e-2)*ax
c   d=d*d
c   d=d*d
c   d=d*d
c   gg=.5/(d*d)
c   if(x) 30,30,40
c 30 gg=1.-gg
c 40 return
c end
c *****
c subroutine outsid(ynot,ynot,isub,r,angle)
c common/inout/ic,ics,ifac,aximf,stime
c common/xyav/xyav(2,50),yav(2,50),sarea(50)
c common/xy1/xy1(4),yl(4),xr(4),yc(4),yc(4)
c common/slip / am(4,50),bb(4,50),cc(4,50),rta(4,50),dist(4,50)
c &,disq(4,50)
c common/slinds/dls(4),perp(4),sl(4),slsq(4),min(4),persq(4)
c &,rsq(4),dls(4)
c dimension jsub(3)
c subroutine for calculating angle when site is outside
c   (quadrilateral) source area.

```

```

90 d=(dis(ii)-sqrt(rsq-persq(ii)))/dist(ii, isub)
c single intersection
c go to 60
c see if this side is closest to point, if so, treat specially.
100 if (ii-ics) i10,i20,i10
c both points are on boundary, calculate angle between them.
110 sign=-1
c go to 130
120 sign=1.
130 angle=sign*2.*as + angle
c see if second point only is on boundary
140 continue
c if(npt.gt.0) go to 150
c if(signal) 310,320,320
150 go to (320,240,160,250),npt
c this is an error unless radius is on a vertex and
c was counted twice.
160 do 170 i=1,3
c j=sub(i)
c if(abs(xc(i)-xc(j)) .ge. .01) go to 170
c if(abs(yc(i)-yc(j)) .le. .01) go to 220
170 continue
c check for one of three points being a vertex
do 210 i=1,4
do 180 j=1,3
c if(abs(xc(j)-xl(i)).gt. .001) go to 180
c if(abs(yc(j)-yl(i)) .le. .001) go to 190
180 continue
c go to 210
190 if(j=2) 200,200,230
200 xc(j)=xc(3)
c yc(j)=yc(3)
c go to 230
210 continue
c go to 320
220 if(i.ne.1) go to 230
c xc(2)=xc(3)
c yc(2)=yc(3)
230 npt=2
c go to 240
240 ad=sqrt((xc(1)-xc(2))*(xc(1)-xc(2))+(yc(1)-yc(2))*(yc(1)-yc(2)))
c angle=angle + signal*2.*asin(ad/(2.*r))
c go to 310
c four intersection points (each on a different side).
c determine angle by finding closest 2 intersections to
c farthest corner, calculate angle between, and add angle
c between other two intersections.
250 dist=10000000.
c i1=0
c i2=0
c i3=0
c i4=0
c do 300 j=1,4
c distn=(xl(ifar)-xc(jj))*(xl(ifar)-xc(jj))
c +yl(ifar)-yc(jj))*(yl(ifar)-yc(jj))
c if (distn-dist1) 260,260,270
260 dist2=dist1
c dist1=distn
c i4=i3

```

```

i3=i2
i2=i1
i1=jj
c go to 300
270 if (distn-dist2) 280,280,290
280 dist2=distn
i4=i3
i3=i2
i2=jj
c go to 300
290 i4=i3
i3=jj
300 continue
c calculate angle between 2 closest points to farthest corner.
ad=sqrt((xc(i1)-xc(i2))*(xc(i1)-xc(i2))
+ (yc(i1)-yc(i2))*(yc(i1)-yc(i2)))
angle=2.*asin(ad/(2.*r))
c calculate angle between 2 points farthest from
c farthest corner and add to previous angle.
ad=sqrt((xc(i3)-xc(i4))*(xc(i3)-xc(i4))
+ (yc(i3)-yc(i4))*(yc(i3)-yc(i4)))
angle=2.*asin(ad/(2.*r)) + angle
310 continue
c compute rate of earthquakes in this annular source
c return
c error printout
320 write(16,330)isub,ic,npt,xnot,ynot,(xl(i),yl(i), i=1,4),r,angle,
&(xc(i),yc(i),i=1,4)
330 format(' **** error in subroutine outsid. source no. ',i3,
&' debug values follow.....',/10x,2i5,5/(10x,4f14.6))
340 call exit
end
c *****
c subroutine insid(xnot,ynot,xc,yc,xc(4),yc(4),
common/xareas/xsav(2,50),ysav(2,50),sarea(50)
common/inout/ic,ics,ifar,azimp,ssize,stepo
common/slinds/dls(4),perp(4),sll(4),slsq(4),persq(4)
&,strsq(4),dip(4)
common/slinp/aa(4,50),bb(4,50),cc(4,50),rta(4,50),dist(4,50),
&disq(4,50)
common/xyl/xl(4),yl(4),xr(4),yr(4),xc(4),yc(4)
dimension jsub(3)
c (xmid,ymid)=point inside quadrilateral source region ii
c r=radius of circle (or radius of annular ring of width dr)
c centered at (xmid,ymid)
c ii = index identifying quadrilateral source region
c the area contained in the intersection of the region and annular
c ring is determined by the fraction of the ring within the region.
c area = pangle x r x dr
c where pangle = the angle (in radians) subtended by the
c part of the annular ring contained within the region.
c subroutine insid determines pangle.
c adapted from subroutine by mcguire
data pi,pi2/3.1415927,6.2831853/
data jsub/2,3,1/
rsqr=r
npt=0

```

```

c
angle=0.
do 170 j=1,3
  if(abs(xc(j)-xl(ii)).le..001) go to 170
  if(abs(yc(j)-yl(ii)).gt..001) go to 180
170 continue
  go to 200
180 if(j=2) 190,190,220
  190 xc(j)=xc(3)
  yc(j)=yc(3)
  go to 220
200 continue
210 if(i.ne.1) go to 220
  xc(2)=xc(3)
  yc(2)=yc(3)
220 npt=2
  go to 230
c 2 intersection points; determine azimuths.
230 if (xc(1)-xnot-r) 240,250,240
240 if (xc(1)-xnot-r-0.001) 250,250,570
250 azim=0.0
  go to 300
260 if (xc(1)-xnot+r) 270,280,290
270 if (xc(1)-xnot+r+0.001) 270,280,280
280 azim=pi
  go to 320
290 azim=acos((xc(1)-xnot)/r)
300 if (yc(1)-ynot) 310,320,320
310 azim=pi2 - azim1
320 if (xc(2)-xnot-r) 330,340,330
330 if (xc(2)-xnot-r-0.001) 340,340,570
340 azim=0.0
  go to 390
350 if (xc(2)-xnot+r) 360,370,380
360 if (xc(2)-xnot+r+0.001) 370,370,370
370 azim=pi
  go to 410
380 azim=acos((xc(2)-xnot)/r)
390 if (yc(2)-ynot) 400,410,410
400 azim=pi2 - azim2
410 pangle=azim2-azim1
  if (pangle) 420,570,460
420 if (azim1-azimf) 430,570,440
430 pangle=pi2 +pangle -angle
  go to 560
440 if (azimf-azim2) 430,570,450
450 pangle=pangle-angle
  go to 560
460 if (azim2-azimf) 470,570,480
470 pangle=pi2 -pangle -angle
  go to 560
480 if (azimf-azim1) 470,570,490
490 pangle=pangle-angle
  go to 560
c four intersection points (each on a different side).
c determine angle by finding closest 2 intersections to
c farthest corner, calculate angle between, and add angle
c between other two intersections.
500 dist=1000000.
  11=0
  12=0

```

```

13=0
14=0
do 550 j=1,4
  distn=(xi(ifar)-xc(jj))*xi(ifar)-xc(jj)
  & +yl(ifar)-yc(jj))*yl(ifar)-yc(jj)
  510 distn=dist1
  dist1=distn
  14=13
  13=12
  12=11
  11=10
  go to 550
520 if(distn-dist2) 530,530,540
530 dist2=distn
14=13
13=12
12=11
11=10
go to 550
540 14=13
13=12
12=11
11=10
550 continue
c calculate angle between 2 closest points to farthest corner
ad=sqrt((xc(i1)-xc(i2))*xc(i1)-xc(i2)
& + (yc(i1)-yc(i2))*yc(i1)-yc(i2))
pangle=2.*asin(ad/(2.*r))
c calculate angle between 2 farthest points from
farthest corner and add to previous angle.
ad=sqrt((xc(i3)-xc(i4))*xc(i3)-xc(i4)
& + (yc(i3)-yc(i4))*yc(i3)-yc(i4))
pangle=2.*asin(ad/(2.*r))*pangle
c angle for this radius is now known, calculate risk
560 return
c anarea=pangle*r**2*size
c error printout
570 write (16,880) isub,ifar,npt,xnot,ynot,xi(i),yl(i),i=1,4),r,
&pangle,xc(i),yc(i),i=1,4)
580 format (' ***** error in subroutine inside. source no.',
& i3,' debug values follow.....',/10x,210,10/(10x,2#12.6))
590 call exit
end
c*****
c subroutine spread(reg,jlo,ll,totac,dmid)
c magnitude smoother for fault ruptures:
c totac acceleration occurrences calculated for ruptures of
c magnitude fm(i1) at distance dmid from site
c are "spread" to acceleration levels corresponding
c to magnitudes in the range
c fm(i1)-delta/2 < fm(i1) < fm(i1)+delta/2
c [where fm(i1)=fm(i1)-delta]
c earthquake occurrences follow a Gutenberg-Richter
c magnitude-frequency relationship, calculated from number
c of earthquakes input at 2 successive magnitudes.
c common/magdis/jent,mdis
c common/rfb/rfab(20)
c dimension reg(106)
c common/acc1/qm(25),exdif(101),atah(20,25),aclog(2)
c common/uds/max,marp
c common/quadr/maxq,ac(105),dumlog,aclog(105)
10 continue

if(dmid,eq,-1.) go to 60
do 20 i=2,mdis
  if(dmid.le.rfab(i)) go to 40
20 continue
c distance exceeds table--we should not even be here
write(16,30) dmid
30 format(' in spread, dmid='e12.5' too large-should not be here')
call exit
40 fr=(rfab(i)-dmid)/(rfab(i)-rtab(i-1))
ilim=i
lp=11*1
do 50 lm=1,lp
  aclog(lim)=atah(i,lm)-fr*(atah(i,lm)-atah(i-1,lm))
  50 ilim=2
  60 acdif=aclog(2)-aclog(i)
  if(acdif.ne.0.) go to 90
  jbot=jlo+1
  70 if(aclog(jbot).le.aclog(i)) go to 80
  jbot=jbot+1
  if(jbot.ge.i) go to 70
  80 jup=jbot+1
  reg(jup)=reg(jup)+totac
  return
  90 add=0.
  jbot=jlo+1
  100 if(aclog(jbot).le.aclog(i)) go to 110
  jbot=jbot+1
  if(jbot.ge.i) go to 100
  c we are slower accel level--for lowest mag of range
  110 aclo=exdif(i)
  jup=jbot+1
  120 upl=aclog(jup)
  if(upl.gt.aclog(i)) upl=aclog(2)
  fr=(upl-aclog(i))/acdif
  iachi=100.*fr+1.00001
  quan=(+exdif(iachi)-aclo)*totac
  add=add+quan
  reg(jup)=reg(jup)+quan
  if(aclog(2).eq.upl) go to 130
  aclo=exdif(iachi)
  jup=jup+1
  if(jup.le.maxp) go to 120
  write(16,125)
125 format(' spread routine needs higher acceleration level,'
  &' rewrite BOX or increase scale factor')
  call exit
130 continue
  return
end
c*****
c subroutine outreg,naf,routs,qouts,sd)
c common/tims/ftim(20),jtim(20),ntims,nwt,iprint
c common/rdis/rdis
c dimension qp(140),sol(20),sols(20)
c common/quadr/maxq,ac(105),dumlog,aclog(105)
c dimension reg(106)
c common/ex/prob,ex,exl
c common/uds/max,marp,maxsp
c dimension regs(106),cdf(105),cdf(105)
  if(iprint.le.2) write(16,10) routs,qouts,rdis,sd

```



```

10 format(' site at long '#B.3', lat '#B.3,
&/' shortest dist to fault='#9.3' km'/10x
&'zero attenuation variability'#B'variability in atten, sigma='
& '#5.2/2( ' g.m. occ/yr exc/yr r(events) r(yrs) ')
do 20 l=1,maxsp
20 regs(1)=0
30 sds=4.4*sd
kls=1
do 70 l=1,max
if(reg(l+1).eq.0) go to 70
asav=(aclog(1)+aclog(l+1))/2.
k=1
40 if(aclog(k1)-asav.ge.-sds) go to 50
k=k1+1
if(k1.le.max) go to 40
go to 70
50 kls=k1
u=(aclog(k1)-asav)/sd
call bbgasus(w,g1)
if(k1.eq.1) regs(1)=regs(1)+(1.-g1)*reg(l+1)
60 k=k1+1
wh=(aclog(k1)-asav)/sd
call bbgasus(w,g2)
regs(k1)=regs(k1)+(g1-g2)*reg(l+1)
if(wh.ge.4.4) go to 70
if(k1.ge.maxsp) go to 70
g1=g2
go to 60
70 continue
c we are really not interested in 1st 4 acceleration boxes but used
c them for computation of variability in acceleration--correct so
c box 5 accumulates all accelerations below ac(5)
c ac(5) is the first value printed
80 continue
regs(1)=regs(1)+reg(1)
regsav=reg(5)
regssv=regs(5)
do 90 i=1,4
regs(5)=regs(5)+regs(i)
90 reg(5)=reg(5)+reg(i)
cdf(maxsp)=0.
cdf(maxp)=0.
maxq=maxp
kmax
do 100 i=5,max
cdf(k)=cdf(k+1)+reg(k+1)
100 k=k+1
mm=maxsp-1
k=mm
do 110 i=5,mm
cdf(k)=cdf(k+1)+regs(k+1)
110 k=k+1
if((cdf(5).ne.0.) and. (cdf(5).ne.0.0)) go to 140
if(print.ge.2) write(16,120)
120 format(' no acceleration events accumulated. statistics calc.
skipped')
do 130 ier=1,ntims
sol(ier)=0.
130 sols(ier)=0.
go to 380

```

```

140 continue
c
c if accelerations do not exceed ac(5), the first acceleration
c of interest, skip statistics calculation to avoid dividing by
c zero or upsetting interpolation routines.
c
cnum=cdfs(5)+regs(5)
cnum=cdf(5)+reg(5)
grncnum=1./cnum
nrep=0
idone=0
if(print.eq.3) go to 330
do 300 i=5,maxsp
if(i.gt.max) go to 150
if(cdf(i).eq.0.0) go to 170
if(reg(i).ne.0) go to 160
150 if(reg(i).eq.0.) go to 190
c
c compute return period in number of events and number of years
c return period in events = average number of earthquakes needed
c to produce an acceleration exceeding a
c r(events)=(yearly earthquakes)/(yearly exceedances of a)
c return period in years = average time between earthquakes
c which cause acceleration a to be exceeded
c r(yrs)=1/(yearly exceedances of a)
c
160 if(i.le.max) retev=cnum/cdf(i)
retyr=retev*yrn cnum
c
c compute acceleration which has probability prob of not being
c exceeded during t years. this is given by the value of a
c for which
c prob = exp(-t * number of exceedances of a per year)
c
go to 180
170 retev=99999.9
retyr=99999.9
180 if(cdf(i).eq.0.) go to 190
retev=cnum/cdfs(i)
retyr=retevs/cnums
go to 200
190 retevs=99999.9
retyr=99999.9
if(cdf(i).ne.0.) go to 200
idone=1
200 continue
qp(nrep+1)=ac(i)
if(i.le.max) go to 210
qp(nrep+2)=0
qp(nrep+3)=0
retev=99999.9
retyr=99999.9
go to 220
210 qp(nrep+3)=reg(i)
qp(nrep+3)=cdf(i)
220 if(retev.ge.100000.) retev=99999.9
if(retyr.ge.100000.) retyr=99999.9
qp(nrep+4)=retev
qp(nrep+5)=retyr
qp(nrep+6)=ac(i)

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qp(nrep+7)=regs(i)
qp(nrep+8)=cdfa(i)
if(retevs.ge.100000.) retevs=99999.9
if(retyrs.ge.100000.) retyrs=99999.9
qp(nrep+9)=retevs
qp(nrep+10)=retyrs
if(nrep.ne.130) go to 260
if(naf.le.3) go to 230
write(16,280) qp
go to 240
230 write(16,270) qp
240 continue
nrep=0
250 format(1h1)
if(idone.eq.1) go to 330
go to 300
260 nrep=nrep+10
270 format(1h #8,2,f9.5,2f8.1,f12.2,f9.5,2f8.1)
280 format(1h #8,4,f9.5,2f8.1,f12.4,f9.5,2f8.1)
if(retev.eq.99999.9 .and. retevs.eq.99999.9) go to 310
300 continue
310 if(nrep.eq.0) go to 330
if(naf.le.3) go to 320
write(16,280) (qp(11),11=1,nrep)
go to 330
320 write(16,270) (qp(11),11=1,nrep)
330 if (.iprint.gt.2) go to 350
write(16,340) cnum,cnums,sd
340 format(2i' total yearly events ',#10.5,12x)
&/10x zero attenuation variability'18x,
& 'variability in atten, sigma='f4.2)
c compute acceleration which has probability prob of not being
c exceeded during t years. this is given by the value of a
c for which
c prob = exp(-t * number of exceedances of a per year)
c
350 reg(5)=regsav
regs(5)=regssv
do 360 jv=1,ntims
maxq=max
call linint(reg,ftim(jv),sol(jv))
maxq=maxsp
call linint(regs,ftim(jv),sols(jv))
if(iprint.le.2) write(16,370)prob,sol(jv),jtim(jv),sols(jv),
&jtim(jv)
360 continue
370 format(#10.3' ext prob ='f7.3' for '15' years'19x,
& #7.3' for '14' years')
380 rat=0.0
if(sol(1).ne.0.0) rat=sol(ntims)/sol(1)
ratsd=0
if(sols(1).ne.0) ratsd=sols(ntims)/sols(1)
if(iprint.le.2)write(16,390)jtim(ntims),prob,jtim(1),rat,ratsd
if(iprint.ge.2) write(2) xouts,youts,(sol(jt),jt=1,ntims),
&(sols(jt),jt=1,ntims)
390 format(' ratio'15' yr'f6.3' extreme value to'15' yr val
&='f7.2,f7.2)
return
end

```





