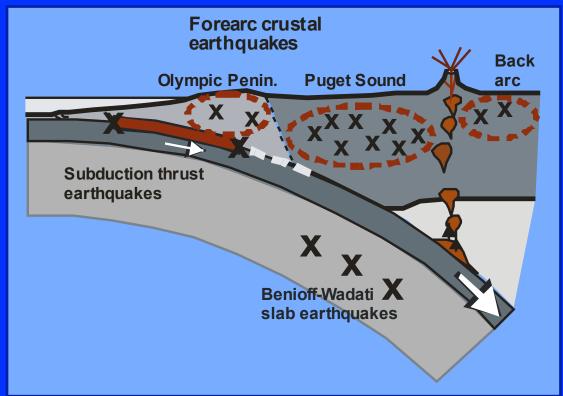
### Seismicity Rates from GPS and other Deformation Rate Estimates in Washington State

Roy Hyndman, Lucinda Leonard & Stephane Mazzotti Geological Survey of Canada, Pacific Geoscience Centre, Sidney, B.C.; SEOS, University of Victoria; Univ. Montpellier, France



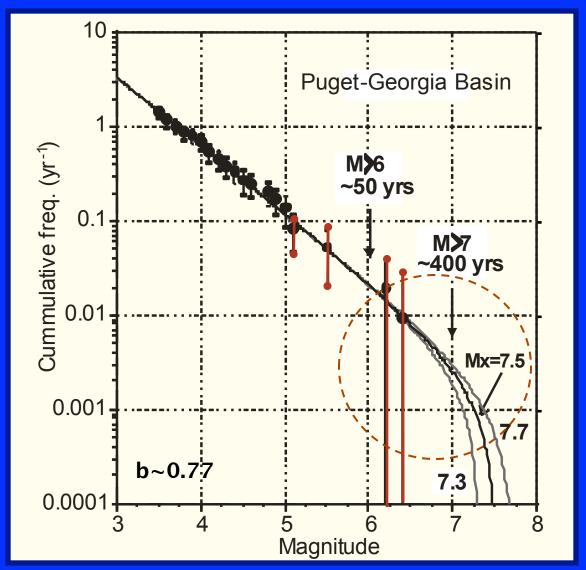




### Catalogue Seismicity Rates Large event rates based on smaller event rates

Are there large events without corresponding smaller magnitude seismicity?
i.e., "characteristic earthquakes"?

Cascadia
megathrust
& central Van. Is.
two M7 with little
other seismicity







## Long-term Crustal Seismicity Rate in Pacific Northwest

- --Can we estimate the earthquake rate from that required to accommodate deformation rate <u>assuming all seismic?</u> (GPS, paleoseismic fault displacement rates, tectonic models, etc.)
  - (1) Puget Sound, (2) Olympic Penin., (3) E. Washington
- --Problem of short duration of catalogue seismicity; possibility of large infrequent earthquakes where few smaller ones, i.e., Cascadia megathrust, M7 central Van. Is. events, etc. "characteristic earthquakes"?
- --also issue of large crustal events in Puget Sound after megathrust as predicted by deformation data





### Seismicity rate from deformation rate

(Kostrov, 1974; Anderson, 1979; Hyndman & Weichert, 1983; Hyndman et al., 2003; Mazzotti et al., 2011)

Convergence Rate:  $s' = C M_o' / (2 \mu A')$ 

where:

A' = WL is the total cross-sectional area (i.e., need seismic thickness)

Mo' = rate of seismic moment release per unit time in area, from recurrence

Mo from magnitude vs moment relation

 $\mu$  = shear modulus

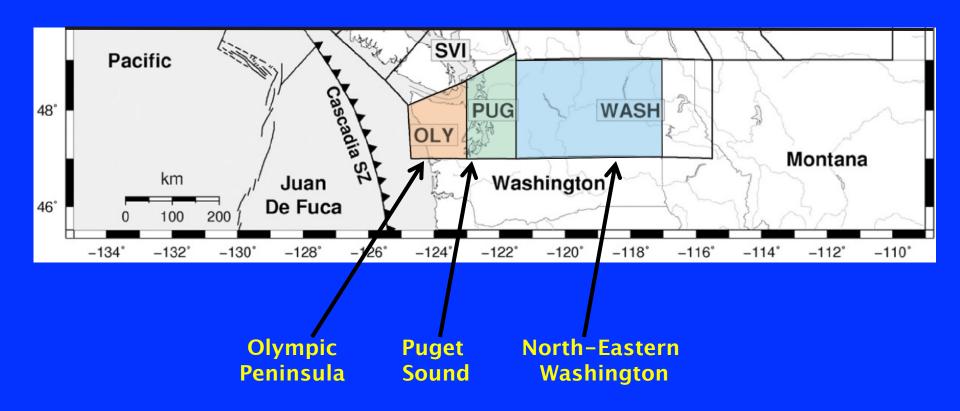
C = faulting orientation constant,  $\sim 1.0$  for 45° faulting

From Hyndman et al. (2003). Better to use strain rates rather than uniaxial shortening but less intuitive, e.g., detailed discussion by Mazzotti et al. (2011)





# Washington Seismicity Areas Studied (mainly from Mazzotti et al., 2011; study with BC Hydro for mainly BC)



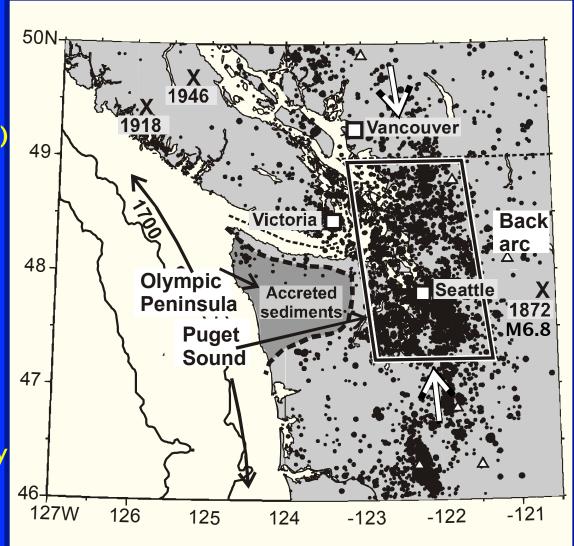




### **Crustal Seismicity Study Areas**

Comparison of long-term seismicity rate from catalogue seismicity and from deformation (GPS etc)

- 1. Puget Sound
  High seismicity, high
  deformation rate
- 2. Olympic Peninsula Low seismicity, high deformation rate
- 3. <u>Eastern Washington</u>
  Low seismicity, moderately
  low deformation rate





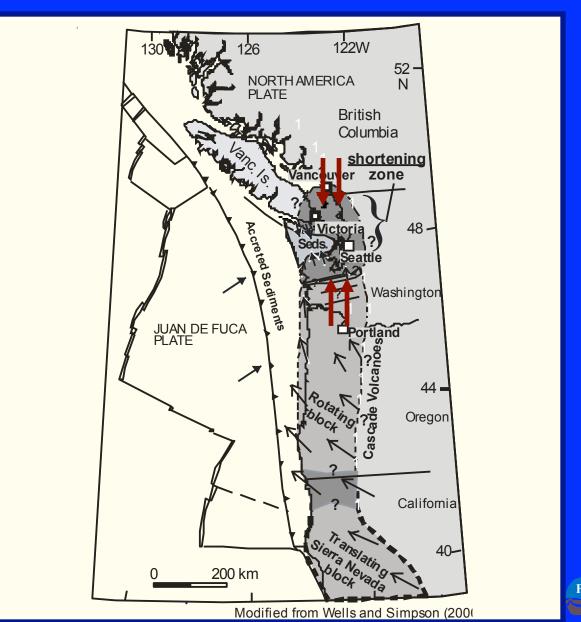


### Puget Sound & Olympic Penin. N-S Deformation

Oregon forearc block is rotating northward compressing Puget Sound against Vancouver Island backstop (Wells &Simpson, McCaffrey, and others)

Weak Olympic accr. sediments focus forearc deformation in Puget area





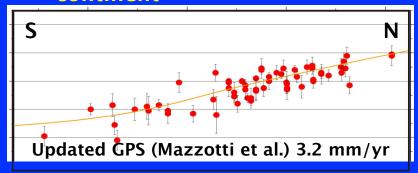


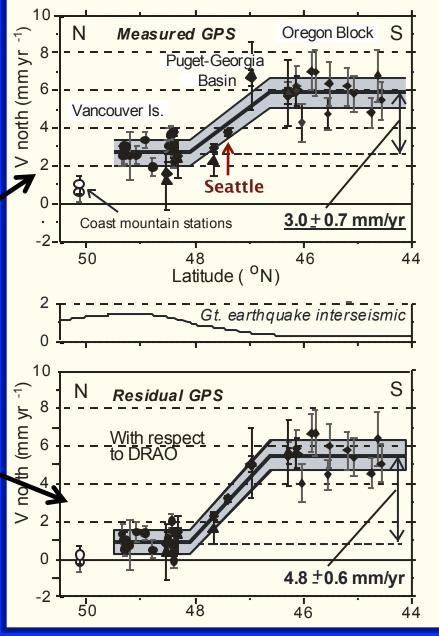
### Puget Sound N-S Deformation Rate from GPS

**Stable continent reference** 

Current rate relative to Vancouver Island (which is moving north due to great earthquake cycle elastic deformation)

Long-term through great earthquake cycle i.e., relative to stable continent







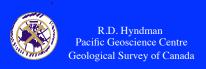
### N-S Shortening Rates for Puget Sound Area

1. GPS shortening rate (Sherrod, Mazzotti, Haugerud, 2008; Hyndman et al., 2003):

Current:  $3.2 \pm 0.8 \text{ mm/yr}$ : Long-term:  $4.8 \pm 0.6 \text{ mm/yr}$ 

- 2. GPS rotation models for Oregon Block (e.g., McCaffrey et al., 2007): Puget Sound northerly motion:  $4.4 \pm 0.3$  mm/yr.
- 3. Paleomagnetic models (e.g., Wells et al., 1998; Wells and Simpson, 2001): Northerly motion of southern Puget Sound: ~6 mm/yr
- 4. Active fault estimates: total north-south rate: 3.6 mm/yr (Sherrod, Mazzotti, & Haugerud, 2008; Wells and Simpson, 2001); Seattle fault: 0.7-1.1 mm/yr, Johnson et al., 1999)

What seismicity rate is required to accommodate this deformation rate (assuming all seismic)?



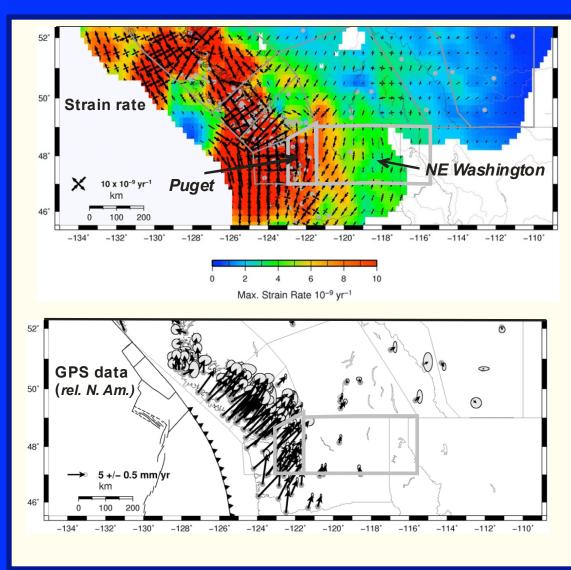


### **GPS Data and Strain Map of Pacific Northwest**

(Mazzotti et al., 2011)

Estimate smoothed strain field (not just shortening)

Estimate seismicity rate required to accommodate strain



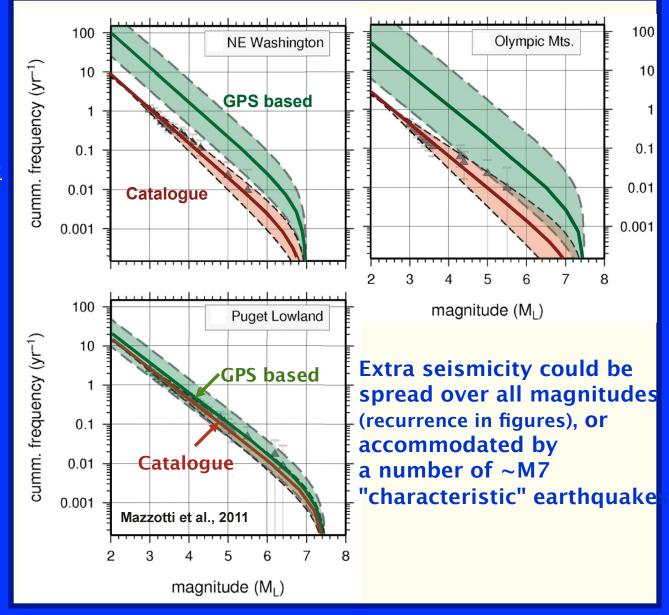




Seismicity
Rate
from GPS
strain rate
(Mazzotti et al., 2011):

Puget-Sound agrees well with catalogue

E. Washington & Olympic Penin. much greater than catalogue (x10)





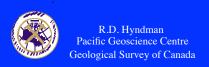


# Moment Rates from Catalogue Seismicity compared to from GPS Strain

(Mazzotti et al., 2011)

 $(10^{17} \text{ Nm yr}^{-1})$ 

Puget Sound	<u>iPS</u>	<u>Seismicity</u>	<u>Ratio</u>
	2.8	2.2	~ 1
Olympic Mtns.	4.1	0.23	18
E. Washington	1.9	0.20	10





### Olympic Peninsula and Central/Eastern Washington Seismicity Rates from GPS Data

- 1. Catalogue seismicity for <u>Puget Sound</u> agrees well with estimate from deformation
- 2. Large N-S shortening rate for <u>Olympic Peninsula</u> (greater than for Puget Sound) but very little seismicity. Is deformation mainly aseismic in this accretionary sedimentary prism (as in such prisms elsewhere)?
- 3. Seismicity for <u>Eastern Washington</u> required to accommodate GPS strain is much greater than that in the catalogue. This is a common occurrence for low deformation rate areas. Explanation unclear; aseismic?

But, more frequent large events than indicated by the catalogue cannot be excluded. ... low weighting?





### Good agreement of deformation estimates and catalogue seismicity:

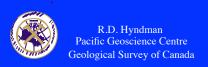
- 1. Puget Sound N-S shortening
- 2. San Andreas fault system
- 3. Queen Charlotte fault zone
- 4. Denali fault zone, Alaska
- 5. Offshore transform faults of J. de F. ridge
- 6. Charlevoix (eastern Canada seismic area)

### **Deformation estimates much greater:**

Olympic Mtns, E. Washington, most of B.C. interior, other slow deformation examples; <u>possible</u> large "characteristic" events not shown by catalogue

### Large recorded earthquakes with few smaller events:

Cascadia, Central Vancouver Island...





# Implications for PSHA: ratios of ground shaking derived from GPS vs seismic data

(Mazzotti et al. (2011)

Puget Sound: ~1:1 good agreement

Olympic Peninsula and E. Washington:

> 10 x seismicity; 2-3 times greater PGA from GPSderived model compared to catalogue seismicity





### **Conclusions:**

(1) Can long-term seismicity rates be estimated quantitatively from deformation data? (GPS, fault slip rates, tectonic models, etc.)

Yes, but... probably only for areas with rapid deformation and well-organized "mature" faults or deformation zones (Hypothesis?)

(2) Deformation confirms seismicity rates from catalogue for Puget Sound

--GPS deformation larger than seismic for Olympic Mtns. and E. Washington, so could have large "characteristic" events, but give low weighting?





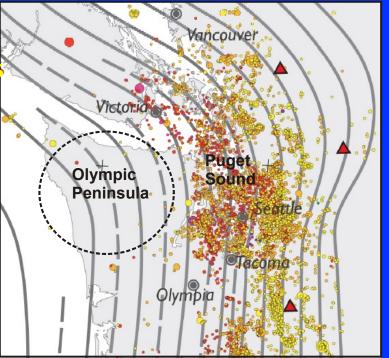


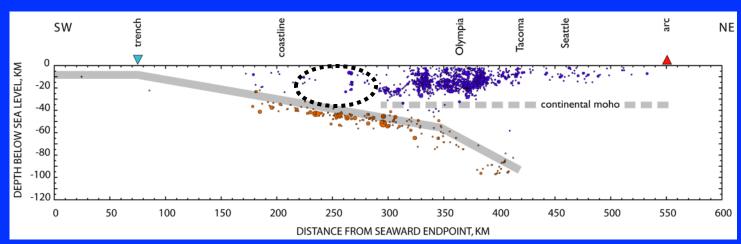






#### Very few catalogue earthquakes in Olympic Peninsula area









### Puget Sound Seismicity Rates from Deformation Data

- 1. Seismicity rate to accommodate deformation is very close to rate from catalogue seismicity.
- 2. Vancouver Island receding to north due to locked megathrust elastic deformation. Will be recovered to south in megathrust. i.e., ~up to seven M > 7 events after megathrust possible.

Crustal events occurring some "short" time after megathrust could be more important hazard than the megathrust itself,

-but not really apparent in paleoseismic data?

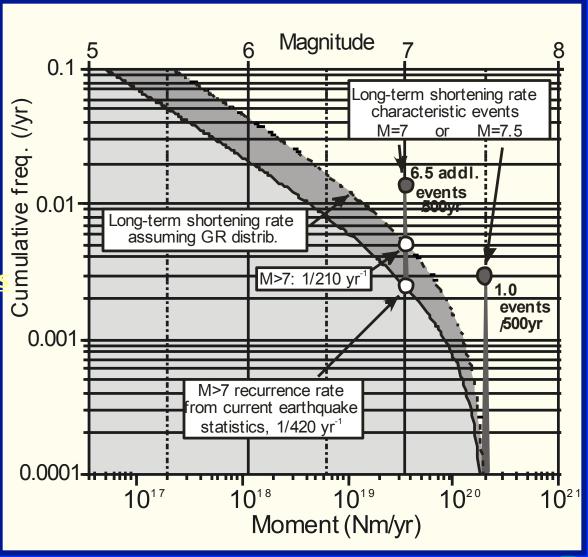
Sherrod et al. conclude about half of Holocene shortening occurred over short time period (~1100 yrs ago), including Seattle fault earthquake M≥7 (1020–1050 cal yr BP, Atwater 1999). Timing does not appear to match a megathrust event.

Pacific Geoscience Centre Geological Survey of Canada

### Long-term Deformation over Great Earthquake Cycle

South component of megathrust coseismic motion results in 0.5-1.0 m short-term N-S shortening

~ 7 additional M=7 events following the megathrust possible to accommodate this shortening?







#### **N-S GPS Puget Sound**

Long-term through great earthquake cycle i.e., relative to stable continent

Current rate relative to Vancouver Island (which is moving north due to great earthquake cycle elastic deformation)



